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A STUDY OF WATER USE EFFICIENCY
IN FIELD CROPS OF MAISE AND HEARS

by

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Declaration

I declare that this thesis has not been submitted for a degree in any other University.

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SUMMARY

crop water use efficiency may be defined as the ratio of total dry matter produced to total evaporation from the crop and the soil. In East Africa, the areas of high photosynthetic potential (highest number of sunshine hours) are also the driest. Success in growing annual food crops like maise and beans in these areas is therefore heavily dependent on the date of planting and the ability of the crop to complete all its stages of growth within the short rainfall seasons.

While breeding of short term and high yielding varieties has been emphasized, little has been done in finding out the water use patterns of the varieties.

information on the water use patterns of one hybrid of maise (H 511) with a medium maturity period (4 months) and one popular variety of field bean (Phaseolus vulgaris var. Canadian Wonder), in all stages of growth, and to observe the effect of reduced soil moisture on the water use and rate of growth of these two crops. Ratios of crop water use (Et) to Penman estimate of open water evaporation (Eo) gave values as high as

1.4 - 1.5 for maise and 1.3 - 1.4 for beans under wet conditions. It is shown that the excess water use, at least in the maise crop, may be due to the combination of large net interseption of rain and low aerodynamic resistance. There was a reduction of \sim 20% in water use and \sim 40% in yield in the dry treatment of the maise experiment.

Measurement of soil moisture in situ by the neutron scattering technique was studied with the intention of using the method for routine determination of crop water use in the field. Although reliable calibrations were obtained for two makes of neutron moisture meters. E.A.L. and N.I.V.-I, the method was shown to be successful only in the absence of drainage. Because of errors in calibration and spatial variations in moisture contents, the precision of soil moisture determination by the neutron moisture meter is not adequate for small differences and the interval between measurements should be at least 7-10 days. In irrigated fields, the inherent poor distribution of irrigation water is a major limitation. Attempts to derive drainage correction data from tensiometer readings were not successful.

been successfully correlated with measured dry matter production in maise and beans. The correlations suggest that in the local environment respiration loss for the two crops is a constant proportion of gross photosynthesis in all stages of growth. This result enables the prediction of the maximum yields of these crops from meteorological data, mainly solar radiation.

Studies of partition of energy in field crops of maise and beans have shown that in the local environment, when the crops are supplied with adequate water, all net radiation may be converted into latent heat, and for periods of 1-2 hours during the day, latent heat may greatly exceed net radiation, the extra energy being derived from the air.

Finally the implications of the above findings on plant breeding and agronomic techniques for maximum water use efficiency are discussed.

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CHAPTER 1

INTRODUCTION

The Climate of Bast Africa

The three territories which comprise East Africa lie between latitudes 5° North and 12° South. The climate of Bast Africa is therefore tropical with 11-12 hours of daylight throughout the year. However, because of the tremendous variation in altitude, air temperatures and average rainfall vary over wide limits from place to place. Mean air temperatures decrease from 25-28° C.at the coast to 13-15° C at 2,500 metres and less than 0° C on top of mountains rising above 5,000 metres. Except for the coast and the area around lake Victoria where total rainfall is higher, average rainfall in East Africa increases with altitude and is bimodel, falling in two distinct seasons with periods of varying length and degree of drought in between. Uganda is much wetter than Kenya and Tansania, but as a region, only about 5 per cent of the East African land surface has average

rainfall greater than the potential open water evaporation in four years cut of five (Dagg, Woodhead and Rijks, 1970).

over most of East Africa, therefore, one of the major factors limiting agricultural productivity is inadequate soil moisture during critical stages of plant growth. With valuable perennial crops like coffee and tea, supplementary irrigation during the dry season has proved economical but for annual crops of less economic value, the answer seems to be in matching the crops to rainfall regimes and the breeding of drought resistant and drought escaping varieties.

Heed for research in crop water use

Maise, Zea mays, is one of the most important food orops in East Africa. This is obvious from the tremend-ous efforts made by local people, particularly in Kenya and Tansania, who insist on growing maise even under very unfavourable environmental conditions. This results in unnecessary wastage of seed and labour. The logical solution to this problem may be the intensification of

maise production in high potential areas, but for a variety of reasons, mostly connected with marketing, it seems that this will take some time to operate successfully. Farmers in the low potential areas will therefore continue to grow maise and the best that can be done to help them at present is to provide them with high yielding varieties and devise agricultural methods which would minimise input costs and changes of crop failure.

There is another important reason why the problem of growing maise - and other grain crops - in the low potential areas of East Africa should be given sufficient attention. The rangelands of East Africa are ideal for animal production, particularly beef cattle; but one of the problems which must be solved before these areas can be fully utilised is how the animals can be supplied with sufficient feed during the dry seasons. Cereals such as maise, sorghum and millet would fulfil this need provided they can be grown under such conditions of low and erratic rainfall.

Rational approach to matching crops to anvisoment

Unlike the sorghums and millets which have over the years evolved several varieties naturally adapted to the different local climatic regimes, maise is not indigenous to East Africa and was only introduced in the 19th century. However, because of its greater resistance to bird damage, maise has achieved such popularity that it has become necessary to produce seed suited to different rainfall regimes.

Significant progress has been achieved in breeding high yielding and drought escaping varieties of maine for areas with short, medium and long rainfall seasons. Much breeding work has also been carried out with sorghum and millet which are more drought resistant than maise.

In his paper entitled "A rational approach to the selection of crops for areas of marginal rainfall in East Africa", Dagg (1905) showed how success or failure in growing a crop of maise requiring 210 days to reach maturity at Muguga was governed not by the heavy long rains concentrated in the months of April and May, but by the small amounts of rain received in June, July and

August. The validity of the above prediction is borne out by the frequent failure of the local long maturity varieties of maise and the reliable performance of the Katumani short-term (3 months) variety. This approach would therefore be very helpful in land-use planning provided there are sufficient data on rainfall reliability, crop water use in all stages of growth and soil moisture storage capacity. Significant progress has been achieved in calculating rainfall variability (Manning 1965; Walker and Rijka 1967; Huxley, Turk and Mitchell, 1969), and where soil data are not available. it is a relatively simple job to take soil samples and work out soil moisture storage capacity. Few data are however available on the water use of crops in all stages of growth. It is therefore one of the main objectives in this thesis to examine the water use patterms obtained for maise in different experiments and to find out what additional information, if any, is required for the valid application of these results.

Water use of maise

Numerous studies have been carried out on the water use of maise (Zea mays). Results from these

experiments have not, however, always been comparable because of varying and often unspecified experimental conditions, e.g. fertilizers applied, plant population and climatic conditions.

Haynes (1948) concluded from his experiments that available soil moisture affected vegetative growth, but did not affect transpiration rate per unit of plant dry weight. This result suggests that transpiration rate is reduced in the same proportion as the loss of dry matter production due to water stress. From their studies on supplementary irrigation for maise, Letey and Peters (1957) concluded that while maintenance of soil moisture tension well above 15 atmospheres tension was desirable, adequate water reserve in the soil profile at planting was also important. Supplementary irrigation for deep rooted crops was then required only when weather conditions favoured a serious depletion of water in the upper two feet of soil.

Denmend and Shaw (1962), however, showed that when potential transpiration was about 6-8 mm/day and soil water potential was greater than atmosphere, maise was unable to maintain either full transpiration rate or full turgor. Fuehring et al. (1966) also showed that

increasing soil moisture stress within the upper half of "available water" decreased transpiration. In their experiments, crop water use in the first week after irrigation was greater than in the second week, and yields were depressed by 4% when a weekly irrigation interval was increased to two weeks.

Evaporation from bare soil is also an important factor in the water use of a maise crop. Peters and Russell (1959) found that in a crop of maise evaporation from bare soil accounted for 50-70% of total water use. They used polyethylene plastic covers to separate transpiration from bare soil evaporation in field crops of maise. Similar results were obtained by Harrold et al. (1959) in lysimeters covered with plastic. In studies on the influence of soil moisture. nitrogen fertilisation and plant density on evapotranspiration and yield of maise, Carlson et al. (1959) also found that because of the effect of surface wetness on evaporation from bare soil, evapotranspiration from irrigated plots was considerably greater than water use in the non-irrigated plots. Bare soil evaporation also interferes with meteorological estimates of crop water use. Garber and
Decker (1961) compared water balance from a ten-acre
maize field with heat budget estimate of evapotranspiration by the method of Penman (1956). The two
estimates were in agreement when the soil surface was
wet, but not when the soil surface was dry even though there
was sufficient water available in the soil profile. These
workers suggested that under dry surface conditions, surface temperature and hence sensible heat were
underestimable.

In spite of the uncertainties in the experiments described above there seems to be a general pattern of water use for maise which shows a gradual increase in the early part of the crop, a plateau of varying duration after tasselling and a decline as the crop matures. Such a pattern is shown by Denmead and Shaw (1959) in their analysis of water use data covering 5 seasons at 11 sites in Iowa. Most of these data, based on State Soil Moisture Survey, were necessarily of low accuracy, and assumptions on runoff and deep drainage may be in error. The resulting pattern of the ratio of water use (Et) to

pan evaporation (E_p) should, however, be representative. Et/E_p ratios increased in sigmoid namer from a value of 0.36 at planting to C.81 at silking. The value 0.81 remained constant for 16 days and then decreased, apparently due to declining physiological activity of the crop.

a maximum during pollination, have been observed by

Pritechen and Shaw (1961), England (1965), and in studies

conducted over two seasons by Cackett and Metelerkamp

(1964). In the latter study, water use was estimated

from soil moisture sampling, but excluded 5 days after

each irrigation to minimise errors due to drainage.

The recorded maximum ratio of Et to Penman Bo was nevertheless 1.10. This figure would be an underestimate if

there was water extraction by maise roots beyond the

48 inches depth of soil profile sampled. These data

were also used to formulate seasonal trends in Et/Bo

as quadratic functions of the age of the crop in

weeks (r) in equations of the form

$$\frac{Et}{Eo} = ax - bx^2 - c.$$

It is, however, difficult to find any physical basis

for such an equation, and hence its usefulness in predicting water use of maise in a different environment is limited.

Tater use of beans (Phaseolus vulgaris)

Field beans are a popular and common food crop in East Africa. The seed has a high protein content and is a useful supplement to the maise diet. Research work on field beans has, however, lagged far behind that on maise and relatively little is known about the water requirements of beans.

Cackett and Metelerkamp (1963) working in Sabi
Valley, Rhodesia, and using a technique similar to that
used for the maise crop as discussed earlier, found that
the water use pattern for variety Red Canadian Wonder
beans was similar to that of maise, except for a much
flatter peak; maximum consumption occurred during the
period from 9 to 12 weeks after planting. The highest
Bt/Bo was approximately 0.90, average for the season
was 0.72. These data are subject to the same limit—
ations observed for the maise data.

estimate of crop water use

water use by crops is primarily an energy dependent process. The rate of vapour transfer away from the site of evaporation, and in some cases, the rate of supply of liquid water to such sites, serve to modify the energy dependence of evaporation from the soil and leaf curiace.

Since the pioneering work of Dalton (1834), Rohwer (1931), Briggs and Shants (1914-1917), and others, good progress has been made in the understanding of the evaporation process, especially evaporation at an open water surface.

The methods of Thornthwaite (1948), Penman (1948), Blaney and Griddle (1950), and Olivier (1961) have all been tried in East Africa, but Penman's formula has provided the best correlation with data from pan evaporimeters in the tropical climate of East Africa.

The basic Penman equation combines energy balance with the efficiency of vapour transfer (sink strength) and can be written as

$$B = \frac{\Delta Rn + \delta Ea}{\Delta + \delta} \qquad \dots (111)$$

where B = evaporation rate (mm/day),

Rn = net radiation (equiv. mm water/day),

 \triangle = slope of saturation vapour pressuretemperature curve (mb $^{\circ}0^{-1}$).

or = psychrometric constant (mb °C-1), and Ea is the equivalent of

0.35 (es - ed)(1 +
$$\frac{u}{100}$$
)

with as replaced by ea, the saturation vapour pressure at the air temperature.

u is windepeed (miles/day).

With the above approach, it is possible to estimate with an error less than 20 per cent for periods
longer than 10 days, the evaporative loss from a free
water surface through a relatively simple integration
of simple meteorological observations (McCulloch, 1965).
Maps of monthly and annual potential open water evaporation based on Penman's formula have been prepared for
Kenya and Tansania by Woodhead (1968) and for Uganda by
Rijas and Owen (1965). keliable estimates of open
water evaporation can also be made for periods less
than a day, but proper measurement and integration of
meteorological parameters become much more complex.

Application of Pennan's formula to evaporating surfaces, other than open water, has met with diffigulties. Penman (1948) found that his meteorological estimates of open water evaporation Bo were well correlated with evaporation from well watered, short homogeneous grass at Rothamsted. The normal field crop is, however, neither short nor well watered, nor does it completely cover the ground. Fortunately there is general agreement on the shape of the evaporationtime curve for bare soll surface drying from initial thorough wetting (Penman 1941; Philip, 1956; Veihmeyer and Brooks, 1954). These curves are based largely on experimental data and although the role of soil capillary conductivity in evaporation from bare soil has been recognised and attempts have been made to include this factor as an additional resistance in the evaporation or transpiration process (van den Honert, 1948; Cowan, 1965; Wangati, 1966), it has not been possible to find representative values of this factor for incorporation in the evaporation formula.

Perman and Schoffeld (1951) estimated the ratio of evaporation from plant cover to that of open water by

Perman and Schofield (1951) estimated the ratio of evaporation from plant cover to that of open water by assuming that the crop canopy could be regarded as part of a large flat leaf, with the stomata fully open during the day and fully closed at night. The effective "length" over each cm² of surface, a parameter defining the efficiency of vapour transfer by turbulent mixing, was therefore assumed to be the same for open water as for continuous crop cover and had the same effect on transpiration as well as carbon dioxide assimilation. The additional stomatal resistance was therefore calculated on the basis of stomatal dimensions and their population.

The final equations arrived at predicted that
evaporation from well watered turf was less than open
water evaporation. This prediction was supported by
experimental data and energy balance estimates, and the
influence of day length on Bt/Eo indicated that where
the above assumptions applied Et/Eo was equal to a
constant, f, which varied between 0.6 and 0.8 according
to seasons (Penman, 1956). Penman calculated irrigation
requirements in Britain on the basis of the above Et/Eo
ratios. The approach was successfully applied by Pereira
(1957) in estimating requirements for supplementary

wallis (1903) found the same procedure applicable in predicting soil moisture deficits under irrigated coffee. Mitchell (1965) however found that coffee fields in northern Tansania could vary in their irrigation requirements, suggesting that there was no simple formula for estimating the frequency and quantity of irrigation water. Evaporation data estimated from catchment area water balance in East Africa (Blackie, 1964; Dagg and Blackie, 1965) suggest that in high rainfall areas where evaporation is not limited by soil moisture deficits, Et/Eo values for perennial vegetation like high montane forest and tea plantation at full cover are relatively constant, varying between 0.7 and 0.9.

Complications arise, however, when the fraction of ground covered by the crop changes with crop development and where available soil moisture is not sufficient to meet potential evaporative demands on the crop. The resulting changes in Et/No-time curves can be very large but there is no method as yet which can be used to make the necessary quantitative allowance in the factor f.

There is a further problem, this time arising from Et/Eo

from 1.0 to 1.8 were reported by Prescott (1938),
Stanhill (1958), Mather (1954), McCloud and Dunavin
(1954), and Hutchison, Manning and Parbrother (1958).
Although there is some doubt on the procedures used
for estimating open water evaporation, Bo, it is
evident that some crops can be more efficient than open
water surfaces in the conversion of net energy into
latent heat.

constants in the Bo formula (Sibbons, 1962), indicates that Et/Eo ratios greater than 1.0, and the obvious differences between crops, might be more related to the physical properties of the crop canopy, especially the aerodynamic resistance referred to earlier. The derivation of Penman formula in a somewhat different way by Monteith (1965) emphasized the role of the aerodynamic resistance (r) and the additional stomatal resistance (r) introduced in a crop when the leaf surfaces are not wet. Measurement of aerodynamic resistance for crop canopies has, however, proved difficult, requiring large representative areas for proper determination of wind-

speed profiles. Numerous measurements with diffusion porometers are also required for determination of r.

However, in spite of the difficulties cited above, this approach provides the best method to date of defining expressions for crop water use in terms of crop parameters and sould be extended in the future to include other factors connected with soil moisture availability.

In conclusion to the above review of present status on the application of evaporation formulae to crop water use. it is perhaps a fair observation that interpretations of many results of crop water use measurements have been biased by the underestimation of the part played by the roughness of crop canopies in the extraction of sensible heat from the air, and the conversion of this energy into latent heat. Values of Et/Eo which turned out to be greater than unity were therefore in some cases too readily explained in terms of advection of heat from surrounding areas, with a suggestion that actual water use in the middle of very large, uniform field erop would always be less than potential evaporation from open water in the same environment. Even after the difference in r between crop and water surfaces was recognised, it was still maintained that the probable increase in Bt would at most just balance the decrease due

for a resident period, band in the barry, to us area of

to surface albedo (0.25 for crop, 0.05 for water).

Advection is, however, a factor to be considered seriously in designing experiments for crop water use measurements. The experiments of Fritschen and

Van Bavel (1964), with lysimeter crops of Sudan grass protruding well above the surrounding crop, have provided a good demonstration of the importance of a uniform crop in water use experiments.

Other methods of estimating crop water use - neutron moisture meters

Direct measurement of soil moisture under field crops is not only necessary for checking the validity of meteorological estimates of crop water use, but also such measurements provide useful information on the soil moisture profile at all stages of crop growth. These data should therefore lead to a better understanding of root activity and hence the optimum frequency of irrigation and depth of placement of fertilisers for maximum uptake by the plants.

ment requires soil sampling in the field, weighing, drying for a standard period, usually 48 hours, in an oven at 105°C, followed by cooling in a desiccator and re-weighing. This technique is tedious because heterogeneity of soil moisture conditions in the field requires large numbers of samples for a reasonable degree of accuracy to be achieved. Soil sampling also tends to destroy the soil profile to the extent that in time, sampling sites become channels of preferential drainage. A further source of error in this method is the necessity to determine soil bulk density or extraction of core samples of known volume for the conversion of gravimetric to volumetric moisture contents.

Specially made mylon resistance units can be used to measure soil moisture. Calibration of these units is, however, difficult and is usually not possible to check either during or after lengthy field experiments. This limitation applies especially to the cheaper gypsum blocks available commercially.

Considerable and fast development of an alternative method of measuring soil moisture content in gitu by the neutron scattering technique has taken place in the last 18 years. Gardner and Kirkham (1952) considered that hydrogen present in the soil, mainly as water, was the

They developed a theory and a method of measuring soil moisture content based on this property and tests indicated that the method was applicable in the range of soil moisture contents between oven-dryness and water saturation. Holmes (1955) presented results of similar tests carried out in Australia and among these early experiments the work of van Bavel, Hood and Underwood (1954) was directed on methods of increasing vertical resolution of the equipment.

Most of the early work on the use of neutron and gamma radiation for this purpose is extensively reviewed in the Commonwealth Bureau of Soils Bibliography, "The Determination of Soil Moisture Using Neutron Probes (1963-1951)". The later developments are mainly in instrument design for both better resolution and reliability in field operations (Bell and McGulloch, 1966), methods of minimising errors in calibration and measurements (Bell and Beles, 1967), and operational precautions for accourate evaporation measurement (Van Bavel and Stirk, 1967).

As a result of successful calibration and field tests of the neutron moisture meters, a meeting of a panel of experts on radiation techniques in soil physics and irrigation studies, held in Vienna in October 1964, considered that the neutron moisture meter represented a considerable advance in technique over any previous soil moisture measuring device, and held promise of being able to yield useful quantitative measurements of crop water use.

Accordingly, it was decided to set up a co-ordinated experiment in which the water balances of the same crop under irrigation in different countries would be compared. The crop water use measurements with neutron moisture meters reported here were part of this co-ordinated experiment.

Dry matter production in relation to crop water use

Weight of dry matter produced per unit of water lost by evaporation in a unit area of field crop (Haise and Viets, 1957). Since in most cases only a fraction of total dry matter is of economic interest in agriculture, a more practical definition would be in terms of market-

able crop produced for a unit depth of water used in evapotranspiration. Both evaporation and dry matter production are dependent on energy derived from solar radiation. Crop water use efficiency can therefore also be defined in terms of the partition of net energy between latent heat and chemical energy stored in the form of dry matter. Several accounts of this approach have appeared in literature: Allen. Youum and Lemon (1964) obtained photosynthetic efficiency of 6.85 in the utilisation of radiation in the 0.4-0.7 wavelength range. Using Beer's Law for light absorption, they calculated potential photosynthesis and compared this value with potential evapotranspiration Et calculated from net radiation data. The result was that 16 times more energy was used in evaporation than in photosynthesis.

Amongst the more interesting data, Yao and Shaw (1964) found that water use efficiency in maise was also sensitive to plant population, being highest (571 lbs/inch) in maise planted two seeds per hill in rows 21 inches apart, and falling to 414 lbs/inch when inter-row spaces were increased to 42 inches. The same workers (1964b) found that net radiation 1 metre above

the crop was higher in 42 inch row crops than in 21 inch row crops, implying greater storage of radiant energy in the denser crop.

The subject of energy conversion in photosynthesis and evaporation is reviewed by Lemon (1966). It turns out that the better the understanding of the physics of energy exchange and the morphology and physiology of plants, the less can be said about water use efficioncy in such general terms as the "transpiration ratio". Each crop, on each site at a particular time, follows the laws governing the interaction crop and environmental factor. The study of crop water use efficiency described here is therefore not meant to provide more data on the weight of dry matter produced per unit of water used, but to find logical and convenient means of estimating both optimum crop productivity and crop water requirements under field conditions from measurable environmental and crop factors. The task of combining crop productivity with crop water use for maximum water use efficiency will be left to the agronomist to work out according to carcumstances environmental social and economic - prevailing at each location.

Evanoration reduction in field erand

While a better understanding of the pattern of crop water use can be used to minimise westage of water in irrigation and would, of course, be very useful in the extension of erop agriculture to drier areas, the scope for control of Bt is limited by the necessity to keep stomata open as long as possible for 00, assimilation. Studies carried out on chemical anti-transpirants are reviewed by Waggoner (1966). It appears that although there are chemical sprays which would control stomatal opening and in a few cases these have been found to decrease transpiration relatively more than photosynthesis, little is known about the feasibility of using these techniques under field conditions. Experience will therefore indicate that at present, the most promising approach to evaporation reduction in field crops lies in finding out the crop water requirement at all stages of growth and the most economic methods of satisfying this requirement.

Estimates of cross photograthesis in field gross

Scientific studies of crop yields date as far back as the earliest fertiliser trials in Britain and America in the 19th century. These studies were confined to final yields of the economically important components, vis. grain, tubers or hay. Interest in the relationships between plant populations and yields subsequently developed, followed by studies of growth rates, which led to the introduction of Leaf Area Index by Watson (1947). Parallel studies of carbon dioxide assimilation in leaves were being carried out in controlled environments, pioneered by the work of Blackman (1895). Brown and Escombe (1905), Maskell (1928), and following experiments by Heath (1951) and Geastra (1959) and others. there has been a welcome but gradual extension of laboratory methods to field crops by workers such as Hesketh and Moss (1965).

Carbon dioxide concentration and light intensity are the two major limiting factors to photosynthesis under field conditions. There is little one can do to increase either the 00_2 concentration in the free

being dependent on windspeed and serodynamic properties of the crop. Much field workshas therefore been directed at the description of the pattern of light interception in the crop canopy and attempts to calculate gross photosynthesis from these profiles using photosynthesis-light curves for single leaves as determined in the laboratory.

Some of the original work (Davidson and Philip, 1958; Sacki, 1960) assumed light profiles based on Beers' Law of light extinction in a homogeneous medium:

$$I_{\chi} = I_0^{\bullet} \qquad \dots \dots (1:2)$$

where I is the incident light intensity

I is the light intensity at depth L in the medium k is the light extinction coefficient in the medium.

In the case of erop canopies, k is dependent on transmission through single leaves and on their geometric arrangement on the plant. The changing spectral composition of radiation due to absorption, transmission and reflection by leaves in the canopy was neglected.

de Wit (1959) based his calculations on a model assuming random distribution of leaf angles within a canopy which

absorbed all radiation falling on it. This approach was modified later (de Wit, 1965) by dividing up the canopy into an infinite number of sections oriented at all angles to direct solar radiation, with a secondary radiation component based on contribution from diffuse sky radiation. The total area of canopy receiving light was obtained by a series of integrations based on the above factors. Success of de Wit's method has been demonstrated for complete camppies of alfalfa (Stanbill. 1962), pasture (Alberda and Sibma, 1962), kale (Watson and Witte, 1959), subterranean clover (Davidson and Donald, 1958; Black, 1963) and rice (Takeda, 1961), but serious discrepancies have also been observed with corn (Williams, Loomis and Lepley, 1965), mixed pasture (Brougham, 1956) and sugar beets (Watson, 1958).

Monteith (1965) eriticised the use of mean light intensities in estimates of photosynthetic rates in erop canopies and the neglect of the changing spectral composition of light as a result of reflection and transmission in the canopy. In order to minimise such errors, Monteith, in the same paper, proposed a new model for light interception in field crops. In this model, the

canopy is divided into layers of unit leaf area index, L, and the light distribution is given by the binomial expansion of the equation

$$I_{I_0} = I_{O}(a + (1 - a)\tau)^{L}$$
(1:5)

where s is a parameter characteristic of the average arrangement and crientation of leaves, and is equal to the fraction of incident radiation that passes through a layer without interception. T is the average transmission coefficient of the leaves over the spectral range of the incident solar radiation.

The model has been criticised on the basis that it stops at fractional areas of sunflecks which are not as accurate in predicting photosynthesis as are the fractional areas of leaves in each layer receiving light at various angles of inclination to their surfaces.

other models, e.g. Duncan et al. (1967), have been proposed, but the main differences are in the degree of detail the authors consider necessary for proper evaluation of the light and carbon dioxide functions and effects on overall estimates of gross photosynthesis.

The work described here placed emphasis on Monteith's

model as the simplest, consistent with reasonable accuracy, for estimating potential productivity of field crops in the local environment.

CHAPTER II

WATER USE OF MAIZE AND BRANS

Site

The experiment was situated on the Muca Irrigation Settlement (latitude 0°58'S, longitude 57°22'B, altitude 1,300 metres), some 110 km from the home base of B.A.A.P.R.O. Although this distance raised many probleme in day-to-day supervision, the site itself was excellent for the purposes of the study.

With some 3,000 hectares under furrow rice irrigation, the site was sufficiently large to give uniform conditions of fetch over a considerable distance.

The position of the experimental plots with respect to the irrigated area is shown in Figures la and 1b. Figure 2 shows a comparison between Perman Bo and evaporation from a raised gridded pan in the meteorological enclosure. Over the period April to November when the rice fields were flooded, Perman Bo was consistently higher than pan

evaporation but when the fields dried out in December,
pan evaporation was more than Penman Bo, indicating
that there might well be an appreciable advected component of evaporation during the period when the rice
fields are not flooded. There may be a residual advection component also when the fields are flooded, but this
is likely to be less than logs.

types both derived from massive volcanic lava flows from Mount Kenya. On well drained sites highly permeable red knolinitie clays have been formed and in areas with impeded drainage, heavy black montmorillonitie clays are found. The experimental site is on an area of level red soil underlain at depths varying from 1.8 to 2.5 metres by tuff rock. Chemical analyses of the experimental soils (Appendix 1) showed no sign of salinity problems. The experimental area had been levelled in the past and it is evident that some 15 cm of soil had been moved from the profile in field 3 on to that in field A. Results of a physical analysis of pore space and water holding capacity are shown in Table 1. The soil is extremely permeatity are shown in Table 1.

eable down to the rock base with a moderately high water holding capacity.

The average rainfall is 800 mm falling in two seasons, April-May and November-December. The annual evaporative demand (Penman Bo) is about 1,800 mm.

Methods and Procedure

hectare plots (Fig. 1a). A hydraulic weighing lysimeter was set in the middle of each plot and a central plot 12 metres square was marked around each lysimeter. Four sub-plots 2 m x 2 m were marked at the corners of each central plot and neutron moisture meter access pipes were installed to a depth of 180 cm in the sub-plots and 170 cm in the lysimeters.

Tensionsters were installed at 170 cm, 150 cm and 70 cm depths in the central plot and at 170 cm and 150 cm in the lysimeters.

Lysimeters

The lysimeters used in this experiment were 2 m x 2 m x 2 m deep, the design, construction and operation

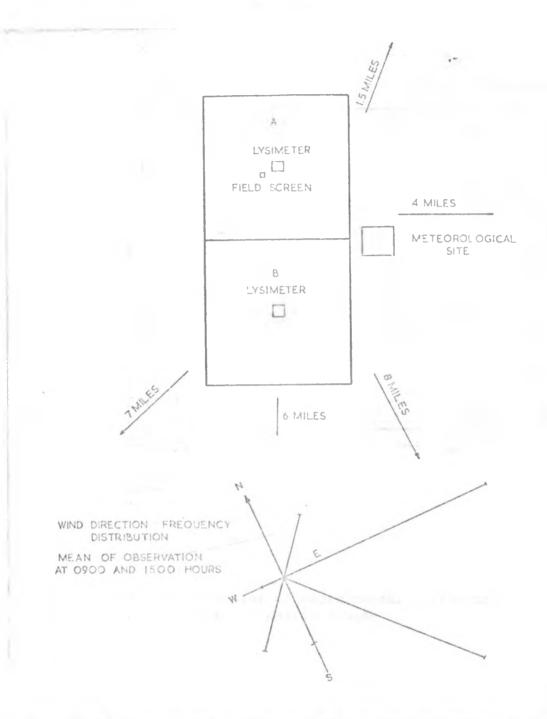


Fig. la: Layout of the experimental field showing direction of prevailing wind. The arrows indicate extent of fetch over flooded paddy fields.

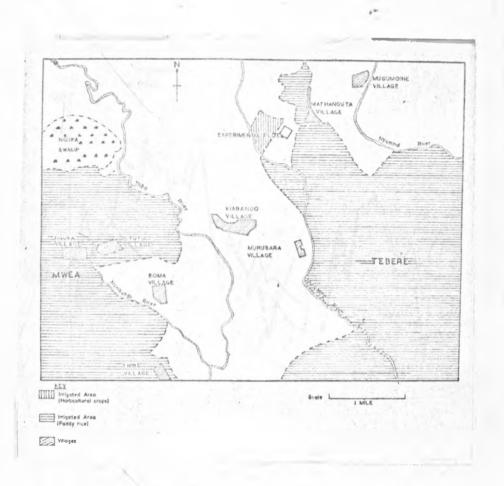


Fig. 1b. Position of experimental plot within the irrigation scheme

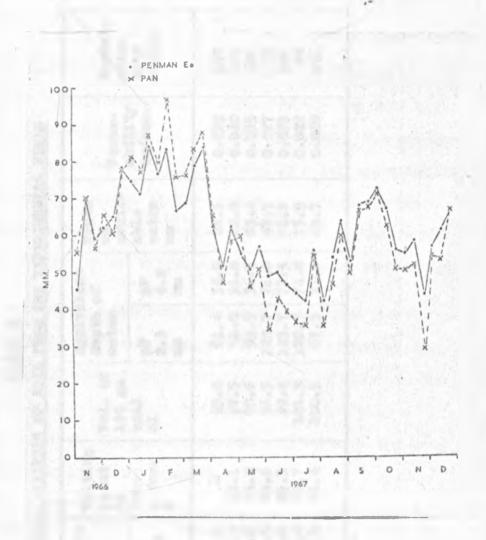


Fig. 2. Comparison between Penman Eo and evaporation from raised gridded pan in the meteorological enclosure

PHYSICAL ANALYSIS OF SOIL FROM THE EXPERIMENTAL FIELD

Depth	Gravimetric moisture content		Wilting point by	Vel. of	Rquivalent depth of water		Depth of avail-	Perme-	Percol- ation
	pF 2.5 \$	pP 4.2 \$	sun- flower	soil	pF 2.5	pF 4.2	able water mm	ability on/sec.	rate cm/hr.
0-15	38.4	25.1	_	93.8	61.4	36.9	24.5	0.039	140
15-30	38.2	23.4	23.1	97.5	58.8	36.0	22.8	0.040	144
30-45	39.4	24.0	24.2	98.6	59.9	36.5	23.4	0.035	126
45-75	39.7	24.4	24.6	96.3	123.7	76.0	47-7	0.027	97
75-105	41.0	25.0	25.5	99.4	123.7	75.5	48.2	0.029	104
105-135	43.7	25.4	25.7	101.9	128.7	74.8	53.9	0.024	86
135-150	45.3	25.3	26.4	100.9	67.5	37.6	29.7	0.028	101

of this type of lysimeter being described by Foregate, Hosegood and McCulloch (1965).

A tank of soil containing the crop planted to match that in the field is supported on flat flexible metal bolsters that are filled with water. The pressure of the tank and soil is supported by a balancing column of water connected to the bolsters. Changes in the water content of the tank are reflected by changes in the height of the balancing column which is read daily.

Several weeks after the lysimeters were first installed, it became apparent that the lysimeters were not performing properly and the levels of water in the balancing columns were rising steadily, implying a reduction in the bearing surface area between the lysimeter tank and the bolshops.

It was suspected that this trouble was caused by stretching the bolsters beyond the elastic limit of the material during installation, and this was partially confirmed by preliminary tests on specimen bolsters. A test rig set up at Muguga was used to subject the bolsters to cycles of pressure up to half an atmosphere to check for leaks and for consistency in performance.

The test rig was unde of two concrete clabs 2 mm x 1.2 m x 15 on cast round flat hoops of steel bar such that one clab was suspended 5.7 cm above the other. Solutors could then be pushed into the slot and considerable pressures imposed without any danger of the bolster swelling excessively. On initial subjection to pressure of half an atmosphere, the bolster took some days to settle down but less time was required on subsequent recycling. All the new bolsters for the experimental lysimeters were tested in this rig before they were used to replace the faulty bolsters. The trouble did not recur and has not been observed in other lysimeters with bolsters treated in the same way.

After a few months of satisfactory performance, the lysimeters showed signs of leakage in the measuring system and were once more excavated. The hard plastic connecting pipes between the bolsters were found to have developed fine cracks at sharp bends. All pipes were replaced with more pliable but durable plastic hosepipe which subsequently proved satisfactory. The lysimaters were back-filled with fresh soil excavated

from pits at the edge of the field. The initial test period under a waterproof tarpaulin was shortened by the necessity to plant the maise crop on time, but this and later tests after harvest indicated a reasonably satisfactory level of performance.

Both lysimeters were functioning properly when
the bean crop was planted at the end of Pebruary 1968,
but after three weeks one of the lysimeters developed
a leak in the weighing system and had to be excavated
and lifted out. The leak was traced to pin holes on
one bolster, apparently caused by electrolytic corrosion
in the metal. This bolster was replaced and the lysimeter tank refilled with soil but although performance
was good at the beginning, sudden drops in column height
continued to occur at random, making estimates of grop
water use from this lysimeter unreliable.

The lysimeter measuring systems were calibrated using field assistants as weights. Figure 5 shows sample calibration lines for the two lysimeters. The calibrations were always linear with 1 mm change in

column corresponding to nearly 1 mm of water on the lysimeter.

A tension drainage system was installed in each lysimeter. This consisted of two sets of 4 ceranic candles, 16 cm long and 4.5 cm in diameter, connected to two plastic tubes dipping in water 100 cm below the level of the bottom of the lysimeter. These candles performed satisfactorily throughout.

Neutron moisture meters

Two different makes of neutron moieture meter, the Electronic Associates Limited (E.A.L.) model and the MIV-I, a Russian make, were tested in this experiment (Figure 4).

Calibration was obtained by digging volumetric soil samples from sets of three profiles 15 cm from neutron probe access pipes placed in dry, moist and wet ground in the same plot. The volumetric moisture contents thus obtained were plotted against neutron count rates in the same soil horison.

the B.A.L. neutron moisture neters made available by
the Institute of Hydrology, Vallingford, U.K., for
eatchment research work in East Africa. Celibration of
this instrument in the soil moisture range 0.25 to 0.38
moisture volume fraction is shown in Figure 5a. The
special batteries required for the MIV-I moisture
meters could only be obtained overseas and this led to
considerable loss of time. Calibration for this
instrument over the same range of soil moisture contents
is shown in Figure 5b.

The range of soil moisture contents used in the above calibrations is comparable with the range of available soil moisture in the top 50 cm (see Table 1). Field capacity figures in Table 1 were obtained by draining core samples under atmosphere tension in the laboratory and although these figures agree with field observations for the top 50 cm, there is no real evidence that the higher values indicated for lower depths were achieved under field conditions.

Tensiometers

Tensionsters were made by sealing a small ceramic tensioneter cup 5.5 cm long, 2 cm diameter, on to one end of a 15 mm diameter hard plantic pipe. A glass reservoir with a narrow side tube was fixed to the other end of the plastic pipe. The side tube was connected to a mercury manometer comprising a glass capillary tubing, 2 mm bore, dipping into a mercury reservoir. This use of narrow glass tubing for the manometer reduced response time and the quantity of exchange water between tensionster cup and the soil. The mercury manometers were mounted inside two wooden shields painted white to reduce heating, and were read only once a day at a time (0900 hours) when day-to-day temperatures were nearly the same. By using freshly boiled water throughout the system. formation of air bubbles was greatly reduced and flushing of the connecting tubos was usually necessary only once every two weeks. All tensiometers were tested in the laboratory to a suction of 47 cm Hg before installation, atmospheric pressure at the site being only 58 on Mercury.

The tensioneter arrangement is shown in Figure 6. The first attempt to use these tensionsters in the field failed due to the unexpected flattening of the strong plastic tubing where it was exposed to the sun, and later the development of leaks at the joints between the glass reservoirs and the plastic pipes. The leaks were subsequently stopped by sliding close fitting rubber sleeves over the joints and applying grease. The flattening of the plastic tubing persisted, though to a lesser extent, after burying the tube in the soil over as much of its length as possible. In 1968, it was possible to replace the flattened connecting tubes with tough translucent P.V.O. tubing. This tubing proved more heat registant but was apparently more attractive to redents which bit holes in the tubing causing frequent losses of data.

when a third treatment was included in the 1968 experiment, additional tensionsters could not be constructed in time and commercial irrowsters with various gauges were used. The depths of measurement were, however, restricted to 120, 100 and 50 cm, these being the lengths of the available irrowsters. These irrowsters had the advantage of registering actual

soil moisture tension without the necessity to correct for a hanging column of water, but the vacuum gauges sould only be read to 1 cm Hg and they deteriorated quickly under continuous exposure in the field; the indicator mechanism became unreliable due to sticking.

Meteorological site

Measurements of maximum, minimum, wet and dry bulb temperatures, hours of sumshine, run of wind, rainfall and radiation (Gum-Bellani radiometer) were made in an enclosure established immediately upwind of the experimental site (Fig. 7).

Open water evaporation was recorded in the enclosure from a raised Kenya type evaporation pan 122 cm in diameter and 45 cm deep covered with chicken wire to keep birds off.

A second meteorological screen was mounted on a strong vertical metal pipe in the field near lysimeter A so that screen height sould be adjusted to 30 cm above the orop. Maximum, minimum, wet and dry bulb temperatures in this screen were recorded at the same time as the meteorological site screen temperatures at the edge of the field.

ords from the two screens (Table 2) over a period of 4 months didnnot reveal any appreciable differences and the difficult operation of the field screen was discontinued.

Data recorded at the meteorological site were used to calculate the Penman estimate of open water evaporation, Bo, using tables published by McCulloch (1965). Monthly means of the above weather parameters are presented in Appendix 2.

Irrigation

Initially water was pumped from an irrigation canal at the edge of the field into an overhead spray system consisting of a single line of seven sprinkler heads covering a strip 10 m x 100 m. The pump was not large and took 1 hour to apply 6 mm on a strip.

Uniformity of application over the central area as assessed by 8 small raingauges was reasonably good, but variations of up to 15% were observed.

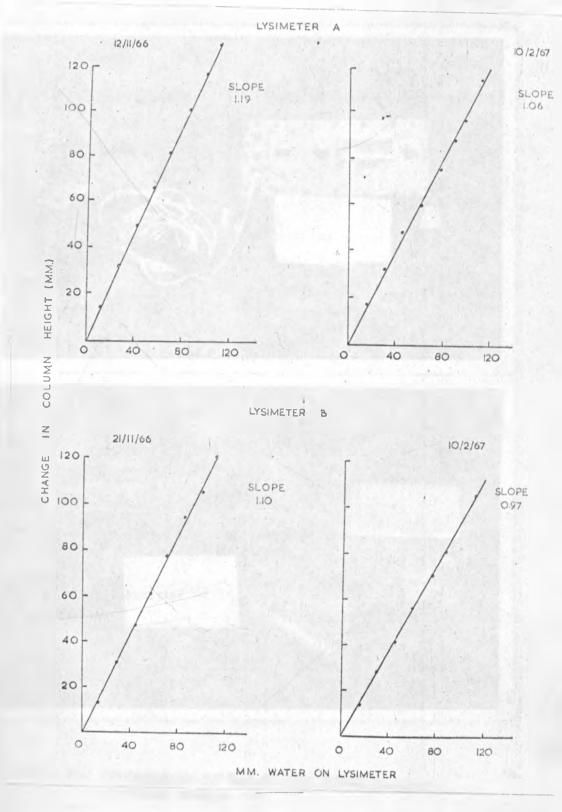


Fig. 3. Sample calibration lines for the two lysimeters

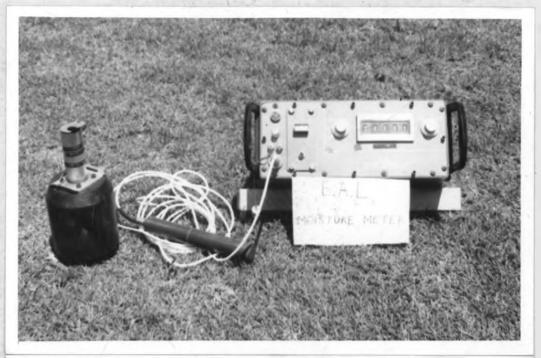




Fig. 4. Photographs of B.A.L. and HIV-I neutron moisture meters

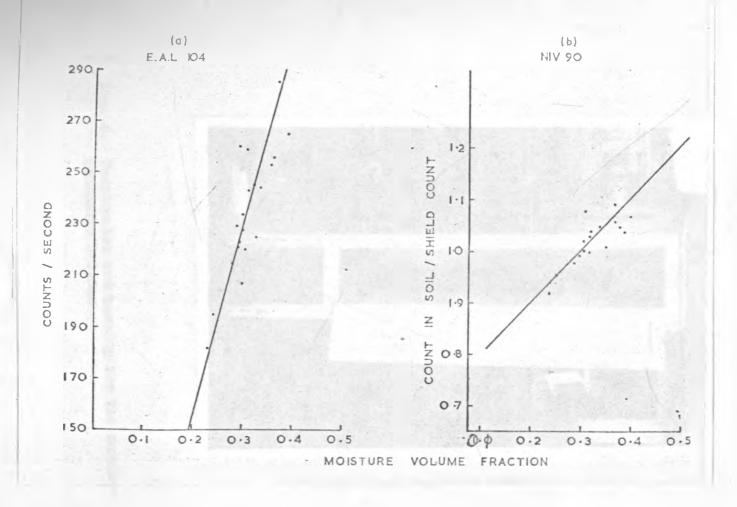
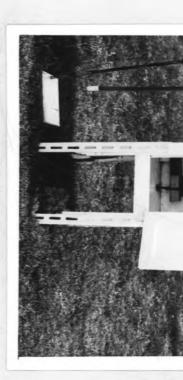


Fig. 5. Calibration of E.A.L. and NIV-I Neutron moisture meter in Mwea soil

Tensioneter and housing for the mercury manometer columns

MAKERERE

THERAPL TO







Pig. 7. Meteorological analogure at the eastern edge of the experimental plot. Paddy fields can be seen in the background.

Comparison between temperature and humidity records from the mutempological site screen (M.S.) and the screen 30 cm above the maise crop (F.S.)

		Ai	r temp	oratur	·e	8.7	Dewpoint mean of 0900 and 1500 hre		mean of delicit		Relative		
Period	Ka	x.	MA	n.	nynn and								and
1967	0	C.	O	c.	O	c.	O	c.	mm.			*	
	M.S.	F.S.	M. S.	P. S.	H.8.	P.S.	M.S. P.S.	M.S.	7.8.	M.S.	J. S.		
Sept.	28.2	31.3	14-3	14-7	21.3	22.2	15.4	14.5	7.47	8.68	40	40	
Oct.	28.0	27.7	15.8	16.5	21.9	22.1	15.6	17.1	6.42	5.32	48	55	
Nov.	26.3	26.6	15.4	16.0	20.9	21.5	17.9	18.3	3.15	3.22	66	66	
Dec.	27.2	27.4	12.5	13.2	19.9	20.3	16.2	16.5	3.61	3.97	56	55	

Although the capacity of this equipment was sufficient to supply routine irrigation requirements. it soon proved completely inadequate to water the 2 hectares of land within the prescribed experimental schedule and was replaced with a new system. A more powerful motor and more appropriate pump, line pipes and sprinklers were installed. In the new system, two lines of 8 sprinklers covered a strip 25 m x 100 m and applied water at the rate of 10 mm per hour in each strip. The uniformity of water application was reasonable (Table 3) but there was a large discrepancy (approximately 25%) between the amount of water collected in raingauges and the estimated irrigation from the metered application. The flow meter had been calibrated by a commercial firm prior to installation but a check carried out in situ by pumping water into a 490-litre tank revealed that the flow meter overestimated the quantity of water flowing through it by approximately 8%. This left 15% to be accounted for by direct evaporation of the fine sprinkler droplets before they reached the ground. In order to minimise this error, most of the irrigation was therefore

PARLE 3

Uniformity of water application

(a) Typical catches (inches) in nine raingauges distributed in an area 20 ft. x 20 ft. receiving water from four sprinklers spaced 40 ft. x 40 ft. and with spray diameter of 80 ft.; bare ground.

Cours We	Irrigation pipe settings									
Gauge No.	1	2	3	4	5	6				
1	3.79	5.75	4.39	5.66	4.38	7.34				
2	4.59	4.38	9.51	4.42	4.57	5.28				
5	3.74	4.39	4.84	3.46	4.14	4.26				
4	3.89	3.78	5.55	5.23	4.21	4.04				
5	3.68	4.27	4.68	4.93	4.94	5.16				
6	3.84	4.32	5.58	4.31	4.88	4.44				
7	3.35	3.71	4.98	5.12	5.10	3.79				
8	3.65	3.87	5.50	4.52	4.94	5.48				
9	4.19	2.81	5.08	4.54	4.85	4.24				
Total	34.72	37.28	50.11	41.99	42.01	43.97				
Mean	3.86	4.14	5.57	4.67	4.67	4.89				
Setimate from flow meter	5.85	5.85	5.85	5.85	5.85	5.85				

(b) Catches (mm) in raingauges placed at the subplots in each central plot. Height of maise approximately 50 cm.

ma air a		lauge in	a subple	*	Central	Retimate		
Pield	1	2	3	4		from flow meter		
A	26.4	27.7	29.5	47.8	31.0	57.6		
3	20.6	19.8	19.8	21.5	25.1	27.5		

carried out at night when evaporation is very low.

Suitable exposure for raingauges proved difficult when
the mains grow higher than 50 cm, the height of the
gauge rim from the ground in normal raingauge exposure.

The first irrigation was applied just before planting and was designed to provide uniform soil moisture
conditions throughout the experimental area. After
sampling both fields to determine the soil moisture
content at all depths up to 180 cm, the profile was
brought up to field capacity by irrigation.

The various treatments were intended to be irrigated on the same day each time, applying 1.1 and 0.7 times as much water as had been evaporated in the wet and dry treatments respectively as recorded by the corresponding lysimeter. The above figures were however changed in 1968 from 1.1 and 0.7 to 1.2 and 0.6 to increase the difference in available water in the two treatments. It was not found feasible to irrigate the two fields on the same day, but by shifting the pipes from one field to the other, it was possible to irrigate at least all of the central strips on the same day. The irrigation interval was set at two weeks but varied

somewhat because of minor breakdowns in equipment. The sprinkler heads had to be raised on extension pipes as the crop height increased.

Because of heavy rain, no irrigation was necessary during the beam crop.

Cropping seasons

The first erop of maise was planted in September 1967. The maise hybrid H 511, maturing in four months and recommended for altitudes around 5,000 feet, was chosen in an effort to keep the growing season within the period when the surrounding rice fields were flooded, and hence to avoid possible advection effects. All measurements proceeded normally but unusually heavy rain in October removed irrigation treatment effects and the two fields became replicates of the wet treatment. In the following long rains season starting in March 1968, the entire field was planted with beans (Phaseolus vulgaria var. Canadian Wonder).

The bean crop was harvested in May and a second maise crop, hybrid H 511, was planted in July 1968.

With the exception of one lysimeter, the equipment did not give trouble and all measurements were made regularly throughout the season. There was very little rain during the major part of the growing season up to maturity and the irrigation treatments were applied throughout.

Land preparation and planting

Before each planting, the land was ploughed and single superphosphate was broadcast at the rate of 382 kg/ha before harrowing. Two mains seeds were planted per hill at 20 cm spacing in rows 1 metre apart. Thinning took place one week after emergence, leaving only one seedling per hill. 560 kg/ha of sulphate of ammonia was applied by hand in two lots - one third as a side dressing on both sides of each maise row at planting and the remaining two thirds along the line halfway between the rows when the maise was 50 cm high. Beans were planted singly at 15 cm spacing in rows 50 cm apart. He fertilizer was applied to the bean crop. The fields were weeded by hand at least twice in the season.

Plant measurements

For each crop of maise, three plants were selected at random in each sub-plot and the lysimeter. Plant heights were measured twice a week and the averages for each plot were calculated.

Crop performence

The 1967 maise crop was adequately supplied with water, germination was 100% and the crop grew uniformly. It was necessary to spray twice with 1% D.D.T. to control stalkborers. The yield from this crop was 5.6 ton/ha. The 1968 maise crop was equally good in germination, uniformity of growth and in yield in the wet and medium treatments. The dry treatment had a good start but growth rate and uniformity of cover were affected after the irrigation treatment was imposed. There was considerable loss of grain due to bird damage and rotting caused by heavy rain at harvest. The bean crop was healthy and grew uniformly. The low yield (1.5 ton/ha) was caused by rotting as a result of heavy rain at harvest time.

Regults

Grop Water use - lysimeters

Results from the 1967 maise erop are shown in Fig. 8 where St/Bo is plotted in running averages of two 5-day periods. Lysimeter A received the higher water treatment at the beginning, but the difference was removed by heavy rain later. There was a high peak St/Bo of up to 1.6 lasting from tasselling to grain formation. A similar presentation (Fig. 9) shows the water use pattern of the bean crop. A broad peak St/Bo of 1.5 was recorded.

shown in Fig. 10. The high peak Bt/Eo observed in the 1967 maise crop did not occur; instead there was a gradual increase, with small peaks corresponding to irrigation. It is interesting that the maximum Et/Eo of 1.2 was obtained late in the season when most of the leaves had dried off and hence there was little transpiration, and that this maximum corresponded with the onset of heavy rain (see discussion on evaporation from "web" and "dry" leaves on page).

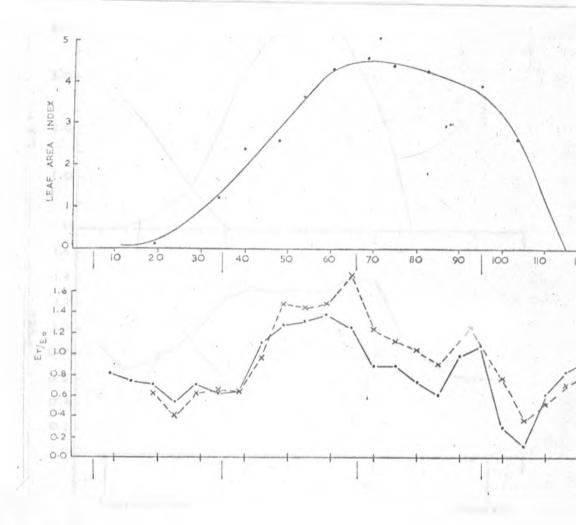
The validity of the above measurements is discussed later. In general, the pattern of water use followed the development of leaf area, but the actual relationship between St/Ro and ground cover was different with each crop. Details of lysimeter data and computation of St/Ro are contained in Appendices 5, 4 and 5. A neutron moisture meter was not available during the 1967 maise crop. The water use data derived from gravimetric soil sampling are shown in Appendix 6.

Cron water was - neutron moisture meters

The neutron moisture meters used in the 1968 beam crop were the B.A.L. 104 and the MIV-I/90.

Details of estimates of moisture content at different depths in each profile are tabulated in Appendices 7a to 7e.

Taking the MIV-I/90 data first and comparing the data in the last four columns of Table 4, there is little agreement between estimates of Et for both fields and the lysimeter either in individual periods or in totals for the period let March to 27th March.



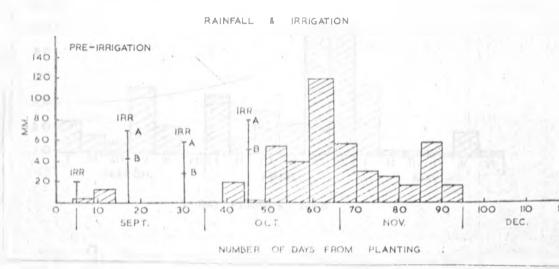


Fig. 8. Water use of maize (1967-68 crop) in Lysimeter A (.-.) and Lysimeter B (x----x)

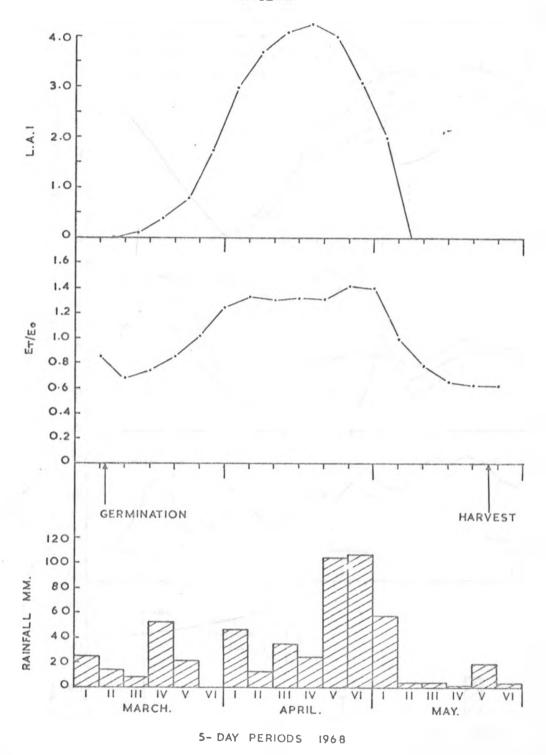


Fig. 9. Water use of beans (1968 crop) Lysimeter A

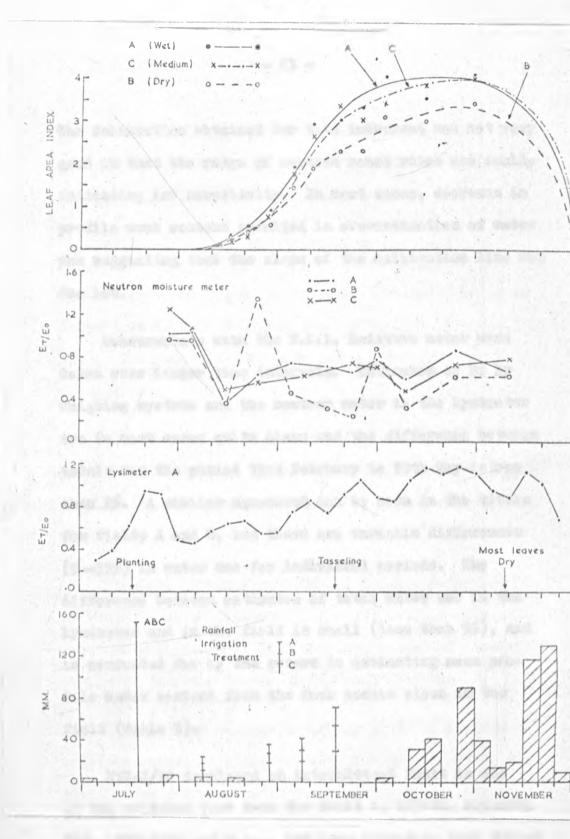


Fig. 10. Water use of maize (1968 crop). Data from Lysimeter A and E.A.L. 104 neutron moisture meter

The calibration obtained for this instrument was not very good in that the range of neutron count rates was small, indicating low sensitivity. In most cases, decrease in profile wear content resulted in overestimation of water use suggesting that the slope of the calibration line was too low.

Measurements with the E.A.L. noisture noter were taken over longer time intervals. Estimates of Et by weighing systems and the neutron meter in the lysimster are in most cases quite close and the difference between totals for the period 25rd February to 29th May is less than 25. A similar agreement can be seen in the totals for fields A and B, but there are variable differences (5-35%) in water use for individual periods. The difference between estimates of total water use in the lysimeter and in the field is small (less than 5%), and is accounted for by the errors in estimating mean profile water content from the four access pipes in the field (Table 5).

MIV-I/90 developed an intermittent fault in one of the switches (see data for field A, 1/3/68, Appendix 7b). MIV-I/96 and B.A.L. 104 were therefore used during

the 1968 maise crop. Detailed data for each profile are shown in Appendix 7e. The equivalent WIV-I/90 readings were obtained from the equation

$$\frac{r}{r}$$
(NIV-I/90) = 1.042 $\frac{r}{r}$ (NIV-I/96) + 0.055

derived from an earlier comparison. — is the ratio of count rate in soil to count rate in shield. The moisture volume fraction was then read off the MIV-I/90 calibration.

proved unsatisfactory and were not comparable with parallel estimates with the E.A.L. moisture meter (Table 6).

The data obtained with E.A.L. 104 are summarised in Table 7, which includes Pennan estimate of open water evaporation, Eo, soil moisture tension and tension gradients near the bottom of the soil profile.

A comparison between estimates of Rt by neutron moisture meter and lysimater weighing system (Table 8) shows large differences in individual periods. The

Comparison between estimates of Et by weighing lysimeter and neutron moisture meters - bean crop 1968

TABLE 4

Date 1968	Mean wate content o profile mm (neutr		Mean w.c. in Lysimeter			se in total content nm. Rainfall Lysimeter drairage mm.			t mm.	Et Lysimeter A		
	Field A	Field B	mm (probe)	A	В	Lys.A			Field A	Field B	Weigning system	Neutron meter
					NIV-9	O NEUTR	ON MOISTURE	METER				
1/3 9/3 13/3 16/3- 20/3 27/3	647 594 614 601 577 566	644 626 550 616 618 633	555* 530 575 558 551 540	+53 -20 +13 +24 +11	+18 +76 -66 -2 -15	+25 -45 +17 +7 +11	39.1 9.1 0.0 52.1 22.9	-	92.1 -10.9 13.0 76.1 33.9	57.1 85.1 -66.0 50.1 7.9	34.6 17.1 10.5 21.2 28.9	64.1 -35.9 17.0 59.1 33.9
Total							123.2		204.2	134.2	112.3	138.2
					EAL 1	04 NEUT	RON MOISTUR	E METER				
23/2 6/3 17/4 2/5 8/5 29/5	518 636 664 778 743 676	566 688 757 843 808 730	50) 639 626 715 691 655	-118 -28 -114 +35 +67	-122 -69 -86 +35 +78	-130 +13 -29 +24 +36	139.5 211.6 198.1 55.9 31.8	19.8 50.8 8.5	21.5 183.6 84.1 90.9 98.8	17.5 142.6 112.1 90.9 109.8	40.8 231.9 106.7 33.4 73.1	9.5 224.6 89.3 29.1 59.3
Total							636.9	79.1	478.9	472.9	485.9	411.8

^{*}estimated from Lysimeter B

TABLE 5

Errors due to spatial variation in the estimates of average Et by the neutron moisture meter

1968 Bean Crop

Date		ure content 4 profiles		ease in l water ent	E	t
1968	1.00.	Standard Error ± mm.	illi.t o	Standard Error	hillig •	Standard Error ± mm.
		II.I.V.	90			
1/3 9/3 13/3 16/3 20/3 27/3	647 594 514 601 577 566	7 7 7 12 11 19	+53 -20 +13 +24 +11	9.9 9.9 13.9 16.3 21.9	92.1 -10.9 13.0 76.1 33.9	9.9 9.9 13.9 16.3 21.9
Total			+81	20.3	204.2	20.3
		E.A.L.	104			
23/2 6/3 17/4 2/5 8/5 29/5	51.8 636 664 778 743 676	9. 12 9 15 15	118 28 114 +35 +67	15.0 15.0 17.5 21.2 20.5	21.5 183.6 84.1 90.9 98.3	15.0 15.0 17.5 21.2 20.5
Total			-158	16.6	478.9	16.6

TABLE 6

Comparison between estimates of Et by the E.A.L. and N.I.V. neutron soisture antage

Rented	Pie	14 A	Fiel	4 3	Pield C		
Period 1968	PAL MM	MIA	PAL		BAL .	MIA	
7/8 - 21/8	18.6	200.8	18.1	-27.8	25.0	•	
21/8 - 28/8	22.6	13.2	47.5	33.3	19.6	56.0	
28/8 - 11/9	61.1	115.5	37.5	190.4	82.8	73.9	
Total	102.5	329.5	103.1	195.9	127.4		

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CROP WATER USE ESTIMATED FROM NEUTRON MOISTURE METER READINGS

(EAL 104. MAIZE CROP)

	Et (nn)			Penman		· Et/Eo			IEAN soil moisture tension T and Tension gradient dT at 160 cm. depth*				
Period 1968	Field A	Field B	Field C	Eo (mm)	Field A	Field B	Field 0	Fie	ďΣ		eld B		ld C dT dz
23/7 - 31/7	31.0	29.1	37.8	30.3	1.02	0.93	1.25		-	-	_	242	5.0
31/7 - 7/8	28.0	25.9	29.1	27.2	1.03	0.95	1.07	146	0.4	151	-0.6	į.	4-4
7/8 - 21/8	18.6	18.1	25.0	49.7	0.57	0.36		155	0.5	140	-0.1		4-4
21/8 - 28/8	22.6	47.5	19.6	35.2	0.64	1.35	0.56	156	0.9	153	-1.5	312	3.9
28/3 - 11/9 11/9 - 19/9	31.1 38.5	37.5 17.1	}84.5	81.8 53.4	0.75	0.45	0.63	160	0.7	145	-6.5 -7.7	315	4.1
19/9 - 26/9	55.3	12.7	37.5	50.1	0.56	0.25	0.75	161	1.0	115	-7.5	420	-3.1
26/9 - 3/10	43.0	47.8	38.4	54-1	0.79	0.83	0.71	165	1.3	135	-5.1	416	-0.1
3/10 - 15/10	51.6	27.3	42.8	86.6	0.60	0.52	0.49	170	2.0	143	-5.9	510	-0.7
15/10 - 31/10	34.1	59.5	70.7	96.9	0.87	0.61	0.73	179	3.5	196	2.5	563	-1.2
31/10 - 19/11	71.6	65.5	82.7	106.3	0.67	0.62	0.78	-	-	163	4.7	4201	12.8

*110 cm. in Field C

belong the the sensor stated by hitty, an ele-

Comparison between estimates of Et by neutron muthod

(B.A.L.104) and the weighing system in the lyminuter

medicine, arthropius his medicaliting the he was about he have

semination of mail melidose profiles (nor Aspealin 7).

MARKET BY MARK	Et (Lysis	meter A)	Bt (Field A)
Period 1968	weighing aystem	neutron meter	neutron meter
23/7 - 31/7	24.9	1,8	31.0
31/7 - 7/8	9.4	7.1	28.0
7/8 - 21/8	25.8	26.5	18.6
21/8 - 28/8	19.0	33.7	22.6
28/8 - 11/9	61.5	55.7	61.1
11/9 - 19/9	40.1	26.1	38.5
19/9 - 26/9	50.2	62.7	35.5
26/9 - 3/10	52.4	27.8	45.0
5/10 - 15/10	81.3	73.7	51.6
15/10 - 51/10	118.6	80.4	84.1
31/10 - 19/11	101.9	79.4	71.6
Total	585.1	474.9	485.4

14 to be named below 16.

earlier, precision in estimating Et in the field by the neutron moisture meter is reduced considerably by spatial variability of soil moisture profiles (see Appendix 7). This variability was increased by the uneven distribution of irrigation water.

Another source of error is the gravimetric determinstion of scil moisture content in the top 20 cm, where the loss of neutrons imposes a limitation to the use of the neutron moisture meter. In the early stages of growth and whenever the soil surface is wet, evaporation from this layer of soil may constitute as much as 90% of total Et. An experiment was carried out to determine the error introduced in this way. Ten soil samples at 5 cm and ten samples at 15 cm were taken at the same time within an area of 100 on radius centred on one of the access pipes. The standard error of the mean moisture volume fraction was 0.01, i.e. 2 mm in 20 cm of soil and the standard error of difference between two such estimates would be 0.014, 1.e. 2.8 mm in 20 cm of soil. A time interval of at least 10 days at 6 mm/day evaporation is therefore required if error due to this source alone is to be reduced below 5%.

A check was also made on the precision with which moisture content could be determined at a particular level in the soil. Yen readings were taken in succession with MIV-I and E.A.L. moisture maters both at 30 cm depth in soil and in their respective shields. From these readings, means of two successive readings were salculated, this being the normal procedure in the field.

The results summarised in Table 9 suggest that
greater precision was achieved with N.A.L. than with MIV-I
moisture meters. Since there was no systematic change in
shield count rate, calibration of N.A.L. moisture meter
was based on count rate in soil, and moisture contents in
the field were obtained from this calibration without
reference to shield count. The shield count was nevertheless recorded twice on each day of field measurements
as a check on the performance of the equipment.

Boil moisture tension

The data obtained with tensionsters are contained in Appendix 8. Average gradients of soil moisture tension have been calculated but, for reasons discussed

Precision in estimates of profile water content

by the neutron moisture meter

Description of error	MIN 36	BAL 104
No. of readings taken	10	10
No. of means of 2 successive readings	9	9
Coefficients of variation		
(a) counts in soil (B) (b) counts in shield (B _B) (c) B R ₈	0.84% 1.85% 2.05%	0.65% 0.67% 0.935%
Mean R	0.978	0.466
Standard errer (S.E.):	al major	MARKET THE
Ra	0.0066	0.0044
Moisture volume fraction	0.008	0.003
mm/20 cm soil mm/soil profile mm/period, i.e. difference between	1.6	0.6
2 profiles	18.1	6.8

in the previous section, no attempt has been made to calculate drainage from these data.

Yield

Grain yields obtained in the sub plots and the lysineters are shown in Tables 10 and 11.

Discussion of Results

Reliability of lyaimeters

The validity of the crop water use data and the assessment of the performance of the neutron moisture meters depend on the reliability of the operation of the lysimeter system which is therefore discussed first.

There are three main factors involved:

- 1. The accuracy and sensitivity of the measuring system;
- 2. The stability of the balancing column over long periods;

TABLE 10
Yield of experimental plots - 1967 maige eren

Treatments and plots	Weight of dry grain kg	Yield ton/ha	Weight of dry stalk kg		
Lysineter A	1.56	5.9	3.4		
A-subplote:					
A	1.42		2.4		
В	2.19		2.8		
C	1.57		4.3 3.4		
	2.71				
Average	1.64 (1.45)*	4.1 (3.7)*	3.2		
Lycimeter B	1.58	3.9	5.5		
B-subplote:			1		
B	1.27		3.7		
2	1.42		3.7 3.4		
G	1.55		2.4		
H	1.62		3.6		
Average	1.46	3.7	3.3		

Total grain yield from both fields = 10.8 metric ton Weight of 1,000 grain at 14% mc. = 3.27 kg.

^{*}Excluding the high value of 2.19

IN STREET

facatrogas 80gf out at ablock salad

No. of cobs	aniara 000 ti 20 fablow	sa/sa	ex 70 plants	Plot	Pield	tanatanati
2°1 2°1 2°1 1°3	475.1 8.64 6.65.6 6.75.7 7.53.4	7,860	5,018 5,175 2,814 5,572 5,144	AA AB AG AB Tield ITOR Town	V	MA
5°5 5°2 5°7 5°7 5°0	6.714 4.854 4.83.4 8.724 2.744	\$ \$ \$ \$ \$ \$ \$ \$ \$ \$ \$ \$ \$ \$ \$ \$ \$ \$ \$	2,945 1,911 5,005 2,867	CI CK CK CL Meen Tield from S rows	э	Margen
5°2 5°2 5°1 5°1	\$21°\$ \$22°\$ \$22°\$ \$24°8 \$72°8	8°218	2,779 2,704 2,995 2,995 2,619	DE BE	æ	IHT

3. The degree to which the crop in the lysimeter represents the field crop in ontogeny, behaviour and constraints.

and McCulloch (1965) for the lysimeters used in this experiment. The balancing column could usually be read easily to 0.5 mm, but fluctuations due to wind sometimes increased this to 1 mm. This can be serious in daily readings but less so in 5-day or 10-day totals. There was little change in sensitivity and error from this source was reduced by using monthly calibrations in the conversion of column readings to water use.

A much more serious source of error was the sero drift, i.e. systematic changes in column height unrelated to changes in the weight of the tank. This could be checked easily before planting by covering the lysimeter with tarpealin and observing any movement in the balancing water column. This check was carried out in 1967 but the period was cut short by planting date for the maise crop, and no conclusion can be drawn from data obtained. The only check that could be applied with a crop on the lys-

imeter was by assuming that the tank and contents would return to the same weight each time the soil was wetted to field capacity. The main difficulty was to decide when this condition was achieved and the start and end of drainage were chosen as the most reliable criteria in this experiment. These cheeks facilitated critical evaluation of lysimeter data as shown in Tuble 12 belows

Column height (cm) corresponding to approximate

field connectty conditions in the lysinsters

Date	Lysimeter A	Date	Lysimeter B
5.9.67	33.2	14.9.67	79.1
1.10.67	24.3*	22.9.67	75.0
9.11.67	26.6	28.10.67	75.2
1.12.67	26.5	10.11.67	71.2
		30.11.67	52.5

^{*}possibly underestimated

The data in Table 12 show a fall in column in lysimeter A in September, but the lysimeter was stable from less satisfactory, starting off with a rapid drop in column. The lysimeter data are reasonable between 22nd September and 25th October, doubtful between 28th October and 24th November, and reliable from November 26th to the end of the season. A sudden column drop of 21.6 on ocsurred on 25th November in spite of 59.4 mm of rain on this day. The five-day period following this day has been removed from St/Bo calculations. The calibration factor changed from 0.88 to 0.98 mm on lysimeter/mm change in column. This reduction in consitivity implies an increase in the bearing area on the bolsters possibly caused by a slight shift of the lysimeter.

harvesting the crop in January 1,68. The data obtained over a period of 38 days are plotted in Fig. 11 together with mean and maximum air temperatures. These data show fall in column equivalent to about 0.5 mm/day in lysimeter A. It is possible that this much evaporation managed to escape through imperfections in the tarpaulin cover. Data from lysimeter 3 show large variations apparently correlated with temperature. Although an attempt

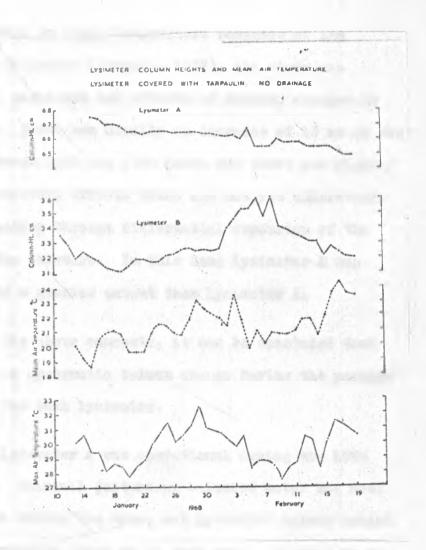


Fig. 11. Lysimeter column height and mean air temperature. Lysimeter covered with tarpaulin and no drainage recorded.

has been made to study temperature response of the hydraulic lysimeter (Wangati, 1965), the study was limited to magnitude and effects of density changes in the water. There was usually an increase of 10 mm on the column between 0900 and 1500 hours but there are clearly other temperature effects which are not yet understood, possibly acting through differential expansion of the tank and the bolsters. In this case lysimeter B was affected to a greater extent than lysimeter A.

From the above commente, it can be concluded that there was no systematic column change during the periods specified for each lysimeter.

Only lysimeter A was operational during the 1968
been erop. Rainfall just about balanced water use most
of the time during the erop, and lysimeter column height
at field capacity (82.0 cm on 28th April and 80.0 cm on
14th May) was sufficiently consistent. Column height
with the lysimeter under terpaulin cover at the end of
the season showed satisfactory stability (72.8 cm on
11th June, 73.3 cm on 25th June).

Although the check with tarpaulin cover after the bean crop had been harvested did not show any drift in the lysimeter column, column height at estimated field capacity conditions during the 1968 maise crop showed a decrease from 76.0 cm on 26th July to 70.0 cm on 21st

November. This means a possible error of 0.5 mm/day in the lysimeter estimate of crop water use.

Another source of error occurred after heavy irrigation. The column took a long and variable time to reach maximum height and then fell suddenly. The data recorded on days of irrigation have therefore had to be discarded and the average for the five-day period used.

The degree to which the lysimeter erop is represented raises problems of a more fundamental nature. When functioning properly, the lysimeter gives an accurate pattern of water use for the crop in the lysimeter. The extent to which this pattern represents the behaviour of a real field crop depends entirely on the differences in soil profile and energy balance between the field and the lysimeter. The results from growth measurements in this experiment (Figs. 12 and 13) are a good demonstration of

higher. This might have encouraged faster growth in the the difficulty in establishing a representative crop in the lysimeter. It was impossible to pack the soil in the lysimeter to the same density as in the field and the water holding capacity of the lysimeter soil was noticeable in the short bean drop, large differences lysimeter crop and although the difference was not were recorded with the maine crop. The difference in height between lysimeter and field crops can affect estimates of crop water use in the lysimeter in two ways:

conditions at the experimental site. ditions. This effect is considered unimportant because of the low wind speed and on-advective ance to evaporation, particularly at high wind would tend to decrease boundary layer resistthe lysimeter considerably in advective con-(a) Increased turbulence is likely to occur within the projecting lysimeter crop. This speeds, and could increme evaporation from

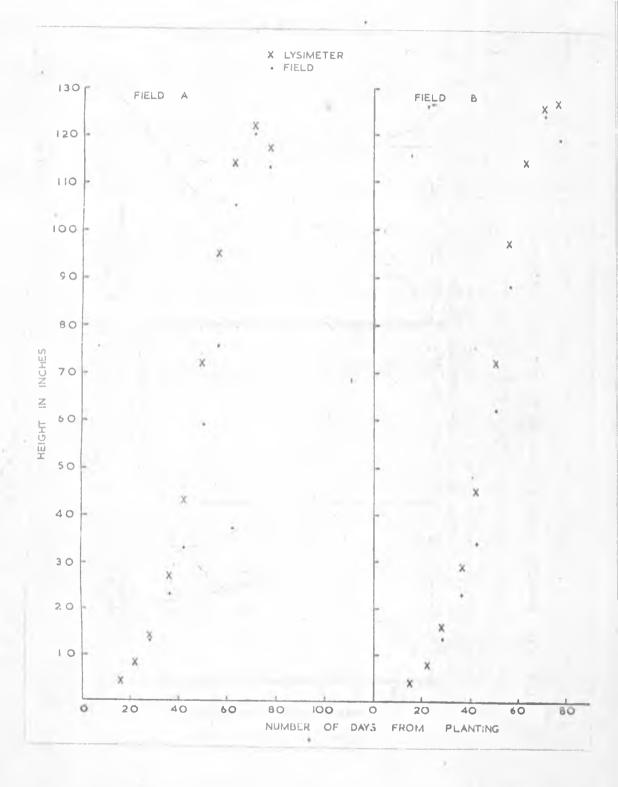


Fig. 12. Average height of maize in the 1967 experiment

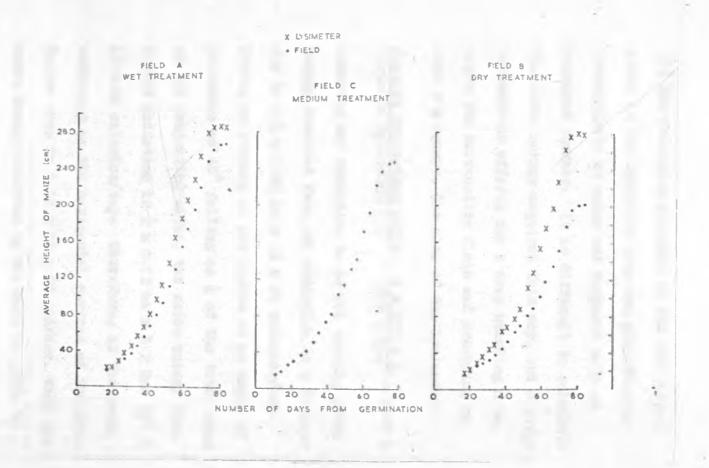


Fig. 13. Average height of maize in the 1968 experiment

(b) The radiation absorbed by the projecting sides of the lysimeter crop can give rise to overestimate of water use compared with an extended surface. It is difficult to calculate the extra energy acquired this way, but a rough estimate is offered for a crop standing h em above the surrounding field and occupying en area 2 m square, the size of the lysimeter.

Area of projecting sides = $\frac{4 \times 200 \times h}{200 \times 200} = 0.02 h$

energy absorbed from sky radiation in a 12 hour day is 0.1 x 0.02 hA x 12 x 60 calories/day.

Taking an average of 1.2 Ly/min at an angle of incidence of 45° falling on 1 of the total area of the projecting sides, the extra energy from direct radiation is 1 x 0.02 hA x 1.2 Gos 45° x 12 x 60 calories/day. Therefore, if the total radiation on the horizontal surface of the lysimeter crop is 1.1A x 12 x 60 cal/day, then the extra energy absorbed by the crop is given by

 $[\]frac{0.0042 \text{ bA} \times 12 \times 60}{1.14 \times 12 \times 60} = 0.386\%.$

Crops	1967 maise crop						1968 maise crop				
Lysimeters	A Oct.	B Oct.	A Oct.	B Oct.	A Oct.	B Oct.	A				
Month:							October		Hovember		r
10-day period:	I	I	II	II	111	III	11	III	I	II	III
Serror in Et due to: "a. sero drift "b. project- ing lysimeter	0	0	0	0	0	o	7	6	10	6	22
crop	4	7	13	10	19	7	8	8	8	8	8
Total	4	7	13	10	19	7	15	14	18	16	30
Average Et/Eo Corrected Et/Eo	0.80	0.80	1.30	1.40	1.40	1.60	1.20	1.20	1.10	0.92	0.80

Assuming that sero drift was constant with time

^{**}Assuming that all the extra energy was converted to latent heat

Prom the above calculation, the extra energy absorbed by the 1967 maise crop (h = 35 cm) was 15%, and for the 1968 maise crop (h = 30 cm), 8% in the wet treatment and 50% in the dry treatment (h = 80 cm).

The above analysis shows that the differences between the height of lysimeter and field crop could have resulted in considerable extra energy falling on the lysimeter crop. The partition of this extra energy into sensible and latent heat is however uncertain, and the actual increase in evaporation is likely to be less than the increases in energy shown above. The effects of these sources of error on Et/Eo are summarised in Table 15.

Evaporation from wet and dry leaves

the high Et/Ro values obtained in these experiments cannot therefore be explained critically in terms of lysimeter performance, advection and differences in the height of the lysimeter and field crops. One factor which has not been considered so far is the evaporation of intercepted water. Peak Et/Ep occurred in periods of heavy rain which, for the 1967 maise crop, coincided

with tasselling, silking and maximum ground cover. For the 1968 maise crop, this peak was absent at the tasselling-silking stage of development, but occurred with the onset of heavy rain when the leaves were drying off, i.e. when the expected transpiration rate would be very low and decreasing. This observation calls for a discussion of the role of aerodynamic and internal resistances in grop water use.

The process of evaporation is the same whether this takes place at an open water surface on the surface of a leaf of a transpiring plant. The rate of evaporation is in all cases governed by the rate of energy input and the resistance to flow of water vapour from the evaporating surface to the surrounding air.

For a surface where the vapour pressure is equal to the saturation vapour pressure at surface temperature, it can be shown (Monteith, 1965b) that

$$\lambda B_{W} = \frac{\Delta H_{W} + \cos(\cos(2) - o)/r_{o}}{\Delta + o} \qquad (211)$$

where E = evaporation rate (gm.cm sec 1)

- λ = latent heat of evaporation (cal.gm⁻¹)
- H = net radiation (cal.ca⁻²sec⁻¹)
- e = density of air (gm.cm -5)
- c = specific heat of air at constant pressure

 (cal.gn -1 oc-1)
- es(T) = saturation vapour pressure (mb) at temp. T
 - e = vapour pressure of the atmosphere (mb)
 - y = paychrometric constant (mb. °C⁻¹) given by
 \[
 \frac{\text{OnP}}{\epsilon\lambda}\] where Cp = specific heat of dry air (per unit volume) at constant pressure, P = atmospheric pressure, and \(\epsilon\) is ratio of density of water vapour to that of dry air at the same temperature and pressure (0.662).
 - r = aerodynamic resistance (sec.cm⁻¹). is

 defined as the rate in which 1 cm³ of air

 exchanges heat, vapour or momentum with 1 cm²

 of the evaporating surface.

The subscript w indicates free water on the evaporating surface.

For a dry leaf surface, Monteith (1965b) also showed that

$$\lambda E_{a} = \frac{\Delta H_{a} + ec\{es(2) - e\}/r_{a}}{\Delta + v(1 + \frac{r_{a}}{r_{a}})} \qquad \dots (212)$$

where the subscript d indicates a surface where vapour pressure is less than saturation vapour pressure at surface temperature and r is resistance to diffusion of water vapour through the stomata, to be known simply as internal resistance.

are the same, the only significant difference between

H_d and H_w would be due to differences in surface radiative
temperature, the wet surface being cooler than the dry.

However, as long as there is adequate soil moisture, the
"dry" leaf surfaces are unlikely to attain a temperature
sufficiently higher than the "wet" leaf surfaces for this
factor to be significant.

Monteith (1965b) argued that since the degree of wetness on the leaf during most of the day had very little effect on their reflection coefficient for solar radiation, H could be equated with H and hence

$$\frac{B_{\text{W}}}{B_{\text{d}}} = \frac{\Delta + \sqrt{1 + \frac{P_{\text{S}}}{P_{\text{B}}}}}{\Delta + \sqrt{2}}$$

$$(215a)$$

The values of $\frac{R}{R}$ would therefore depend on mean air temperature (for \triangle), mean air temperature and altitude (for \vee), and $\frac{R}{R}$.

Monteith's method can be extended to give the where to is the open water evaporation, provided aerodynamic resistance for open water surface (r) is known. It is, however, now necessary to maintain the identity of I and I, since the albedo for the two surfaces may be quite different.

Following from equations (2:1) and (2:2) above.

$$\lambda H_0 = \frac{\Delta H_0 + \cos(es(T) - e)/r_{eq}}{\Delta + \delta} \qquad \dots (215b)$$

and esabining (2:2) and (2:5b), we have

$$\frac{\mathbf{E}_{\mathbf{d}}}{\mathbf{E}_{\mathbf{0}}} = \left\{ \frac{\Delta \mathbf{E}_{\mathbf{d}} + \{\mathbf{e}_{\mathbf{S}}(\mathbf{T}) - \mathbf{e}\} \mathbf{e}_{\mathbf{0}} / \mathbf{r}_{\mathbf{n}}}{\Delta \mathbf{H}_{\mathbf{0}} + \{\mathbf{e}_{\mathbf{S}}(\mathbf{T}) - \mathbf{e}\} \mathbf{e}_{\mathbf{0}} / \mathbf{r}_{\mathbf{n}}} \right\} \left\{ \frac{\Delta + \delta}{\Delta + \delta} \right\}$$

....(214)

The components of total net radiation H are given by

$$H = (1 - ol)R + L_D - L_S$$
(215)

where D and U represent downward and upward fluxes of long wave radiation.

Monteith and Secios (1961) obtained the equation

$$L_{D} - L_{U} = \frac{b}{a} - E(\frac{1-a}{a})$$
(2:6)

or
$$L_p - L_q = L = \frac{b}{a} + \beta H$$
(217)

where the constant β could be regarded as a heating coefficient, and if L=L when R=0, then

$$\mathbf{H}_{d} = (\frac{1-\alpha}{1+\beta})\mathbf{R} + \mathbf{L}_{0} \qquad \dots (2:8)$$

The factor $L_0 = L_0 - L_0$ at mean air temperature T_0 may be calculated from

$$L_0 = (1 - C)L_0 + C\delta(T_0 - 2)^4 - \delta T_0^4 \qquad(209)$$

On cloudy days (0 >1), $\frac{1}{0}$ becomes very much smaller than $\frac{1-\alpha}{1+\beta}$ and can be neglected.

A comparison of published values of β by Ideo, Baker and Blad (1969) suggests an average value of β close to

0.096 for several crops. Therefore, for $\alpha=0.20$, the coefficient of R in equation (2:8) becomes 0.75.

In the case of an open water surface, the ratio of daily back radiation component to the net short wave radiation in the Penman formula was approximately 0.12 at lives; hence

$$H_0 = (1 - <_0)R = 0.12R$$
(2:10)
and for $<_0 = 0.05$, $H_0 = 0.63R$.

If, however, the assumption is made that the main difference between H_d and H_o is in the albedo ($\propto 0.20$, $\propto 0.05$), equation (2:4) becomes:

$$\frac{E_{d}}{E_{0}} = \left\{ \frac{0.80 \Delta E + (0\{e_{0}(T) - e\}/T)}{0.95 \Delta E + (0\{e_{0}(T) - e\}/T)} \right\} \left\{ \frac{\Delta + \pi}{\Delta + \pi'(1 + \frac{\pi}{2})} \right\}$$

where R is now the incoming radiation (cal.om sec-1), and

$$\frac{E}{E_0} = \frac{E}{R_0} = \frac{0.80 \, \Delta R + \frac{1}{10} \cdot \frac{1}{10}}{0.95 \, \Delta R + \frac{1}{10} \cdot \frac{1}{10}} \cdot \dots \cdot (2:12)$$

The corrections contained in (2:8) and (2:10), if applied to (2:11) and (2:12) would result in 10-15%

increase in $\frac{R_{o}}{R_{o}}$ and $\frac{R_{o}}{R_{o}}$, but because of the uncertainties in almost all the long wave radiation factors, these corrections are omitted in the following treatment.

Values of P. P. and P.

Representative values of r, r, and r are required before proceeding with (2:11) and (2:12). These measurements are rare and an attempt was made in this experiment to derive r, from wind profiles, assuming neutral stability, above a nature maise crop. The details of these measurements can be found in Appendix 9. The result was that during the four days of investigation, the value of r, changed with diurnal variation in windspeed. r, was highest (0.12 - 0.15) in the morning before 1100 hrs. and decreased to 0.05 in the afterneon. These values can only be approximate and an average of 0.10 would presumably be valid for a large part of the day. No measurements are available for the bean crop (50 cm high) but a value near 0.15 may be representative.

Similar windspeed profile measurements over open water were not possible and it is necessary to estimate value of r from the Penman equation.

The expression used by McCulloch (1965) for the computation of Penman Bo (open water evaporation) is

$$B_0 = \frac{\Delta}{\Delta + \sigma} \left\{ B_{\underline{u}} (1 - \underline{v})(0.29 \cos \phi + 0.52 \frac{\underline{u}}{\underline{u}} \right\}$$

$$-\frac{\Delta}{\Delta + \sigma} \left\{ \delta \underline{u}_{\underline{u}}^{4} (0.10 - 0.90 \frac{\underline{u}}{\underline{u}})(0.56 - 0.06/6) \right\}$$

$$+\frac{\underline{v}}{\Delta + \sigma} \left\{ 0.26 (1 + \frac{\underline{u}}{20,000})(1 + \frac{\underline{u}}{100})(\cos(\underline{x}) - \underline{o}) \right\}$$

$$\dots (2.15)$$

h = altitude (metres)

u = windspeed (miles per day).

Comparing (2:13) with (2:3b), and recognizing that the term $\frac{e_0}{\Delta + \sigma}$ has to be modified to take assount of altitude,

$$(1 + \frac{h}{20,000}) \frac{8}{\Delta + 8} \left\{ 0.26(1 + \frac{u}{100})(es(T) - e) \right\} = \frac{e}{\Delta + 8} (1 + \frac{h}{20,000}) \frac{(es(T) - e)}{\lambda r_{a_1}} cm/sec. \qquad (2014)$$

But
$$\sigma = \frac{G_{0}}{\epsilon \lambda}$$
;
the term $\frac{e_{0}}{\lambda(\Delta + \sigma)}$ can therefore be replaced by $\frac{e_{0}\epsilon}{(\Delta + \sigma)P}$

giving 0.26(1 +
$$\frac{u}{100}$$
)(es(T) - e) = /6ay
= $\frac{e \in (es(T) - e)}{e \in (es(T) - e)}$ cm/sec.(2:15)

and
$$r = \frac{e \in x \cdot 3.600 \times 24 \times 10}{P \times 0.26(1 + \frac{W}{100})}$$
 (see.em⁻¹)(2:16)

Equation (2:16) can now be used to determine values of r in terms of windspeed for a given location, the only other variables being air density e and atmospheric pressure ?.

Values of r for various windspeeds have been salculated for a few stations (see Appendix 10). The sensitivity of r to the "Balton" factor (1 + 100) is highest
(>2.0) at windspeeds below 60 miles per day, and decreases
rapidly to less than 0.5 at windspeeds above 160 m.p.4.

At Mwea, the average windspeed during the 1967 mains erop and 1968 bean crop was 68 miles per day; r was therefore 1.45. Average windspeed during the 1968 maiss crop was 64 m.p.d., and r had the value 1.48.

There are relatively few measurements of r reported in literature and the data compiled by Monteith (1965b) show wide variations in r. Monteith showed, however, that the empirical relationship

$$\frac{\lambda_{\rm B}}{\rm N-S} = \log(\frac{25}{r_{\rm g}})^{\frac{1}{\rm B}} \qquad \dots (2:17)$$

gave very good fit for the data for $0.25 \langle r_g \langle 10 \rangle$, and suggested that although the constants in the equation may show some variation for different crops, the equation could be used to estimate internal resistance of closed crop canopies.

By rearranging (2:17), Sseics and Endrich (1969) used the equation

$$\log_{10} r_s = 1.40 - \frac{22\lambda E}{N-S}$$
(2:18)

to calculate values of r_s for different vegetation. They found a marked seasonal variation in r_s , but values of the order of 0.9 - 1.5 for pine forest, 0.5 - 1.3 for potatoes and 0.5 - 0.7 for lucerne were similar for south-east England and Califernia.

Values of internal resistance r calculated from hourly energy balance
(Bewen's Ratio) measurements over maise and been crops

Local time	Values of rs (sec.cm ⁻¹)									
	1967 maise crop		19	68 maise c	1968 bean crop					
	20.12.67	21.12.67	12.10.68	13.10.68	14.10.68	6.5.68	7.5.66			
0700-0800	5.5			0.35	1.38	-				
0800-0900	-	0.25	0.07	2.40	0.44	-	-			
0900-1000	0.35	0.22	1.10	1.10	2.51	0.32	0.38			
1000-1100	0.21	0.19	0.01	0.72	0.18	0.24	0.26			
1100-1200	0.19	0.19	1.10	0.35	0.21	0.25	0.23			
1200-1300	0.17	0.21	0.74	0.01	0.22	0.35	0.23			
1300-1400	0.17	0.29	0.10	0.01	0.36	0.36	0.17			
1400-1500	0.24	0.23	1.00	0.42	0.32	0.48	0.36			
1500-1600	0.24	0.29	1.32	0.83	0.35	0.22	0.35			
1600-1700	0.23	0.30	0.42	0.17	0.42	0.29	0.52			
1700-1800	0.32	0.38	0.33	0.24	0.60	0.33	0.38			
1800-1900	-	1- 1	•	-	-	-	0.35			
Mean days from germin- ation	0.24	0.26	0.62	0.60	0.64	0.32	0.32			

Values of r in Table 14 have been calculated from energy balance measurements (see Chapter IV) over mature crops of maise and beans at Mrea. Ground cover was complete in the bean crop but was at most 80% in the maise crops.

Average values of r for the 1968 maise erop (0.6) are higher than the values for the 1967 maise erop (0.25). The average value of r for the bean erop was 0.52.

Calculation of $\frac{R_w}{B_d}$ and $\frac{R_w}{B_0}$

Values of r_a , r_{a_1} and r_{a_2} derived above can now be used in equations (2:5), (2:11) and (2:12), to obtain theoretical or the expected values of the above ratios.

Fig. 12 shows values of $\frac{E}{d}$ for different values of $\frac{r}{d}$ and for different mean air temperatures (altitude 1,500 metres). The altitude dependence of these values is higher for the lower temperatures and higher values of $\frac{r}{d}$. The correction for altitude is, however, at most $\frac{a}{d} = \frac{r}{r} = 6.0$, $\frac{r}{d} = 10^{\circ}$ C, for the altitude range 0-2,000 metres.

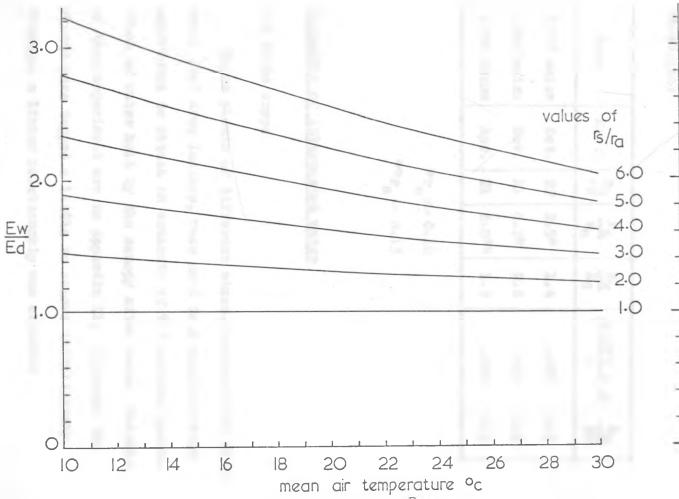


Fig 14: Changes in computed values of $\frac{E_W}{E_d}$ with mean air temperature for different values of r_s/r_a

The expected values of $\frac{B}{B_0}$ and $\frac{B}{B_0}$ (Figs. 14 and 15) for the three crops at lines during peak Bt/Eo are therefore:

Orop	Month	Ta OC	$\frac{r_a}{r_a}$	E _w	(**(T) - *)	E _w
1967 maise	Oct.	22	2.5*	1.4	.015	1.7
1968 maime	Nov.	21	6.2*	2.6	.012	1.5
1968 beans	Apr.	21	2.0**	1.3	.008	1.5

*r = 0.10

Butimates of Intercepted water

(a) Maise crops

Maine plants of different sizes, representing different Loaf Area Indices, were used in a supplementary experiment to obtain an estimate of the maximum quantities of water held by the canopy after rain. Details of this experiment are in Appendix 11. Although the plants were tapped a little to remove transitionary storage, a linear relationship was obtained

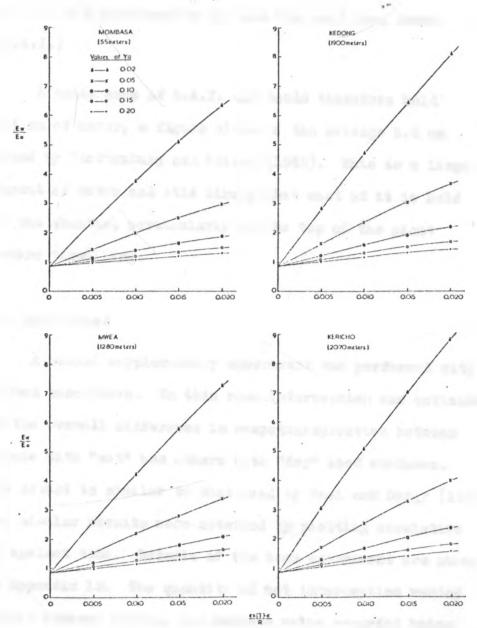


Fig. 15: The response of calculated values of $E_{\it H}$ to aerodynamic resistance $r_{\it a}$ at different altitudes

I = 0.12 L.A.I. + 0.07, r² = 0.89
between met interception (I) and the Leaf Area Index
(L.A.I.)

A maise crop of L.A.I. 4.0 could therefore hold
5.5 mm of water, a figure close to the average 5.8 mm
found by Stottenburg and Vilson (1950). This is a large
amount of water and itis likely that most of it is held
in the sheaths, particularly at the top of the plant
before tasselling.

(b) Bean cropt

A second supplementary experiment was performed with potted bean plants. In this case interception was estimated as the overall difference in evapotranspiration between plants with "wet" and others with "dry" leaf surfaces.

The method is similar to that used by Paul and Burgy (1961) and Biailar results were obtained by plotting cumulative. Bt against time. Betails of the bean experiment are shown in Appendix 12. The quantity of net interception varied widely between plants, the maximum value recorded being about 50 gm. Taking as average 25 gm per plant, net

interception in the lysimeter srop for row spacing 50 em, seed spacing 15 em within the row, would be approximately 0.4 mm.

The net interception by the bean plant is therefore small and does not explain the high Et/Bo figures unless:

- (1) the intercepted water was evaporated within 1-12 hours under field conditions (compare Appendix 11 for isolated bean plants),
- (ii) there were frequent periods of rain during the day, and
- (iii) the Bo for the day was low, so that net interception constituted a large percentage of daily Bt.

It is, however, more likely that the field bean crop, at full ground cover and with overlapping leaves, retained much more water than the above calculations indicate.

Results from other workers

The effect of intercepted water on crop water use has been investigated by several workers, but in most cases

the experimental conditions were far from representative of a field crop; the experimental plants usually stood in isolation. Burgy and Pomeroy (1958) in nutrient solution studies found that in vigorously growing grass, the evaporation of intercepted water caused an equal reduction in transpiration from the plants, hence total moisture loss was approximately the same in plots with "wet" and "dry" leaf surfaces. Similar results were obtained by McMillen and Rurgy (1960) and McIlroy and Angus (1964) for grass under field conditions. Monteith (1965) suggested that such results with well watered grass are due to low r, the relatively high r making release to unity, and relatively high r making release to unity, and relatively high r making

Summary of discussion on Bt/Ro ratios

The validity of crop water use data obtained with lysimeters has been established. Errors associated with these measurements have failed to account for all the apparent excess crop water use which caused high Et/Eo ratio under wet conditions. The theory relating excessive Et to aerodynamic resistance factors for the two crops has been discussed and advanced as a likely explan-

ation. It has also been shown, at least for the maise crop, that the extra water would be available on the plant. The source of energy for the extra evaporation will be discussed in Chapter IV which deals with partition of energy in the crop.

Validity of estimating crop water

One of the main objectives of this experiment was to evaluate the chances of obtaining a fairly good estimate of actual evaporation under well-defined conditions through soil moisture measurements with a neutron moisture meter, supplemented by gravimetric soil moisture measurements in the top soil layer. With E.A.L. 104 moisture meter, agreement with lymmeter estimates was good during the beam crop, and the 100 mm difference during the 1968 maise crop (see Table 9) is accounted for by the possible error due to systematic sero drift in the lymmeter.

It is now necessary to specify the "well-defined conditions" referred to in the preceding paragraph. Any attempt to deduce crop water use from differences in soil moisture input through irrigation, rainfall or upward flow

from deeper layers in the soil profile and loss through drainage. The poor distribution of water during irrigation at once sets a limit to the reliability of this method in irrigated crops, but eventif this source of error could be eliminated, the drainage component is difficult to determine accurately. Suggestions have been made that it may be possible to calculate drainage correction from records of soil moisture tension at a suitable level in the soil profile. This correction depends on obtaining representative uneaturated conductivity data but, as demonstrated by van Bavel, Stirk and Brust (1968). the variability of such measurements is such that they are unreliable for calculating such corrections. A more serious objection to this procedure arises from the presence of roots. Unless the water potential gradient is measured truly below the root range, any attempts to calculate drainage correction would be invalid. The root mystem of maise is capable of penetrating and drawing water from more than 200 on depth. Although observations indicate that the average maximum soil depth in the experimental plot was about 200 cm, the underlying layer of tuff rock was usually saturated with seepage

water from irrigation canals and could be a source of water for any roots reaching it.

CHAPTER III

DRY MATTER PRODUCTION IN MALZE AND BEAM

The successful application of any photosynthesis model in predicting potential agricultural crop production in a given area depends on the ability of the model to follow increase in total dry matter in all stages of growth. Since grain is the most important end product of maise, several attempts have been made to relate dry matter production and grain yield.

Hanway (1962) found that dry weight of the whole
plant and of the grain were directly related to and
highly correlated with the weight of leaves on the plants.
Thus at maturity, the weight of grain constituted about
one half of the total dry weight of the plant. According to Dale and Shaw (1965), the number of days in the period from six weeks before to three weeks after silking,
on which maise was under no moisture stress, were highly
related to grain yield. The effect of plant density on
the relationship was only apparent when there were more
than 40 days of moisture stress in this period. Regland

et al. (1965), nowever, showed that the rate of dry matter accumulation by maine was roughly linear in the period between three and four weeks after emergence to the onset of semescence.

Kalju Rik and Hanway (1966) have found that Leaf
Area Index at silking time and Leaf Area Index days over
grain formation period are linearly related to grain
yield. The maximum Leaf Area Index for this linearity
was four.

Maximum or potential yield of maise has therefore proved difficult to define, and Earley (1965) suggested the use of "relative maximum yield" of maise expressed as

Maximum weight of maise per unit leaf area of the top three leaves.

This definition was based on the observation that the top three leaves were the most efficient photosynthetically during grain formation. To complicate the situation even further, Ragland et al. (1965) found that the rate of growth d'maise was also correlated with air temperature at 150 cm and soil temperature at 5 cm depth. The usefulness of the transpiration ratio has already been discussed in Chapter I, and it has been shown that only in very special conditions could such a relationship be reproduced (de Wit, 1958; Monteith, 1965; Lemon, 1966). The methods cited above suffer from one defect: they are all empirical and therefore strongly influenced by the behaviour of a given crop growing under particular environmental conditions.

However, potential photosynthesis, like potential evaporation, is governed by meteorological and plant factors. A more logical approach to estimating potential crop productivity could therefore be the measurement and integration of these factors assuming that soil fertility is not limiting. There are three basic steps in this process:

(a) Determination of the light photosynthesis
function - that is the rate of photosynthesis
for different temperatures, light intensities
and caroon dioxide concentrations. This is
usually done with single leaves or whole plants
in controlled environment chambers.

- (b) Determination of light intensity and carbon dioxide concentrations within the camppy of the field crop;
- (o) Calculation of potential dry matter production by combining (a) and (b).

The work described in this chapter was not designed to produce either original information on the above parameters, or a new photosynthesis formula, but to test the usefulness of existing photosynthesis formulae in predicting yields of field erops.

Theoretical background to photosynthesis formulae

The photosynthesis function

Between the lowest (completely limiting) and highest (completely saturating) light intensities, there is a wide range of values in which photosynthesis of leaves in normal air is affected by variation of both light intensity and carbon dioxide concentration (Gaastra, 1959; Chapman and Loomis, 1955).

The light and carbon dioxide response curves are usually characteristic of plant species, but actual values on these curves depend also on the health and age of the leaf being studied. Two of the more recent forms of the photosynthesis functions are considered here.

1.
$$P = (a + \frac{b}{I})^{-1}$$
(5:1)

At a given atmospheric concentration of CO₂, where P is the gross photosynthesis, a and b are constants which may be regarded as proportional to resistances in the diffusion of earbon dioxide, and in photochemical reactions respectively. I is the light intensity at leaf surface.

2.
$$A = H(H + HH)^{-1}Amax$$
.(5:2)

at a given concentration of 00_2 . A is the gross photosynthesis, H is the absorbed light intensity, and HH is the light intensity at which half of the maximum photosynthesis (Amax) occurs.

Light distribution functions

In using (5:1) or (5:2) tosalculate photosynthesis of a field crop, it is necessary to know the pattern of

The photosynthesis function of a plant in the middle of a field crop will depend on such functions for individual leaves which may vary widely because of differences in leaf age and position within the canopy. The situation is complicated further by the fact that the effect of reduced light intensity due to shading of lower leaves is also partially compensated for by a higher efficiency of energy conversion.

Monteith (1965) examined earlier attempts to describe the mean intensity of solar radiation, I_L , below a canopy of Leaf Area Index, L. He found that the assumptions of a constant light extinction factor following Beer's Law,

$$I_{\mathbf{k}} = I_{\mathbf{0}} e^{-\mathbf{k}\mathbf{L}} \qquad \dots (515)$$

and the neglect of the changes in spectral composition of radiation as it filters into the canopy, led to serious errors. Monteith therefore suggested a new formula for radiation extinctions

$$I_{L} = I_{0} \{ s + (1 + s) \gamma \}^{L}$$
 (5:4)

where I is the radiation intensity after I layers of unit Leaf Area Index have been penetrated;

8 is the fraction of incident radiation passing through unit layer of leaf without interception;

T is a mean light transmission coefficient for leaves of the given plants;

I is the incident radiation at the top of the canopy.

Since T is small, the effective radiation for photosynthesis is that intercepted by the leaves which are directly sunlit, and that intercepted by "once shaded" leaves, i.e. the light reaching a lower leaf after only one transmission through an upper leaf.

The area of sumlit leaves can therefore be shown to be

$$A_0 = \frac{1 - 8^L}{1 - 8} \qquad(315)$$

and that of "once shaded" leaves

$$A_1 = \frac{1 - S^L - (1 - S)LS^{L-1}}{1 - S} \qquad \dots (5:6)$$

Monteith used equation (5:1) to derive instantaneous photosynthesis of sunlit leaves:

$$P_0 = (1 - 8)A_0 \left\{ a(1 - 8) + \frac{b}{1_0} \right\}^{-1}$$
(517)

and for "once shaded" leaves

$$P_1 = (1-8)A_1\{a(1-8) + \frac{b}{1-5}\}^{-1}$$
(5:8)

so that the total P is given by

$$P_{n} = P_{n} + P_{1}$$
 (519)

Daily totals may be found by integrating (5:9) over the daylight hours. Assuming that solar radiation varies sinusoidally between sunrise and sunset, the daily total of solar radiation S is then given by

$$R = \int_{0}^{h} I_{o} dt = 2 I^{*} \frac{h}{\pi}$$
 (5:10)

where R is the total solar radiation measured by

a solarimeter;

h is the daylength (min)

(cal.on⁻²min⁻¹).

Therefore the final equation derived by Monteith takes the form

$$P_{m} = \frac{h}{a} \left[A_{o} \{ 1 - 2(n_{o}) \} + A_{1} \{ 1 - 2(n_{1}) \} \right]$$
(5:11)

where $\{1 - f(n)\}$ is a function of I° and may be regarded here as a saturation efficiency.

The above derivation of (5:11) is only meant to highlight the major steps; the reader is referred to the original paper by Monteith for the detailed analysis and, more important, for the assumptions and conditions implied at each stage.

Methods and Procedure

The measurements described in this chapter were made on the same site and crops described in Chapter II.

Measurement of Leaf Area Index

Destructive sampling was necessary for the determination of Leaf Area Index. The same samples were, however, used for estimating total dry matter production and because of the large plant population no significant gaps were caused by these samplings.

Measurement of total leaf area was a very laborious task and only two mains and four bean samples were taken

consisted of all plants in an area of 1 square metre for maise (i.e. 5 plants), or I square metre for beans. It was not possible for measurements to be made on the day of sampling. The plants were therefore preserved by dipping the roots in water overnight and all measurements were carried out on the following day. There was therefore a gap of 24 hours between field sampling and actual measurements.

complete leaves, half leaves and stalks. A good sakithcation-was obtained between leaf area (measured by planimeter) and the products of maximum length of the mid-rib
and maximum width (Appendix 15). Leaf area was therefore
calculated from the linear measurements which were simpler
to make. The half-leaves were treated simply as triangles.

It was assumed that only half the maise stem area contributed to interception of direct light at any time, the
contribution being virtually mil at midday when the sun
was everhead. Stem area for beans was considered
insignificant. The sampling interval was 1-2 weeks and
covered all stages of growth.

Light interception in the crops
determination of the factor 8 in (5:4)

Two simple solariseters (Monteith, 1959) were used to measure the fraction of short wave radiation reaching various levels in the canopy. One solariseter was mounted above the crop while the second solariseter was placed on horisontal rails (Fig. 16) at the required height in the crop. Both instruments were connected, via a changeover switch, to a millivolt meter. By reading the top and bottom solariseters alternately, the fraction of radiation intercepted by the crop could be calculated. Spatial averaging within the crop was achieved by placing the bottom solariseter at twenty-one different positions on the rails at each setting.

The Leaf Area Index associated with each light interception was obtained by sampling the crop in layers and measuring L.A.I. for each layer. This method was successful with the maise crop, but only one setting was possible in the bean crop and the measurements may not be representative.



Fig. 16: Solarimeter on rails as used to estimate radiation interception in the maize crop

The transmission coefficient τ in (5:4) was taken as 5.5% (Yosum and Lemon, 1964) for maise. The same value of τ was assumed for beans in the absence of published data. Duily totals of solar radiation were obtained from a Gunn-Bellani radiation integrator (see Chapter II).

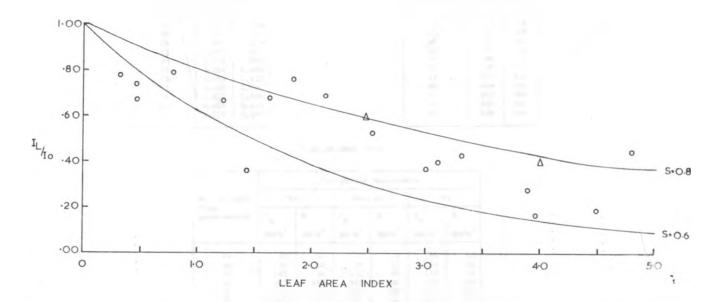
Pollowing Monteith's (1965) calculations based on published photosynthesis data, values of a = 0.25, b = 0.05 were assumed for maise and a = 1.0, b = 0.05 for beans.

A sample calculation of P_m is shown in Appendix 14.

Daily values of L.A.I. were obtained by interpolation on the L.A.I. curves.

Measurement of total dry matter

The dry weight of above-ground parts of the crop was obtained by cutting up and drying separately the leaves, stems, cobs, tassels, pods and grain in a ventilated oven at 70°C to constant weight. The optimum drying time was found to be 36 hours for leaves and 48 hours for stalks,



Pig. 17: Fraction of solar radiation I /I penetrating the camopy of a maise crop at different Leaf Area Indices.

Comparison between measured (M) and calculated (P) total dry matter production from planting to maturity

(a) MAIZE CROP 1967

No. of days from germination	P gm/m ²	m gm/m ²
17 32 38 46 52 59 67 73 81 94 102	38 324 560 965 1,271 1,714 2,254 2,677 3,200 3,917 4,414 4,759	15 106 323 336 657 889 1,484 1,426 1,560 1,896 2,112 2,248

(b) REAR CROP 1968

No. of days from germination	P 2	m/m ²
5 9	2 7	11 15
12	14	18
16	31	41
20	57	43
23	86	49
27	137	63
43	351	225
48	419	285
54	459	305
58	496	321
65	638	406

(c) MAIZE CROP 1968

		Irr	rigation	treatmen	ı ta		
No. of days	We	t	Medi	LINK	Dry		
from germination			P _m	М	P _m	М	
	gm/m²	gm/m ²	gm/m²	gm/m²	gm/m ²	gm/m²	
24	60	12	17	6	50	12	
29	92	32	35	15	74	20	
36	183	98	111	65	144	40	
44	467	165	361	126	350	129	
50	718	453	604	382	548	261	
58	1,152	512	1,035	491	902	376	
65	1,585	818	1,454	605	1,258	425	
72	2,060	1,198	1,906	947	1,649	900	
84	2,856	1,290	2,664	1,219	2,271	1,020	
100	3,876	1,786	3,642	1,684	3,154	1,636	

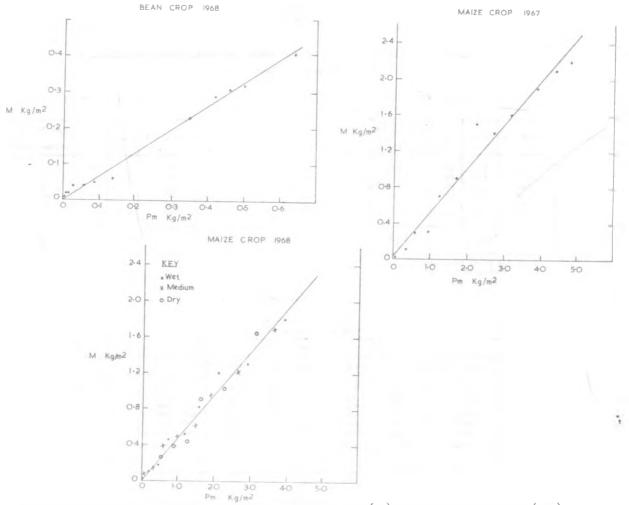


Fig. 18: Relationships between measured (M) and calculated (P $_{\rm m})$ dry matter in the Mwea experiments

Rield of grain (saleable product) compared with total dry matter (M) in make and hann

All figures are in gm/m2.

(a) Maize Crop 1967

No of days flue germination	М	Grain	M - Grain
67	1,484	0	1,484
73	1,426	0	1,426
81	1,560	148	1,412
94	1,896	314	1,582
102	2,112	406	1,706
111	2,248	583	1,665
TOTAL		0.26	

(b) Bean Crop 1968

rom ermination	М	Seed	M - Seed
27	63	0	63
43	225	- 4	225*
48	285	69	216
54	305	109	196
58	321	137	184
65	406	200	206

"Young seed could not be separated from pods.

(c) Maise Crop 1968

	Irrigation treatments											
No. of days		Wet			Mediu	m		Dry				
germination	М	Grain	N - Grain	М	Grain	M - Grain	М	Grain	M - Grain			
65	818	0	818	605	0	605	425	0	425			
72 1	1,198	0	1,198	947	0	947	900	0	900			
84	1,290	26	1,264	1,219	11	1,208	1,020	31	989			
100	1,786	248	1,538	1,684	196	1,488	1,636	268	1,368			
119	2,459	757	1,702	2,005	589	1,416	1,837	662	1,175			
130	2,426	798	1,637	2,310	784	1,526	2,191	781	1,410			

but for maise a correction was applied to the above measurement by assuming that the root component was 20% for young plants up to tasselling and 16% after tasselling. These are approximate values from root/top ratios obtained at Muguga with closely related varieties of maise (Gwynne, 1968).

Besults

Results of light interception studies in the maise crop are shown in Fig. 17. Most values of 8 were between 0.6 and 0.8, and 8 = 0.7 was used in all calculations of P for the maise crop. Measurements in the bean crop indicated a value of 8 close to 0.8. The calculated values (P are compared with measured dry matter (M) in Table 15. The relationship between the predicted and measured total dry matter production (Fig. 18) can be expressed closely by linear equations:

For the beans the equation is

$$H = (0.64 \stackrel{?}{=} 0.02)P_{m} + (5.7 \stackrel{?}{=} 5.2), r^{2} = 0.99, n = 12.$$

For the 1968 mains crop the equation is

$$M = (0.48 \pm 0.01)P_{m} - (5 \pm 21), r^{2} = 0.98, n = 30,$$

which is very similar to that for the 1967 maise crop:

$$H = (0.48 \pm 0.05)P + (34 \pm 68), r^2 = 0.97, n = 12.$$

The difference between the slope for mains and that for beans may be due either to the values assigned to the factor 'a' (proportional to carbon dioxide assimilation) for which there are few field data, or to differences in the fraction used in respiration.

Assuming the only difference between $P_{\rm R}$ and R is the respiration loss R, and that R is a constant fraction K of $P_{\rm m}$,

$$\frac{R}{P_m} = \frac{P_m - R}{P_m} = K$$
, and $M = (1 - K)P_m$ (3:12)

The value of K in this experiment was therefore 0.52 - 0.05 for maise and 0.56 - 0.02 for beans. The

magnitude of the respiration loss is not known with certainty except that it increases from about 25% in cold to 50% in hot climates (Gasstra, 1965). The values of K obtained above are therefore of the right order and could probably be improved by adjustments based on experimental determination of a number of factors, especially the factor 'a'.

Ratio of saleable product to total dry matter

Although total dry matter is economically important where the whole crop is turned into silage, the most important part of maise or beans in East Africa is the grain, and increase in grain yield is a major aim in the cereals breeding programme. It would therefore be of interest to look into the ratio of grain to total dry matter for the varieties of maise and beans used in these experiments.

The data in Table 16 show that during the last 50 days for maise and 20 days for beams, almost all net photosynthesis was channelled into grain formation.

The final grain yields constituted approximately 25-36% total dry matter in maise and about 50% for beans.

Discussion of Results

In spite of the numerous assumptions made in calculations, the estimates of potential photogynthesis (equation (3:11)) were very closely correlated with the actual dry matter production in the field crops of maise $(r^2 = 0.97, 0.98)$ and beens $(r^2 = 0.99)$ in all stages of growth. The success of this approach in the warm climate of the Mwea experiments may be partly due to removal of temperature as a major variable in growth and this result needs to be tested in the cooler climates of the East African highlands. The correlation of dry matter production and leaf Area Index in these experiments was also suite high (r2~ 0.80) and it may be argued that it is not necessary to perform the laborious task of following equation (3:11) to account for an additional 17% of variance. The L.A.I. approach is however empirical

and can only be expected to yield comparable results where light intensity is not limiting.

de Wit's me thod

calculating light distribution functions for different leaf distribution functions (canopy structures). The incident light intensity on a leaf due to direct light is proportional to the sine of the angle (LS) between the leaf and the ray of the sun. In this method, the light distribution function is expressed in terms of sine (LS) for 10° increments of the sun and leaf angles in all combinations. The leaf distribution function, i.e. the frequency distribution of leaf angles in the camppy, is first calculated from direct measurements, and then used to calculate light distribution functions.

In calculating photosynthesis, equation (5:2) is used and instead of using typical values for various factors to minimise laborious calculations, the method assumes a computer is freely available and the best

fitting values are found empirically. de Wit does, however, give a table of light and photosynthesis values for a canopy under standardised conditions, from which photosynthesis at any place and time of year can be calculated. The chosen standard conditions are:

Leavens

Scattering coefficient 0.50

mm 0.56 cal.on min-1

Amex 20 KgCH₂0 ha⁻¹hr⁻¹

Canopy :

Spherical leaf distribution

Canopy density 0.1

Leaf Area Index 5.0

Lights

Clear sky

Inclination of the sun 45°

Direct light 0.484 cal.cm min

Diffuse light 0.092 cal.cm 2min-1

CO .

Aerodynemic diffusive resistance r

0.5 sec.em-1

de Wit suggested that likely variations from most of the above values would have little effect on the ealeulated photosynthesis, but the choice of L.A.I. of 5.0 implies complete canopy. The application of this table of standard values has already been discussed in the introductory chapter.

In attempt was made to determine the light distribution function in the maise crop by the system of two
spheres suggested by de Wit (1965, loc. cit.). Results
were inconclusive and failed to confirm whether the maise
canopy was "spherical" as suggested by Nichiporovich
(1961) or plagiophile, as found by de Wit (1965 loc. cit.).
It is therefore not surprising that photosynthesis
calculated on the basis of the table of standard values
did not yield values comparable to observed dry matter
production at the beginning and towards the end of the

growing season. The method should, however, be applicable to campies of perennial crops like ten and coffee previded all the relevant parameters can be determined in situ.

CHAPTER IV

THE PARTITION OF SCHAR ENERGY IN PIEID CROPS OF MAIZE AND BRANS

Evaporation and photosynthesis are both dependent on energy, the principal source of which is the incoming solar radiation. The partition of solar energy in the field crop is therefore an important factor in studies of crop water use efficiency. The experiments described in this section had two main aims:

- 1. To obtain some indication of the bourly distribution of the major energy components in field crops of maise and beans;
- 2. To compare the latent heat component on a daily basis with lysimeter observations of crop water use.

Methods and Procedure - 1: Derivation of Bowen's Ratio from gradients of temperature and humidity above the evaporating surface

The basis equations for the transfer of heat, water vapour and momentum ares

Sensible heat:
$$Q = -\frac{\partial Q}{\partial p} \frac{dT}{ds}$$
(4:1)

Water vapour:
$$B = -e^{-\frac{dq}{ds}}$$
(4:2)

Momentum:
$$N = -eK_m \frac{du}{ds}$$
 (4:5)

where q = specific humidity (gm. water vapour/gm. moist air) $K_{v}, \quad k_{v} = \text{eddy transfer coefficients } (cm^{2} \text{sec.}^{-1}),$ and all other terms retain their meaning in Chapter 5.

Bowen's ratio is defined as

$$\beta = \frac{Q}{\sqrt{B}} = \frac{Q}{\lambda} \frac{\frac{QT}{D}}{\frac{D}{D}} \frac{\frac{QT}{D}}{\frac{dQ}{D}} \dots (484)$$

If q is expressed in terms of dry air, it becomes the mixing ratio x where

$$x = \frac{\epsilon_0}{2 - \epsilon_0} \qquad \dots (415)$$

• - Vapour pressure (mb)

P = atmospheric pressure (mb)

and for o &P

$$x = \frac{\epsilon_0}{P}$$
and $\beta = \sqrt[K_h]{\frac{dT}{dz}}$

$$\frac{dT}{dz}$$

$$\dots (416)$$

The values of K_h and K, have been shown to be identical under stable and neutral conditions (Pasquill, 1949; Swinbank, 1955; Taylor, 1960; Crawford, 1965).

Under unstable conditions, variations in the ratio K_h are not significant (Rider and Robinson, 1951; Suomi and Tarmer, 1958), and there is evidence to show that the ratio is close to unity. Equality of K_h and K_y is therefore assumed in the analysis of data from the following measurements.

The two quantities $\frac{dT}{ds}$ and $\frac{de}{ds}$ should be measured at the same point above the evaporating surface.

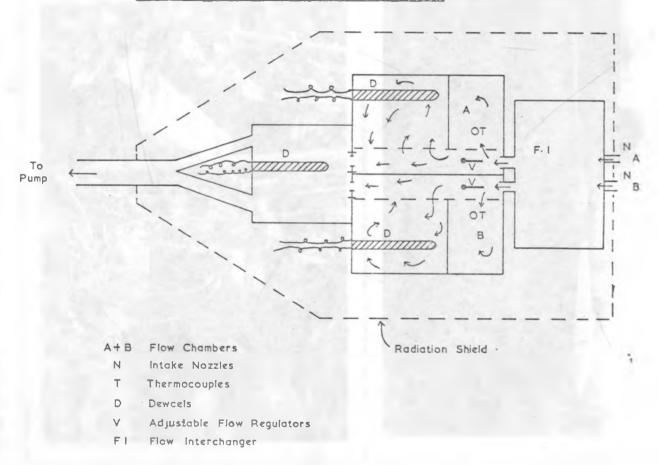
Apparatus

The system used was similar to that of Fritschen (1965), and the equipment was designed and constructed by

B. A. Ripley at Muguga. Figs. 19 and 20 show the components, construction and operation of this equipment.
To improve spatial averaging of the measurements, a
hydraulic system was used to swing the entire apparatus
to and fro horisontally over an angle of 160 degrees.
The period of each swing was sometimes affected by the wind
but was on the average two minutes. This system of spatial
averaging was only possible over the low bean crop and
sould not be operated 5 m above the ground over the tall
maise crop.

the two nossles with a vacuum pump. The two air streams were kept separate as they passed through chambers A and B for measurements of humidity and temperature differences. The difference in dewpoint was measured by a differential thermocouple system attached to lithium chloride dewcels (Tanner and Suemi, 1956). Two thermocouples, soldered on to copper tubes to increase time constant, were used to measure the temperature difference, dT. The two air streams were then mixed in a final chamber containing the third dewcel for the determination of mean dewpoint,

FIG. 19
Diagram Of Bowens Ratio Equipment — not to scale



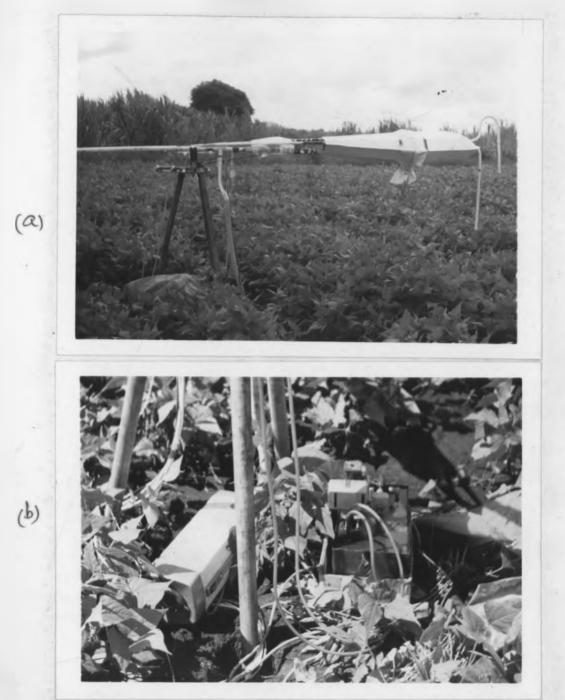


Fig. 20: Operation of Bowen's Ratio equipment. (a) rotating boom above bean crop, (b) details of tripod support, vacuum cleaner for sucking air, and hydraulic system for rotating boom.

Td. The vapour pressure difference, de, was obtained from Td and the difference in dewcel element temperatures, D.E.T. (see Appendix 15). The response times for the dewcels and thermocouples were 4 and 1-12 minutes respectively.

Continuous recording

Because of the rather long time taken by the dewoels to reach equilibrium, the time interval between flow changes was fixed at 10 minutes. The differential thermocouple outputs were first passed through a D.G. amplifier which changed the range of a 12 channel Leeds and Morthrup centre zero recorder to - 50, 100, 200, 500 and 1,000 at as required. Unfortunately there was no provision in the amplifier for separate amplifications for different inputs. The maximum amplification chosen was therefore dictated by the rather large absolute temperature differences between the two dewoels.

Fig. 21 shows a typical record of these parameters for a chart speed of 6 inches/hour. The output from

therefore 10 readings of each parameter for each 10 minute period. Avoiding the first four readings, corresponding with equilibration period, and the last reading before air interchange, the average reading for each 10 minute period was calculated from the remaining five readings.

Mean dewpoint 74 was recorded directly on a separate calibrated Leeds and Northrup 16 channel recorder.

All recorders, timer and power stabiliser were located in a mobile meteorological laboratory. Power supply was from a portable 1.5 KW generator.

Elimination of errors by interchanging airstreams

The vapour pressure and temperature differences over 50 cm are quite small (~ 0.01 mb and $0.05^{\circ}0$ respectively), and errors introduced by differences between sensors can therefore be very important. To overcome this source of error, one can either interchange the sensors regularly or interchange the airflow between fixed sensors. The

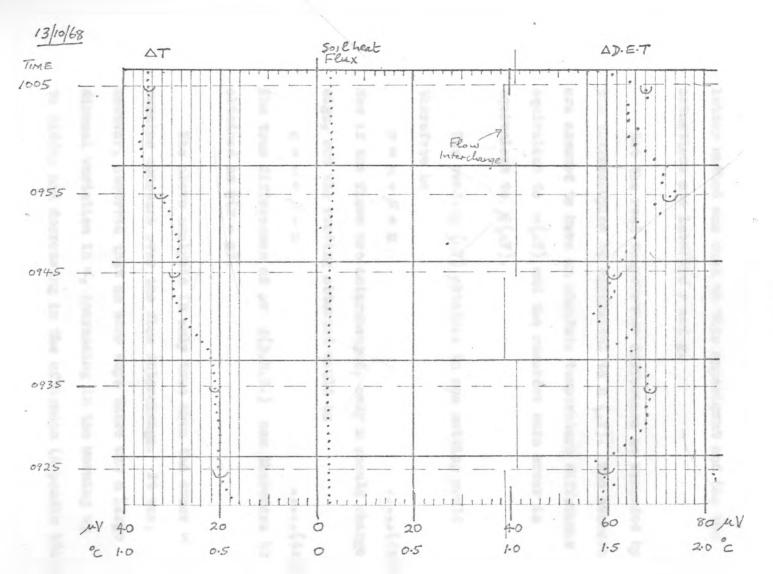


Fig. 21 A part of chart record (actual size) of measurements with Bowen's Ratio equipment above maize prop on 13/10/68.

latter method was used in this experiment and the two situations are labelled r and g.

Let the real temperature difference as recorded by the thermocouples be equivalent to $x(\mu V)$. The sensors are assumed to have an absolute temperature difference equivalent to $x(\mu V)$ and the recorder zero error is equivalent to $\beta(\mu V)$.

The reading (AV) obtained in one setting would therefore be

$$\mathbf{r} = \mathbf{x} + \boldsymbol{\beta} + \mathbf{z} \qquad \dots (418)$$

Now if the flows are interchanged, only x should change sign, and the reading obtained would be

$$\mathbf{g} = \mathbf{x} + \boldsymbol{\beta} - \mathbf{z} \qquad \dots (419)$$

The true differences dT or d(D.E.T.) can therefore be obtained as $\frac{1}{2}(r-g)$.

The above analysis is only true when the error of remains constant over the flow interchange. It was, however, observed that on some days there was a large diurnal variation in of, increasing in the morning up to midday and decreasing in the afternoon (Appendix 16a).

This could have been due to solar radiation affecting one chamber more than the other, in spite of the radiation shield. The analysis in Appendix 16b shows how this error has been reduced by averaging readings over two rather than one stream interchange.

solar radiation, not radiation and soil heat flux

Fotal short wave solar radiation and not radiation were measured with an Eppley pyrheliometer and Funk not radiometers respectively. Calibrations of these instruments were obtained by comparing the vertical component of direct radiation with the Eppley model of the Angetrön electrical compensation pyrheliometer. The sensitivity of the Eppley pyrheliometer No. 4284 varied slightly with solar angle, being lowest in the morning before 9 a.m. and in the evening after 5 p.m. but relatively constant for the hours in between. The average for the day was 6.62 mv.min.Ly⁻¹. Calibration of the Funk net radiometer No. 470 was 8.85 Mv.min.Ly⁻¹.

Solar radiation and not radiation were recorded separately on two Kent Millivolt recorders with a range -5 to +15 mv. The recorders were fitted with mechanical integrators and the total radiation for a given period could be read off the totalising counter or on the automatic hourly printout on EMM counters. Radiation values obtained this way agreed well with planimeter integration of chart records.

heat flux plates (Deacon, 1950) buried horisontally 1 on below the soil surface and arranged so that one plate was halfway between the rows, the second plate was fof the distance between 2 rows, and the third plate was between plants in the row. The plates were installed one week before measurements commenced and they were connected in series to improve spatial averaging. The average sensitivity of the plates in series was 11.5 mv.min.Ly⁻¹. The output was continuously recorded on the same chart as the temperature and humidity gradients discussed in the previous section, but without amplification. Soil heat flux was calculated from the average chart reading for each 10 minute period.

Methods and Procedure - 3: Energy storage in shotosynthesis

The energy used in photosynthesis is only a small fraction of daily radiation; the fraction of energy stored in dry matter is even smaller, and though almost insignificant in comparison with other components of surface energy balance, it is the basis of crop production. An estimate of stored energy can be made by converting dry matter increments to energy equivalent per unit land area. In accurate work, it would be necessary to split up the total dry matter into the major chemical components, i.e. carbohydrates, fibre, protein and fat, and then calculate energy equivalent from the data reported by Maynard (1947) as follows:

Fibre and carbonydrates 4.15 K cal/gr

Pats 9.40 K cal/gm

Protein 5.65 K cal/gm.

The maise data in Table 17, compiled from Morrisson (1956), show that about 77% of dry matter comprises

fibre and carbohydrates. The figures for beans were obtained at Muguga. The energy equivalent of dry natter for the beans was found to be 4.25 K cal/gn. In view of the rather large variability in field samples and the attendant errors in the measured dry matter, the value 4.2 K cal/gn was used for both maiss and beans in the conversion of dry matter to energy equivalent.

TABLE 17
Chemical composition of mains and beans
2 D.M.

Maise	Cartohydrates	Grude Pibre	Crude Protein	Pat
Stover (sten + leaves)	46.5	50.8	5.9	1.6
Grain	69.2	2.0	8.7	4.0
Gobs	54.0	32.1	2.3	0.4
Complete ears (grain + husks + cobe)	65.3	10.5	7.8	3.0
Beans	49.1	19.3	14.5	19.5

TABLE 18a

Energy belance (cml.cm⁻²) over bean crop at lives 6.5.68

Period (local time)	Solar	Net	Soil heat flux s	H - 8	$\beta = \frac{Q}{\lambda E}$	1 +β	λE	Q	Net Solar	Albedo
0700-0800	10.1	7.1	-0.7	7.8	_	1/2	-	-	0.70	0.30
0800-0900	17.8	12.7	0.6	12.1	-	-	-	-	0.71	0.29
0900-1000	31.4	23.5	1.5	22.2	0.05	1.05	21.1	1.1	0.75	0.25
1000-1100	65.5	48.5	2.3	46.2	-0.01	0.99	46.6	-0.4	0.74	0.26
1100-1200	77.0	67.8	3.7	64.1	0.00	1.00	64.1	0.0	0.88	0.12
1200-1300	82.9	72.0	5.1	66.9	0.07	1.07	62.5	4.47	0.87	0.13
1500-1400	86.4	74.9	4.0	70.9	0.09	1.09	65.1	5.8	0.87	0.13
1400-1500	74.8	66.7	2.5	64.2	0.16	1.16	55.3	8.9	0.89	0.11
1500-1600	35.6	30.5	-0.2	30.7	-0.03	0.97	31.7	-1.0	0.86	0.14
1600-1700	39-3	52.8	-0.1	32.9	0.03	1.05	31.9	1.0	0.83	0.17
1700-1800	3.0	0.7	-1.2	1.9	0.06	1.06	1.8	0-3	0.25	0.77
Total	523.8	437.2	17.3	419.9		Equiv.	380.2	19.9		
						(=):	6.5		Meant	0.19

*excluded from the mean

TABLE 18b

Trerry belance (cal.cm⁻²) over bean crop at Mue.

Period (local time)	Solar	Net N	Soil heat flux	N - 3	$\beta = \frac{Q}{\lambda E}$	1+β	λB	Q	Net Solar	Albedo
0900-1000	27.8	17.5	0.7	16.6	0.10	1.10	15.1	1.5	0.62	0.38
1000-1100	37.6	26.9	1.2	25.7	0.01	1.01	25.5	0.2	0.72	0.28
1100-1200	65.7	51.6	2.4	49.2	-0.02	0.98	50.2	-1.0	0.79	0.21
1200-1300	73-7	60.5	4.5	56.2	-0.02	0.98	57.3	-1.1	0.82	0.18
1300-1400	81.1	67.0	4.3	62.7	-0.08	0.92	68.1	-5.4	0.83	0.17
1400-1500	70.5	52.9	3.0	49.9	0.09	1.09	45.8	4.1	0.75	0.25
1500-1600	57.1	45.5	1.4	44.1	0.08	1.08	40.8	3.3	0.80	0.20
1600-1700	33.4	25.3	0.0	25.3	0.19	1.19	21.3	4.0	0.76	0.24
1700-1800	16.9	4.3	-1.0	5.3	0.10	1.10	4.8	0.5	0.25	0.75
1800-1900	0.6	-3.5	-1.2	-2.5	0.07	1.07	-2.1	-0.2	-	-
Totals	464.4	347.8	15.1	332.7		Equiv.	326.8	5.9		
						(mm):	5.6		Meant	0.21

"excluded from the mean

TABLE 19a

Energy balance (cal.cm⁻²) over maize crop at Mwea 20.12.67

Period (local time)	Solar	Het	Soil heat flux 8	1 - 8	$\beta = \frac{Q}{\lambda R}$	1 + β	AB	Q	Het- Bolar	Albedo
0900-1000	67.9	49.1	0.8	48.3	0.07	1.07	45.2	3.1	0.72	0.28
1000-1100	76.6	63.2	1.1	62.1	-0.04	0.96	64.7	-2.6	0.82	0.18
1100-1200	87.9	72.5	1.2	71.1	-0.06	0.94	75.6	-4.5	0.82	0.18
1200-1300	87.0	74.4	1.2	75.2	-0.07	0.93	78.8	-5.6	0.86	0.14
1300-1400	83.4	71.2	2.0	69.2	-0.07	0.93	74.5	-5.5	0.85	0.15
1400-1500	66.4	58.3	1.1	57.2	-0.01	0.99	57-7	-0.5	0.88	0.12
1500-1600	47.8	40.9	0.5	40.4	-0.01	0.99	49.8	-0.4	0.86	0.14
1600-1700	29.8	22.0	0.2	21.8	-0.02	0.98	22.2	-0.4	0.74	0.26
1700-1800	9.2	-2.2	-0.1	-2.1	0.05	1.05	-1.9	-0.2	0.23	0.77
Total:	556.0	449.2	8.0	441.2		Equiv.	457.6 7.8	-16.4	Moant	0.18

excluded from the mean

TAHLE 19b

Inergy balance (cal.cm-2) over maise erop at Muca

21.12.67

Period (local time)	Solar	Het H	Soil heat flux S	N - 8	$\beta = \frac{Q}{\lambda B}$	1 + \$	AB.	Q	Not Solar	Albedo
0830-0930	52.5	44.5	0.6	43.9	0.00	1.00	43.9	0.0	0.85	0.15
0930-1030	69.4	60.1	1.0	59.1	-0.03	0.97	60.9	-1.8	0.87	0.13
1030-1130	75.4	66.8	1.3	65.5	-0.06	0.94	69.6	-4.1	0.89	0.11
1150-1250	96.5	87.8	1.2	86.6	-0.06	0.94	92.2	-5.6	0.91	0.09
1230-1330	82.0	76.9	1.7	75.2	-0.04	0.96	78.5	-5.1	0.94	0.06
1350-1450	63.0	61.8	1.7	60.1	0.03	1.03	58.5	1.8	0.98	0.02
1430-1530	56.7	45-5	0.7	44.8	-0.02	0.98	45.7	-0.9	0.80	0.20
1530-1630	39.1	51.1	0.4	30.7	0.03	1.05	29.8	0.9	0.80	0.20
1630-1730	18.1	9-3	0.1	9.2	0.04	1.04	8.9	0.3	0.51	0.49
1730-1800	-	-2.3	-0.1	-2.2	0.10	1.10	-2.0	-0.2	-	-
Total:	552.7	481.5	8.6	472.9		Equiv.	485.6	-12.7	Meant	0.21

"excluded from the mean

Energy balance (cal.cm-2) over maise crop at Hwea

TABLE 20a

11.10.68

No record of soil heat flux. All therefore overestimated by at least 25.

Period (local time)	Solar	Net	$\beta = \frac{Q}{\lambda E}$	1 +β	λz	Q	Bet	Albedo
0730-0800	16.5	11.1		-		_	0.67	0.52
0800-0900	26.1	20.4	0.10	1.00	20.4	0.0	0.78	0.22
0900-1000	56.2	43.8	-0.16	0.84	52.1	-8.5	0.78	0.22
1000-1100	84.5	72.5	0.06	1.06	68.4	4.1	0.86	0.14
1100-1200	88	78.1	0.42	1.42	55.0	3.1	0.87	0.13
1200-1300	88.9	77.3	-0.22	0.78	99.1	-21.8	0.87	0.13
1300-1400	87.9	76.5	0.31	1.51	58.4	18.1	0.87	0.13
1400-1500	62.2	53.5	0.87	1.87	28.6	24.9	0.86	0.14
1500-1600	46.4	37.1	1.80	2.80	13.3	23.8	0.80	0.20
1600-1700	21.6	16.2	-0.20	0.80	20.3	-4.1	0.75	0.25
170 -1800	26.1	5.5	0.30	1.30	4.2	1.3	0.21	0.79
Totals	606.0	492.0	Y	Equiv.	419.8	41.1		7
				(mm):	7.2		Weant	0.18

excluded from the mean

TABLE 20b

Energy belance (cal.cm²) over maise crop at Nues - 12.10.68

Period (local time)	Solar	Net N	Soil heat flux 8	H - S	$\beta = \frac{q}{\lambda E}$	1 +\$	λg	Q	Net Solar	Albedo
0800-0900	-	19.5	0.9	18.6	-0.22	0.78	23.8	-5.2	-	-
0900-1000	69.0	53.8	2.0	51.8	0.47	1.47	35.2	16.6	0.78	0.22
1000-1100	69.8	60.0	2.1	57.9	-0.42	0.58	99.8	-41.9	0.86	0.14
1100-1200	82.9	72.1	2.5	69.6	0.46	1.46	47.7	21.9	0.87	0.13
1200-1300	91.8	79.9	4.1	75.8	0.32	1.32	57.4	18.4	0.87	0.13
1300-1400	86.2	75.0	5.0	70.0	-0.17	0.83	84.3	-14.3	0.87	0.13
1400-1500	60.1	51.7	2.0	49.7	0.43	1.45	34.8	14.9	0.86	0.14
1500-1600	38.7	31.0	1.4	29.6	0.57	1.57	18.9	10.7	0.80	0.20
1600-1700	21.9	16.4	1.1	15.3	0.13	1.13	13.5	1.8	0.75	0.25
1700-1800	7.1	1.5	0.7	0.8	0.06	1.06	0.8	0.0	0.21	0.79
Totals	527.5	460.9	21.8	439.1		Equiv.	416.2	22.9	ą	
	1			T see		(mm):	7.1		Means	0.17

"excluded from the mean

200

Sherin belance (==1.cm 2) ever makes even at Mean

3,10,68

Perties (Joseph	Pales.	of m	288-	•	N N N N N N N N N N N N N N N N N N N	8	2	ø	923	41146
700-0000	6.3	4.2	0.1	4.3	0.08	3.05	5.0	0.5	0.67	0.53
0060-10	27.2	Z	6.0	21.0	0.8	8	10.0	10.4	0.62	0 13
0001-0060	50	9.8	1.3	57.5	0.46	1.46	25.5	11.0	0.75	0.24
0011-000	74.6	61	9.7	8.65	0.30	1.30	46.0	15.8	0.63	0.17
100-1200	38.6	917	2.4	7.0	0.0	1.00	73.1	9	0 83	0.17
1201-1300	96-0	100	6.1	76.6	9.0-	0.60	127.7	-51.1	0.84	0.16
1300-1400	711.4	6.8.0	4.1	62-69	-0.42	2.0	110-1	-46.2	90.0	0.14
40 -1500	72.5	67.9	2.3	9 %	0.13	1.13	52.7	3	0 85	0.15
900-1600	50-0	32.1	1.7	37.4	0.35	1.35	27.7	P 55	0 76	0.22
600-1700	36.7	27.1	7.7	26.0	-0.01	0.93	28.0	-2.0	0 74	0.26
200-1800	13-3	2.7	0.1	2.0	-0.01	0.99	2.0	0.0	0.19	0.61
Totals	6-509	487.2	20.5	466.7			507.2	5-04-		
							6.7		No.	0.18

exeluded from the mean

TABLE 20d

14.10.68

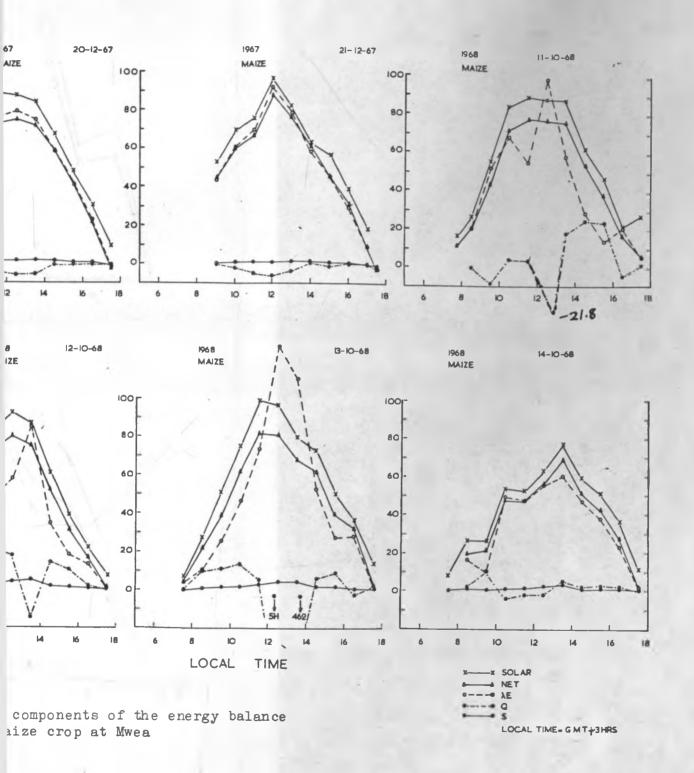
Period (local time)	Solar	Net	Soil heat flux 5	8 - 8	$\beta = \frac{Q}{\lambda E}$	1 +β	λz	Q	Ret	Albedo
0700-0800	8.9	•	0.4	-	0.58	1.58	_	•	-	-
0800-0900	26.7	20.0	1.0	19.0	0.14	1.14	16.7	2.5	0.75	0.25
0900-1000	26.6	21.6	0.8	20.8	1.01	2.01	10.3	10.5	0.81	0.19
1000-1100	53.8	47.8	1.6	46.2	-0.07	0.93	49.6	-5.4	0.89	0.11
1100-1200	52.4	47.6	1.8	45.8	-0.04	0.96	47.7	-1.9	0.91	0.09
1200-1500	63.0	56.2	2.4	55.8	-0.03	0.97	55.5	-1.7	0.89	0.11
1300-1400	77-3	69.2	3.3	65.9	0.09	1.09	60.5	5.4	0.90	0.10
1400-1500	59.0	51.2	1.8	49.4	0.05	1.05	47.0	2.4	0.87	0.13
1500-1600	51.2	42.6	1.6	41.0	0.07	1.07	38.3	2.7	0.83	0.17
1600-1700	36.2	27.3	1.2	26.1	0.12	1.12	23.3	2.8	0.75	0.25
1700-1800	11.8	2.7	0.7	2.0	0.24	1.24	1.6	0.4	0.25	0.77*
Totals	466.9	386.2	16.6	370.0		Equiv.		19.5	Hean:	0.16

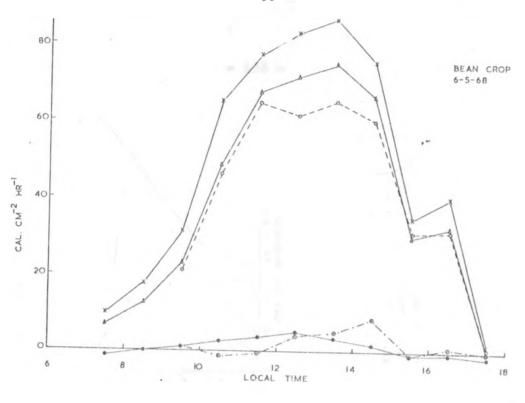
"excluded from the mean

TARLE 21

Comparison between evaporation derived from energy balance measurements in field A and lusinater A estimate of crop uniter una

		Bt	Rt				
Date	Crop	Igeineter A	Energy bolance	Period			
6.5.68 7.5.68	Beans	5.4	6.5 5.6	0900 –1 800			
Total		9.7	12.1				
11.10.68 12.10.68 13.10.68 14.10.68	Matse	6.9 7.5 8.0 8.9	7.2 7.1 8.7 6.0	0800-1800 0800-1800 0700-1800 0800-1800			
Total		31.3	29.0				
20.12.67	Maise	7.5 9.7	7.6 8.3	0900 -18 00 08 30-18 00			
Total		17.2	16.1				





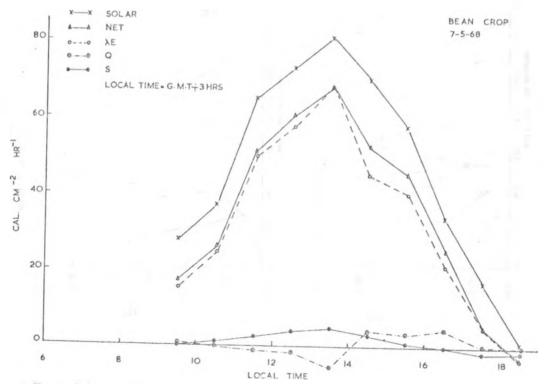


Fig. 23: Hourly components of the energy balance over bean crop at Mwea

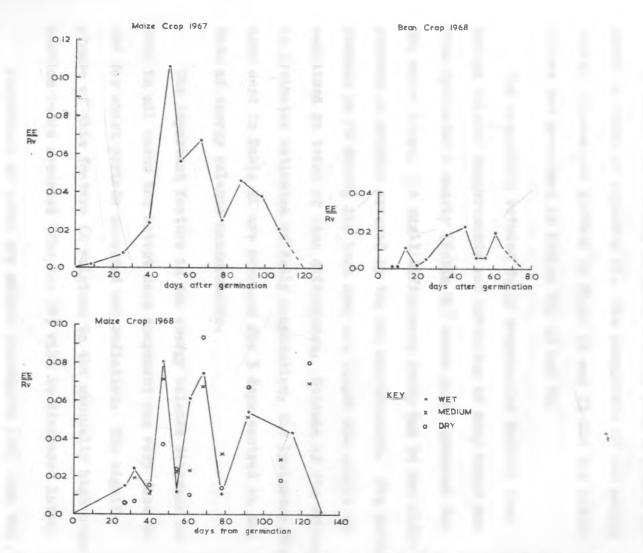


FIG. 24: Fraction of solar radiation (0.4 - 0.7) stored in dry matter at various stages of crop growth

Results

A sample calculation of Bowen's ratio from the data recorded is shown in Appendix 17. The components of the hourly emergy balance are plotted in Figs. 22 and 25 and the actual values are presented in Tables 18, 19 and 20.

The comparison between daily evaporation derived from energy balance measurements and estimates of crop water use from Lysimeter A nearby (Table 21) shows good agreement for the maise crops, the small differences being due to the differences in periods covered by the two measurements. The differences in Et during the bean crop are bigger and cannot be explained in terms of time differences. Because of the errors in lysimeter estimates of daily evaporation, the lysimeter figures used in Table 21 are averages for 5 days centred on the date of energy balance measurements.

One important feature in the energy balance diagrams is that in all cases the latent heat component was very close to and for short periods exceeded net radiation. The significance of this feature in relation to the unusually high evaporation rates recorded (Chapter II) will be discussed later.

Conversion of total dry matter increments (AN) into the equivalent energy, and the comparison of these values with the fraction of total short wave radiation in the photosynthetic $(0.4-0.7\mu)$ wavelength (approximately 0.47), are

shown in Appendix 18. The results (Fig. 24) show that the fraction of light energy stored in dry matter per unit area was higher in maise (0 - 8%, seasonal average $\simeq 4\%$) than in beans (0 - 2%, seasonal average $\simeq 1\%$). The corresponding fractions of total short wave radiation are 1.9% and 0.5% for maise and beans respectively.

Discussion of results

The maximum error in the measurement of solar and net radiation was 3%, all of which was in the calibration and integration. Soil heat flux measurements were less precise and although errors in calibration and integration were no more than 5%, the representativeness of the measurements could not be assessed and the spatial averaging by the use of three flux plates in series may not have been sufficiently effective in smoothing out effects of differential shading by plants and by surface clods. It has been shown (Tanner, Peterson and Love, 1960) that for short periods during the day soil heat flux can be a large proportion of net radiation. The values of soil heat flux recorded in these measurements (highest 5 - 8% of net radiation) are therefore no more than rough estimates, but it is considered that with evapotranspiration rates near potential, the errors in \ F from this source were less than 10%.

The major source of error in the calculation of λB was in the Bowen's Ratio where even without considering errors due to inadequate spatial averaging, the temperature and humidity differences were so small that it was necessary to use a rather elaborate system of flow interchanges. Differences of more than 100% between successive 10-minute estimates of β were common, but probably genuine variations.

The agreement obtained between energy balance and lysineter estimates of crop water use in these experiments is therefore remarkably good and consistent, increasing confidence in the working of the lysimeter and the high Bt/Eo values recorded (Chapter II).

Fig. 24 shows two large drops in photosynthetic efficiency of the two maise crops. The dips are less in the
1967 crop but occur at the same stages of crop growth, 55 and
75 days after germination, in both crops irrespective of water
treatments. In the absence of adverse soil conditions there
are at least two possible causes:

- (a) that these changes are entirely due to experimental error in dry matter determination; in which case the occurrence at similar stages of growth was purely a coincidence.
- (b) that these drops represent genuine reduction of

photosynthetic efficiency due to crop physiological factors.

Since all plant semples were randomly selected and treated in a similar manner, it is unlikely that the experimental error exceeded 20%. This is much less than the 50-90% drops in Fig. 24, but could serve to explain the rather high peaks of energy recovery which approach the theoretical limit of (Monteith, 1965c). The seasonal averages are however larger than the usual 1-2% quoted for agricultural production (Monteith, 1965d; Huxley, 1965). Alternative (b) above is therefore a possibility. The two periods coincide with tasselling and the start of the grain filling respectively, but apart from possible reduction of light penetration by the inflorescence, the cause of this phenomenon, which is not reflected in the linear growth curves (Figs. 12 and 15), is not immediately obvious.

In the case of the bean crop, the drop between 50 and 60 days after germination (Fig. 24) coincides with 212 mm of rain in 10 days and 60 mm in the following 5 days. Since maximum available water in the depth 0-150 cm is 250 mm (Table 1) the addition of the 212 mm to an already wet profile could have led to soil saturation; anaerobic conditions, however temporary, could have resulted in reduced rate of dry matter formation. The canopy was already fully developed for this effect to show on the leaf area index curve.

CHAPTER Y

CONSTITUTE AND SUGGESTION FOR FURTHER VOICE

1. Crop Water Use

The hydraulic weighing lysimeter of the type described by Foregate, Hosegood and McCulloch (1965), if properly sited and operated, can provide reliable estimates of crop water use in all stages of crop growth. The accuracy of these estimates, however, diminishes rapidly as the time interval is reduced below 5 days. The main cause of this seems to be a variable response time apparently dependent on both the magnitude and spatial distribution of the change in load on the lysimeter. The exact mechanism resulting in erratic movement of the water level in the balancing column remains to be investigated. Sensitivity of the weighing system can be increased but the benefit would be minimal unless the drift in solumn height can be climinated.

The patterns of crop water use presented in Chapter II, and the subsequent discussion, have however revealed two major considerations much more important than the proper operation and reliability of lysimeters:

- (a) the difficulty in growing, in a lysimeter, a crop truly representative of the surrounding field crop, and the proper corrections to be applied if this condition is not satisfied;
- and (b) the importance of the frequency of rainfall or irrigation, for a valid extension of crop water use data to environmental and agronomic conditions different from those obtaining at the experimental plot where the measurements were originally made.

No proven answer has been found for (a), but for (b) comparison between the patterns of crop water use for the 1967 and 1968 maise crops suggests strongly that the occurrence of high Bt/Ro values when rainfall is frequent is due to direct evaporation of water intercepted by the crop. The magnitude of this effect is governed mainly

by the ratio of crop surface resistance r to aerodynamic resistance r. There are very few published values of these resistances and more work is required to determine representative values of these factors. Various methods of determining r are discussed by Szeics and Long (1969) whose prediction of large seasonal changes in r for the forest canopy in Kericho, Kenya, appears to be well supported by the bimodal rainfall pattern. Until these factors can be determined with certainty, extrapolation of crop water use measurements between different environments should be carried out with caution and with full realisation that in the absence of supporting meteorological data, the predictions of Et/Ro may be entirely unrealisatio.

Fortunately, agriculture and land use planning do not have to wait until the above corrections can be applied rigorously in every case. Results of this work suggest strongly that in irrigated crops substantial water use efficiency can be achieved by applying fewer but reasonably heavy irrigations direct to the soil surface, notwithstanding the difficulties of uniformity

of water application, and the water storage capacity of the soil profile within root range.

A simpler but interim application of these results is is the incorporation of the number of rain days in the prediction of crop water use. Promising results of this approach have been obtained by Dagg (1970) in the case of tea where an equation of the form

$$\frac{2t}{E_0} = a_1 t + a_2 n(1 - t),$$

where a = a = 0.90

f = fractional number of min days per month,

has given values of Et/Eo closely correlated with measurements of Et/Eo with a lysimeter. The main response to

rain days in the case of young tea is in evaporation from

bare soil, suggesting that — for tea has a value close

to unity. This may not be the case with other crops and

the presence of empirical factors like a and a is still

unsatisfactory. The general usefulness of this formula

should however be checked where accurate measurements of

crop water use and percentage ground cover are available.

2. The use of Neutron Moisture Meters

Of the two makes of neutron moisture meter tested, the H.A.L. and the HIV-I, the former has proved more accurate and better suited to detection of small changes in moisture content in soil with large water holding capacity. Under the right conditions, these changes in moisture content can readily be used to calculate a reasonable estimate of crop water use in the field, but in irrigated fields, uniformity of water application is an important source of error. Another limiting factor is the lack of reliable methods of estimating the upward and downward flux of soil water. This is more serious in annual crops where the maximum root depth at a given time is variable and often unknown, and where sub-surface cracks in the soil profile may be dominant in downward drainage.

The effect of moisture content on the sphere of influence of neutron moisture meters becomes critical near the soil surface and this can result in serious errors in water use measurements when most of the water

for evaporation comes from this layer. It is possible to apply a separate calibration for the first 30 cm (van Bavel and Stirk, 1967) where a relatively small area such as a lysimeter is concerned, but the applicability of this procedure in the open field is limited by the spatial variability of soil density and organic matter content in this layer. The use of standardised neutron reflectors for attachment to normal depth probes has been tried elsewhere without much success, partly because of the effect of such reflectors on the spherical symmetry of the slow neutron cloud around the detector. The alternative method of soil sampling in this layer was tried in the Muca experiments. The results and subsequent examination of the method suggest that it is unsatisfactory since many samples are required in order to achieve spatial average moisture content. This procedure could also change soil conditions near the neutron probe access sine to the extent that the entire soil profile water content may not be representative. Development of a simple but reliable and rapid method of soil moisture determination in the top layer of soil is therefore called for.

3. Estimating Potential Bry Matter Production

A high degree of correlation (98% of variance accounted for) has been obtained between measured dry matter production in all stages of growth and theoretical estimates of gross photosynthesis based on the method of Monteith (1965a). The only measurements required are daily totals of sclar radiation, leaf area index, and the relationship between light extinction in the canopy and leaf area index. The fraction of gross photosynthesis used in respiration under field conditions will be difficult to measure but results of Mwea experiments indicate that this fraction is relatively constant in a given climate and values of 0.5 for maine and 0.4 for beans would be applicable in a warm climate similar to that of Mwea.

These findings should be tested over a wide range of environments. The possibility of predicting maximum yields, not only of stover but also of grain, in crops of maise and beans will be a valuable aid not only in land use planning, but also to plant breeders who would welcome the possibility of assessing the potentialities

of genetic material before final yields are obtained. It is evident that a good variety of maise will be one which develops full leaf area quickly, mintains this condition for the larger fraction of the growing season, and in which all not photosynthesis after silking is stored in the grain.

4. Partition of energy in crops of maise and beans

Results from the Muea experiments have shown that if soil moisture is available, almost all net radiation in maise and bean crops may be converted into latent heat and that for short periods latent heat in the maise crop can exceed net radiation with consequential cooling of the air. The method of Fritschen (1965) for the determination of hourly values of Bowen's Ratio, β , has yielded consistent results. Analysis of chart records proved cumbersom and development of a simple data handling system is very desirable. The main obstacle will be the subjective quality control found necessary before integration of chart records. The fraction of total radiation stored in the form of total dry matter is very small compared with the latent and sensible heat components. This component does not appear to be correlated with the three soil moisture treatments in the 1968 maise crop and does not therefore seem to be a suitable parameter for defining water use efficiency of maise.

The drop in rate of dry matter formation at certain stages of growth in both maise and beans could, however, be a valuable indicator of the effects of physiological (in the case of maise) and /or excessive soil moisture (in the case of beans) conditions on the efficiency of solar energy utilisation for photosynthesis in field crops. These aspects are described in greater detail in Chapter IV, but further detailed experiments are required to check the reproducibility of such effects and, if confirmed, to work out the mechanism of these phenomena.

5. Water Use Efficiency

that water use efficiency of maise and beans cannot be defined in quantitative terms without reference to the environment. Crop management is a significant factor which cannot be quantified, but optimum plant populations and direct application of irrigation water to the soil beneath the vegetation should result in increased water use efficiency.

Provided soil moisture storage capacity is good, heavy doses of irrigation applied less frequently would be more economical than frequent applications of small amounts of water.

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Chemical Analysis of Soils from the Experimental Fields

APPENDIX 1

Field	Depth cm	pH (soil paste)	Organic	Exerengeacle Cations (m. eq. % CDS)				Conductivity EC ₁ x 10 ³ mhos at	Sodium adsorption	Exchangeable acdium	Yeonanical analysi (pipette method) (% ODS)		
			(% ODS)	7.	Ca	Ng	Mn	250 C (1:1 suspension)	ratio SAR	percentage ESP	Sand	Clay	Silt
	0-15	2.1	3.58	2.95	10.34	7.1	0.194	0.113			25.6	36.6	34.2
	15-30	6.2	3.58	2.67	10.33	7.0	0.167	0.072	Less than 1-2		29.0	36.6	30.8
A	30-45	5.9	3.ce	2.46	8.19	5.9	0.452	0.110		2033	25.6	42.0	29.3
	45-75	5.5	2.04	0.84	5.93	4.9	0.647	0.130		than	21.2	54.0	22.8
	75-105	5.6	1.36	0.36	4.50	4.2	0.139	0.682		1-2	19.7	57.8	21.1
	105-135	5.8	1.50	0.17	3.36	4.2	0.141	0.120			20.1	58.6	19.1
	135-150	5.9	1.26	0.17	2.58	4.2	0.085	0.140			20.4	55.9	22.4
	0-15	5.6	3.44	1.95	3.56	5.0	0.482	0.180			28.3	38.5	29.8
	15-30	5.6	2.46	1.29	7.50	4.0	0.450	0.300		8	26.0	44.0	27.5
	30-45	5.6	1.83	0.76	6.49	4.2	0.551	0.172		Less	23.3	49.7	25.1
В	45-75	5.6	1.60	0.45	5.08	4.1	0.227	0.144	Less than 1-2	than	24.4	54.0	20.0
	75-105	5.7	1.26	0.30	4.30	4.1	0.140	0.164		1-2	21.7	55.8	21.2
	105-135	5.7	1.10	0.17	3.67	4.8	0.108	0.170			20.0	57.2	21.6
	135-150	5.6	1.04	0.22	3.35	4.5	0.108	0.258			23.3	55.1	20.6

APPENDIX 2

Summary of Meteorological Observations at Mwea Irrigation Scheme

	Rainfall	y Konthly means			Air humidity		Wind speed	Redistion	Sumshine	Open water evaporation	
Period	Monthly totals			eans	Dewpoint mean of 0900 and 1500 hrs	Saturation deficit mean of 0900 and 1500 brs	Relative humility at 1500 nrs	Wean wind speed at 6 ft acove ground	Gunn Bellani	Year daily subshine hours	Raised pan with grid cover
	201.00	Sax.	Win CC	CC C	°c	mm of mercury	d	mph	Ly/day	hrs	EE
1967		11,									
Jul.	21.1	24.8	15.0	19.9	14.4	5.12	53	2.1	420	4.3	121.7
Aug.	20.6	25.5	14.2	19.9	13.3	5.97	48	2.6	462	5.0	141.5
Sep.	15.5	28.2	14.3	21.3	13.4	7.47	40	3.3	568	7.2	182.1
Oct.	236.5	28.0	15.8	21.9	15.6	6.42	48	3.4	556	6.7	183.1
Nov.	206.0	26.3	15.4	20.9	17.9	3.15	66	2.5	515	6.7	138.2
Dec.	С	27.2	12.5	19.9	16.2	3.61	56	3-1	530	10.1	178.8
1968											
Jan.	0	29.2	11.1	20.2	13.1	5.45	39	3.8	695	11.1	225.0
Pec.	149.6	28.6	14.7	21.7	15.3	6.43	43	3.5	574	7-1	175.3
Mari	123.2	27.6	15.4	21.5	17.2	4.52	59	3.1	544	6.7	160.8
Apr.	305.5	26.0	15.4	21.0	17.6	3.56	62	2.5	491	5.8	135.9
Nay	92.5	26.0	16.3	21.1	16.8	4.42	61	2.3	491	6.C	128.8
Jun.	17.0	25.2	15.0	20.1	15.1	4.77	58	1.8	435	4.6	112.0
Jul.	19.8	23.4	15.0	19.2	13.9	4-77	59	3.8	341	2.4	87.5
Aug.	e.9	23.5	14.7	19-1	13.4	5.06	55	1.9	353	2.4	102.4
Sep.	2.5	23.5	14.7	21.6	13.1	8.04	39	3.2	603	7.1	189.7
Cct.	168.4	28.6	17.1	22.8	15-3	7.78	45	3.2	562	6.4	179.6
Nov.	334.5	25.8	16.9	21.4	16.9	4.68	52	3.0	468	5.7	126.7
Dec.	136.1	25.7	14.0	20.4	16.3	4.05	58	2.8	591	9.4	153.9

appendix 3a

Lysimeter A and B readings and calculation of Et for the 1967 maize crop

					SEPT	EMBE	R 196	7						
Date			Lysime	ter A			Lysimeter B							
	Column height cm.	Column change x 1.03 mm.	Drainage	Rainfall	Et mm.	5-day total	Column height cm.	Column change x 0.08	Drainage	Rainfall	Et	5-day total		
1 2 3 4 5	34.1 33.6 33.7 33.2 33.2	5.2 -1.0 5.2 0 17.5	2.1 2.1 1.1 0.7 0.6	0 0 0 0 2.0	3.1 -3.1 4.1 -0.7 18.9	22.3								
6 7 8 9	31.5 33.8 31.3 30.9 30.4	-23.7 25.8 4.1 5.2 -4.1	0.3 0.2 0 0	8.1 5.1 0 0.3	-15.9 30.7 4.1 5.5 -4.1	20.3								
11 12 13 14 15	30.8 30.2 # 31.7 32.9 31.2	6.2 - 17.5 0	0 0 3.9 3.8 6.8	0 0 0 0	6.2 4.4* 4.4* 13.7 -6.8	21.9	79.1 78.0	9.7 2.6	1.2	0	8.5			
16 17 18 19 20	31.2 30.7 30.9 30.0 28.6	5.2 -2.1 9.3 14.4 8.2	4.0 2.8 1.7 1.4 0.6	0 0 0	1.2 -4.9 7.6 13.0 7.6	24.5	77.7 77.2 77.1 76.5 76.0	4.4 0.9 5.3 4.4 5.3	2.8 1.9 1.2 1.0	0 0 0	1.6 -1.0 4.1 3.4 4.7	12.8		
21 22 23 24 25	27.8 27.3 26.2 26.5 26.7	5.2 11.3 -3.1 -2.1	0.5 0.4 0.2 0.1	0 0 0 0	4.7 10.9 -3.3 -2.2 2.5*		75.4 75.0 74.6 74.3 74.2	3.5 3.5 2.6 0.9	0.4 0.2 0.1 0	0 0 0	3.1 3.3 2.5 0.9 2.5*			
26 27 28 29 30	* 26.5 28.2 27.2 25.6 24.7	10.3 16.5 9.3 4.1	3.2 1.8 2.9 1.2 0.9	0 0 0	8.4* 8.5 13.6 8.1 3.2	12.6	# 73.7 75.4 74.9 74.0 73.3	4.4 7.9 6.2 7.0	0.1	0 0 0	6.4* 4.4 7.9 6.2 7.0	12.3		

^{* 5-}day average

^{*} Days with irrigation

^{* 5-}day average

^{*} Days with irrigation

APPENDIX 3

Date			Lysine	ter A		Lysimeter B						
	Column height cm.	Column change x 1.03 mm.	Drainage	Rainfall	Et ==+	5-day total	Column height em.	Column change x 0.89	Drainage	Rainfall	Et mm.	5-daj total
1	27.5	22.7	10.4	1.8	14.1		74.3	17.6	6.4	1.8	13.0	
2	25.3	-38.1	2.5	41.7	1.1		72.3	-29.0	6.3	41.7	6.4	
3	29.0	10.3	13.6	4.8	1.5		75.6	11.4	6.3	4.8	9.9	
4	28.0	13.4	11.3	4.1	6.2		74.3	11.4	6.3	4.1	9.2	
5	26.7	4.1	5.8	5.3	3.6		73-0	5.3	6.3	5.3	4.3	
-						26.5						42.8
6	26.3	-15.5	3.7	24.1	4.9		72.4	-16.7	5.1	. 24.1	2.3	
7	27.8	7.2	6.3	2.0	2.9		74.3	7.9	6.1	2.0	3.8	
8	27.1	5.2	4.7	4.3	4.8		73.4	7.9	5.6	4.3	6.6	
9	26.6	9.3	3.3	0 1	6.0		72.5	11.4	3.5	0	7.9	
10	25.7	4.1	1.2	0	2.9		71.2	5.3	1.9	0	3.4	
10	23.1	* · I	1.5	1 1	2.3	21.5	12.6	7.7	1.3		2.4	24.1
11	25.3	9.3	0.4		8.9	21.7	70.6	7.9	0.7	0	7.2	24-1
12	24.4	9.2	0.4	6.6	6.6			1.8		6.6		
13		0		1.8			69.7		0.1		8.3	
	24-4		0		1.8		69.5	4.4	0	1.8	6.2	
14	24.4	1.0	0	0	1.0		69.0	1.6	0	0	1.8	
15	24.3	-6.2	0	16.5	10.3		68.8	-0.9	0	16.5	15.6	
1						28.6						39-3
16	24.9	-6.2	0	11.9	5-7		68.9	0.9	0	11.9	12.8	
17	25.5	-2.1	0	3.8	1.7		68.8	0	0	3.8	3.8	
18	25.7	-3.1	0	1.3	-1.8		68.8	-3.5	0	1.3	-2.2	
19	26.0	4.1	0	0	4.1		69.2	3.5	0	C	3.5	
20	25.6	4.1	0	0	4.1		68.8	3.5	0	0	3-5	
						13.8						21.4
21	25.2	-12.4	0	12.7	0-3		68.4	-12.3	0	12.7	0.4	
22	26-4	1.1	0	0.8	1.9		‡ 69.8	2.0	0	0.8	2.8	
23	26.3	5.4	0	0	5.4		69.6	8.8	0	0	8.8	
24	25.8	-1.1	0	5.6	4.5		68.7	1.0	0	5.6	6.6	
25	25.9	-35.3	0	39.4	4.1		68.6	211.7	0	39.4	4.7*	
				1 1		16.2						23.3
26	29.2	-10.7	0	9.7	-1.0		47.0	-36.3	0	9.7	-	
27	30.2	11.8	7.2	7.9	12.5		50.4	-5.9	0	7.9	_	
28	29.1	12.8	6.4	0	6.4		51.0	-11.8	0.7	0	_	
29	27.9	10.7	4.5	0 !	6.2		52.2	-2.9	3.5	a	_	
30	26.9	4.3	2.0	0	2.3		52.5	-9.8	2.3	0	_	
		1.7				26.4	//	2.0	/	~	_	

^{* 5-}day average

^{*} Calibration factor changed to 1.07

⁺ Calibration factor changed to 0.98

APPENDIX 3 ...

Lysimeter A and B readings and calculation of Et for the 1967 maize crop

			D	ECE	MBER	1 9 6	7					
		L	ysimeter A			Lysimeter B						
Date	Column height	Column change x 1.03 mm.	Drainage	Et mm.	5-day total	Column height cm.	Column change x 0.98	Drainage	Et mm.	5-day total		
1	26.5	1.1	1.1	0		53.5	4.9	1.3	3.6			
2	25.4	8.6	0	8.6		53.0	8.8	0	8.8			
3	24.6	. 8.6	0	8.6		52.1	2.9	0	2.9	1		
4	23.8	3.2	0	3.2		51.8	10.8	0	10.8	Ì		
5	23.5	7.5	0	7.5		50.7	14.7	0	14.7			
.				-	27.9					40.8		
6	22.8	-4.3	0	-4.3		49.2	-4.9	0	-4.9			
7	23.2	-7.5	0	-7.5		49.7	-4.9	0	-4.9			
8	23.9	-6.4	0	-6.4		50.2	-1.0	0	-1.0			
9	24.5	8.6	0	8.6		50.3	14.7	0	14.7			
10	23.7	-1.1	0	-1.1		48.8	-2.0	0	-2.0			
- 1					-10.7					1.9		
11	23.8	-1.1	0	-1.1		49.0	2.0	0	2.0			
12	23.9	3.2	0	3.2		48.8	3.9	0	3.9			
13	23.6	13.9	0	13.9		48.4	13.7	0	13.7	1		
14	22.3	-3.2	0	-3.2		47.0	-4.9	0	-4.9			
15	22.6	5.4	0	5.4		47.5	4.9	0	4.9			
					18.2					19.6		
16	22.1	10.7	0	10.7		47.0	3.9	0	3.9			
17	21.1	4.3	0	4.3		46.6	7.8	0	7.8			
18	20.7	-4.3	0	-4.3		45.8	-2.9	0	-2.9			
19	21.1	-1.1	0	-1.1		46.1	-1.0	0	-1.0			
20	21.2	4.3	0	4.3		46.2	3.9	0	3.9			
					13.9					11.7		
21	20.8	19.3	0	19.3		45.8	15.7	0	15.7	1		
22	19.0	5.4	0	5.4		44.2	3.9	0	3.9			
23	18.5	7.5	0	7.5		43.8	5.9	0	5.9			
24	17.8	-1.1	0	-1.1		43.2	0	0	0			
25	17.9	5.4	0	5.4		43.2	4.9	0	4.9			
					36.5					30.4		
26	17.4	5.4	0	5 - 4		42.7	5.9	0	5.9			
27	16.9	13.9	0	13.9		42.1	2.0	0	2.0			
28	15.6	6.4	0	6.4		40.9	6.9	0	6.9			
29	15.0	-4.3	0	-4.3		40.2	-1.0	0	-1.0			
30	15.4	8.6	0	8.6		40.3	5.9	0	5.9	1		
31	14.6	-5.4	0	-5.4		39.7	2.9	0	2.9			
					24.6					22.6		

There was no rainfall in December.

APPREDIX 30

Orop water use Bt. from lysimster compared with ben evaporation. Sp. and Penman estimate Bo 1967 mains crop

Per	ied	3		Bo	Ep
		Lys. A	Lys. B	ma	100
Sep.	1-5	22.3		26.4	29.0
	6-10	20.5	-	26.3	19.6
	11-15	21.9	•	27.1	24.4
	16-20	24.5	12.8	33.8	36.0
	21-25	12.6	12.3	34.3	31.5
	26-30	41.8	31.9	34.3	35.5
Oct.	1-5	7.0	16.7	35.8	38.0
	6-10	37.0	34.1	35.9	32.5
	11-15	45.8	40.3	33.3	32.0
	16-20	46.6	61.8	33.4	30.0
	21-25	30.9	26.0	25.3	19.8
	26-51	46.8	58.6	30.4	30.7
Hov.	1-5	26.5	42.8	27.0	24.6
	6-10	21.5	24.0	27.0	25.5
	11-15	28.6	39.1	29.2	26.0
	16-20	13.7	21.4	29.1	25.5
	21-25	16.2	23.5	22.9	20.9
	26-50	26.4	•	21.9	17.0
Dec.	1-5	27.9	40.8	28.4	27.0
	6-10	-10.7	1.9	28.5	27.0
	11-15	18.2	19.6	30.3	24.0
	16-20	13.9	11.7	50.4	29.0
	21-25	36.5	30.4	50.4	30.5
	26-51	24.6	22.6	36.5	36.1

APPENDIX 4.

		K	ARCH					AFRIL						KAY			
Date	Column height cr.	Column change x 0.88	Rainfall	17	5-day total	Column neight cm.	Column change x C.86	Drainage	Rainfall	Et .	5-day total mr.	column telent cm.	Column change x 0.78	Drainige	Rainfall	î. î.	5-day total
1	76.3	0	4.3	4.3		74.9	-11.4	0	19.0	7.6		E7.0	16.4	12.6	6.9	10.5	
2	78.3	1.8	4.1	5.9		76.2	-10.6	0	18.8	8.2		84.9	-9.4	10.6	27.2	7.0	1
3	78.1	4.4	0	4.4		77.4	2.6	0	7.1	9.7		bb.1	17.9	12.6	0	5.1	
4	77.6	-6.2	9.4	3.2		77.1	2.6	0	1.8	4.4		63.8	-7.0	6.4	20.5	5.1	
5	76.3	-1.8	7.6	5.8		76.8	5.3	0	0	5.3		84.7	11.7	5-4	3-1	0.4	
,	10.5	-1.0	7.0	7.0	23.6	10.0	1 2.2			41.4	35.2	04.1	11.	0.4	7.2	0.7	34.2
ó	76.5	-4.4	9.4	5.0	27.0	76.2	7.0	0	0.5	7.5	22.2	83.2	10.9	6.4	0	2.5	77.0
7	79.0	-0.9	1.0	0.1		75.4	-5.9	0	6.4	7.5		61.8	6.2	4.0	3-1	5-3	
8	79.1	2.6	3.3	5.9		75.5	5.3	0	0	5+3		81.0	5.5	2.4	C	3.1	
	75.A	5.3	1.0	6.3		74.9	8.8	0	0	8.8		60.3	6.2	1.5	0.6	5.5	
10	76.2	2.6	0	2.6			1.8	0	4.1	5.9		79.5	6.2	0.9	0.5	5.8	1
10	76.2			2.0		73.9	1.8	0	4+1	7.9	35.0	19-5	0.2	0.9	0.5	7.0	24.2
11	77.9	3.5	1.3	6.8	19.9	73.7			7.6	7.6	33.0	76-7	1.0	0.6	4.3	5.3	2400
							0	0				78.5	3.9	0.6	0	3.3	
12	77.5	-3.5	6.9	3-4		73-7	0		4.3	4.3							1
13	77.9	4-4	0	4.4		73-7	10.6	0	2.5	13.1		76.0	3.9	0.7	0	3.2	
14	77 - 4	3.5	0	3.5		72.5	0.9	C	1.0	1.9		77-5	4-7	0.4	0	4.3	1
15	77.0	2.6	С	2.6		72.4	-17.6	0	20.8	3.2		70.9	2.3	0.4	0.5	2.4	
					18.7						30.1						10.5
16	76.7	5.3	0	5-3		74.4	-16.7	0	18.0	1.3		70.6	4.7	0.3	. 0	4.4	1
17	76.1	0.9	0	0.9		76.3	7.0	0	0	7.0		76.0	3+2	0.2	0.6	3-7	1
16	76.0	-21.2	30.2	9.0		75.5	4.4	0	1.6	6.2		75.6	2.3	0.2	0.5	2.9	
19	78.4	-15.8	21.6	6.0		75-0	7.9	0	0	7.9		75.3	3.1	0.1	C	3-0	-
20	80.2	3.5	1.0	4.5		74.1	2.6	0	5-3	7.9		74.9	1.6	0.1	0.8	2.3	١
					25.7						30.3						10.3
21	79.6	-11.4	13.7	2.3		73.8	4.4	0	0.5	4.9		74.7	3-1	0.1	0.3	3-3	
22	81.1	2.6	0	2.6	i	73.3	-51.9	0	57.0	5.1		74.3	3.9	0	0.3	4.2	
23	80.6	7.0	C	7-0		"9.2	-19.4	-	23.8	4.4		73.8	-0.8	0	0	-0.8	
24	80.0	4.4	0.5	4.9	1	61.4	7.0	С	0.5	7-5		73.9	0.8	0	2.3	3.1	1
25	79.5	-5-3	7.6	2.3		80.6	0	0	4.3	4.3		73.6	-12.5	٥	16.0	4.3	
					19.1						20.2						14.1
26	80.1	5.3	0	5-3		80.6	-7.8	0	11.7	3.9		75 - 4	3-1	C	0-3	3.4	
27	79.5	6.2	0	6.2		81.6	4.7	0	0	4.7		75.0	3.9	0	0	3.9	
25	76.6	10.6	0	10.0		81.0	-27.3	0	40.5	13.2		74-5	-0.8	0	3.3	2.5	-
29	77.6	7.9	0	7.9		84.5	3.9	0	1.1	5.0		74-6	2.3	0	0.3	2.0	4
30	76.7	9.7	0	9.7		84.0	-23.4	7-0	44.7	14.3		74.3	2.3	0	0	2.3	
31	75.6	6.2	0	6-2								74.0	3.1	0	0	3.1	
					45.9					İ	41.1						17.8

N.B. No drainage was recorded in March.

APPENDIX 4b

Vater use of beans (Canadian Wonder), Nece Experiment 1968

Month	5-day periods	Lys.A Bt	Lys.B	E. Pan	Po mm	Ft/Ro Lys. A	Et/Eo Lys.A 10 day averages	Rainfall	L.A.I.
March	I	23.6	26.0	25.4	25.7	0.92		25.4	0.0
	II	19.9	18.6	22.8	25.8	0.77	0.85	14.7	0.02
	III	18.7	8.3	32.0	31.8	0.59	0.68	8.2	0.10
	IV	25.7	LEAK	29.5	29.0	0.89	0.74	42.0	0.36
	▼	19.1	BECA-	15.8	23.5	0.81	0.85	21.8	0.80
	VI	45.9	VATED	35-3	37.0	1.24	1.02	0.0	1.75
April	I	35.2		24.9	27.8	1.27	1.25	46.7	3.00
	II	35.0		21.6	25.4	1.38	1.33	13.0	3.70
	III	30.1		22.6	24.8	1.21	1.30	36.2	4.1
	IA	50.3		19.1	21.9	1.43	1.32	25.1	4.25
	▼	26.2		18.5	22.2	1.18	1.31	105.1	4.0
	AI	41.1		17.3	24.8	1.66	1.42	106.9	3.1
May	I	34.1		21.6	30.3	1.13	1.40	57.7	2.0
	II	24.2		23.6	27.8	0.87	1.00	4.4	1.0
	III	18.5		22.1	27.4	0.68	0.78	4.8	0.0
	IA	26.3		22.1	25.6	0.64	0.66	2.4	0.0
	V	14.1	13	18.0	22.6	0.63	0.63	19.7	0.0
	AI	17.8		20.5	28.2	0.63	0.63	3.9	0.0

Germination in both lysimeters on 6th March. Harvest in 'A' 23rd May. Leaves dried off from 13th May.

APPENDIX S

	Social Both	u			7+6	15.1	2	5
ĺ	i i	40044	2.4 3.1 1.4	46940	94004	C 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	W 0 0 0 W	0.0
	Rain ell	00000	20400	00000	8 8 8	00000	00000	o
AUGUST	9 21 84 C	00000	00000	00000	00000	00000	00000	0
	Col.ma cnan z O s mm.	0,0044	04411	22.28	W 000 H	20067	11000	3.0
	Column height	5.5.5.5 5.5.5.5 5.5.5.5 5.5.5.5 5.5 5	7722	24EEE	722.7	71170	770.7	73.59
	5-d to: Em.	,	4	4	00	4 - 12	n n	25.4
	Bt	200041	1.2	5 K O O H	1 1 1 0 0 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	1.00FM	500000 50000	7:-7
	Reisfall	00000	110000 ww	0,000	00000	00040	40000 800	0
JULY	Draina e	00000	00000	00000	0 0 0 0 0 0 0 0 0 0	W41-W0	00000	0.4
	Col.mn change x 0.89	0000d	00000	1-00 1-00	1 1 2 2 4 5 4 5 4 5 4 5 4 5 4 5 4 5 4 5 4 5	00040	24000	7. B
	Column height cm.	772.25	72.1	411111	1888 1900 1900 1900 1900 1900 1900 1900	77.7	44.04.0 6	73.5
	D B t t				_			

ž Days with irrigation

APPENDIX 5 (continued)

Lysimeter A regions and calculation of Et for the 1968 maize crop

			SEPTEMBER					EZHOTSO					NCVE	GER		
Date	Column height cm.	Column change x 0.89 mm.	Rainfall	Et	5-day total	Column height cm.	Column cmarge x C.89	Rainfall	Ξt	5-day total	Column height cm.	Column change x 0.89	Drainage	Reinfall	Ēt	5-day total
2	71.5	4.5	0	4.5		62.1	6.2	2.2	8.4		50.7	-17.9	0	20.0	2.1	
2	71.0	1.8	0	1.8		61.4	5.3	1.2	6.5		58.7	1.9	0	7.9	9.7	
3	70.8	3.6	0	3.6		50.8	6.2	0	6.2		58.5	6.2	0	1.5	7-7	
4	70.4	6.2	0	6.2		60.1	6.2	0	6.2		57.8	5.3	0	1.0	6.3	
5	69.7	4.5	0	4.5		59.4	3.6	0	3.6		57.2	-4.5	0	9.6	5.1	
1					20.6					30.9						30.9
6	69.2	6.2	0	6.2		59.C	7.1	0	7.1	, , ,	57.7	1.8	0	0	1.8	20.2
7 13	68.5	-	0	5.3*		56.2	6.2	0	6.2		57.5	3.6	0	1.3	4.9	
8 !	71.6	-	0	5.30		57.5	7.1	0	7.1]	57.1	-2.7	0	6.9	4.2	
9	71.1	3.6	0	3.6		56.7	7-1	0	7.1		57.4	-3.6	0	3.3	-0.3	
10	70.7	6.2	0	6.2		55.9	4.5	0	4.5	1	57.8	0.9	0	2.3	11.2	
					26.6					32.0						21.6
11	70.0	7.1	0	7.1		55.4	6.0	1.2	9.2		56.8	3.6	0	0	3.6	
12	69.2	1.8	0	1.8		54.5	7.1	0	7-1		50.4	-0.9	0	13.5	12.0	
13	69.0	5-3	0	5.3		53.7	5.2	0	6.2		50.5	-7.1	0	5.3	-1.8	
14	68.4	8.0	0	8.0		53.0	-18.7	29.5	10.8		57-3	5.3	. 0	0	5.3	
15	67.5	4-5	٥	4.5		55.1	9-9	0.8	9.7		50.7	6.2	0	0.5	0.7	
- 1					26.7					43.0						26.4
lo l	67.0	3.0	0	3.6		54-1	-17.8	28.0	10.2		50.0	4.5	0	0.8	5.3	
17	66.6	6.2	0	6.2		56.1	0	7-4	7.1		55.5	-10.7	0	16.7	6.0	
18	65.9	3.6	0	3.6		56.1	-5-3	7.1	1.8		56.7	-4.5	0	14.2	9.7	
19	65.5	-	0	4.5*		56.7	3.6	0	3-6		57.2	-24.9	0	27.8	2.9	
20 1	65.0	-	0	4.5*		56.3	7.1	0	7.1		60.0	-44.5	0	56.0	11.5	
					22.4					29.8						35.4
21	70.9	8.0	C	8.0		55.5	9.8	0	9.8		65.0	-41.0	0	47.5	6.5	
22	70.0	10.7	0.3	11.0		54.4	6.2	0	6.2		69.6	-2.7	6.9	18.3	8.7	
13	100.5	7.1	0	7.1		53.7	1.3	0	1.8		69.9	-0.2	14-1	19.6	-0.7	
24	68.0	7.1	0	7.1		53.5	8.9	0.8	9.7		70-6	-3.6	15.6	27.0	7-6	
25	67.2	8.0	0	8.0		52.5	7.1	1.2	0.3		71.0	-7-1	19-0	17.0	-9.1	
1					41.2					35.8						13-0
6	60.3	4.5	0	4.5		51.7	-71.2	80.0	0.8		70.2	16.9	18.2	4.3	3.0	
27	65.8	a.c	0	8.0		59-7	9.8	0	9.6		68.3	10.7	10.9	3.1	2.9	
28	04.9	3.6	0	3.5		56.6	10.7	0	10.7		67.1	8.9	6.3	1.5	4.1	
29	64.5	13.4	0	13-4		57.4	8.0	0.5	8.5		66.1	7.1	3.9	2.3	5.5	
50	63.0	8.0	0	8.0		56.5	-3-6	9-1	5.5		65.3	1.3	2.5	5.6	4.9	
11]						50.9	1.8	0	1.a							
j					37.5					45.1						20.4

^{*57}day average

Bo drainage recorded in September and October.

[&]amp; Days with irrigation

Water use (Et) for the 1968 Maise Grop (Wet Treatment)

Month	5-day periods	Et(mm) Lys. A	Penman Bo (mm)	Bt/Bo	Rainfall (mm)	.Irrig- ation (ma)	Pan Bo (mm)
Jul.	A AI IX III III III	8.1 5.4 8.8 - 19.5 15.4	26.2 16.7 19.4 22.7 17.4 19.9	0.51 0.52 0.45 - 1.12 0.77	3.6 1.8 2.5 0.8 1.5 6.9	152	20.3 11.9 15.2 18.0 11.2
Aug.	A A A A A A A A A A A A A A A A A A A	5.6 7.0 9.1 12.1 14.0 19.3	22.0 13.0 22.1 15.2 22.9 33.8	0.25 0.54 0.41 0.30 0.61 0.57	0.3 4.6 0.0 3.5 0.5 0.0	25	20.1 8.6 18.3 10.2 16.8 28.4
Mep.	AI A IA III II	20.6 26.6 26.7 22.4 41.2 37.5	36.7 27.4 33.9 32.8 35.6 39.5	0.56 0.97 0.79 0.68 1.16 0.95	0.0 0.0 0.0 0.0 0.3	42 71	34.5 26.0 30.5 31.0 31.0 34.0
Oct.	AI A A II II II	50.9 52.0 43.0 29.8 55.8 45.1	37.5 36.6 34.3 27.4 31.7 37.0	0.82 0.87 1.25 1.09 1.13 1.22	3.4 0.0 31.5 42.2 2.0 89.6		50.0 51.7 52.8 24.9 27.8 50.6
Bov.	AI A A IA III II	30.9 21.8 26.4 35.4 13.0 20.4	30.2 24.7 29.5 26.5 19.1 24.5	1.02 0.88 0.89 1.34 0.68 0.83	40.0 13.8 19.3 115.5 129.4 8.9		22.4 19.5 24.9 24.9 15.5 11.4

APPENDIX 6

Comparison betwee: Dt estimated from soil sampling and from lysimaters A.& B

1967 maize crop

					···		Perio	d 1967						
Depth in	12/9	- 16/9	26/9	- 10/10	10/10	- 24/10	24/10	- 8/11	8/11 -	- 22/11	22/11	- 5,12	5/12 -	21/12
ft.	A ins	B ins	A ins	B ins	A ins	B ins	A iis	B ins	A ins	B ins	A ins	.13	A ins	B
1 1 2	01 08	.31	.C4 .11 .31	.10 .21 .22	53 24 05	84 28 25	.06 13 25	27	12 .14 .52	.06	.50 .08 10	.16	.33 .56	.63 .59
3 4 5 6	21 .06 13	10 .00 .15	.00 21 06 14	.30 .10 .14	.01 .06 .03	26 15 14 15	49 21 43	37	.50 .49 .26	.30	27 54 .02 20	.43 .36 01 39	.51 .43 .63	.25 .13 .25
Total water	30	-73	.05	1.04	-1.11	-2.08	-1.	7.10	2.34	1.99	51	.74	3.31	2.51
Total water removed (mm)	-7.6	18.5	1.3	26.4	-28.2	-52.8	- :.5	-58.4	59.4	50.5	-13.0	18.8	84.1	63.8
Irrigation (mm) Rainfall(mm)	68.8	43.8	57.6 19.4	27.5	79.9 83.1	51.7 83.1	0 218.0	0 218.0	0 58.9	ò 58.9	0	o 63.4	0	0
Et (mm)	61.2	62.3	78.3	73.3	134.8	82.0	189.8	159.6	118.3	109.4	50.4	82.2	84.1	63.8
Et (lysimeter) (mm)	57.1	43.6	79.0	71.3	109.5	117.7	92.6	125.7	53.4	75.0	68.3	85.9*	40.7	48.9

Estimate
* 60.9 + 5 days 0 5 mm/day

APPENDIX 7a

Estimates of soil moisture with FAL 104 moisture meter down erop 1968

23.2.68

					100000	-				
				Wast	er cont	ent (.u	1.)			
Depth cm.			Field A				Fi	eld 1		
	ΛA	A.B	ΛC	(LA	Lym.A	103	BF	nG	rili	Lya.
30 60 90 120 150	48.0 66.0 90.0 93.0 106.5	51.0 78.0 93.0 96.0 106.5	60.0 90.0 93.0 93.0 93.0	48.0 78.0 95.0 96.0 90.0	55.5 106.5 99.0 111.0 136.5	55.5 78.0 96.0 99.0 100.5 121.5	60.0 75.0 96.0 106.5 100.5	66.0 79.5 99.0 102.0 102.0	66.0 81.0 99.0 106.5 111.0 117.0	66.0 75.3 73.0 81.0 93.0
Total Lean	510.0	531.0	535.5	495.∪	508.5	550.5	552.5 5	565.5	530.5	393.0
					0.3.6	3			1	
30 60 90 120 150 180	114.0 114.0 103.5 96.0 103.5 103.5	118.5 126.0 106.5 93.0 103.5 100.0	118.5 129.0 111.0 100.5 90.0 111.0	103.5 118.5 106.5 99.0 90.0 85.5	96.0 130.5 118.5 126.0 159.0	1: 1.5 1: 1.5 99.0 99.0 106.5 121.5	126.0 126.0 111.0 103.5 103.5	129.0 129.0 121.5 103.5 103.5 121.5	129.0 126.0 114.0 106.5 111.0 118.5	129.0 103.0 96.0 85.1 88.1
lotal	634.5	647.5	660.0	603.0	(.39.0	654.0		70u.0	705.0	502.
					17.4.6	3				
30 60 90 120 150 180	139 98 102 100 111 112	137 114 105 104 113 107	138 110 109 104 99 113	121 102 113 107 99 97	115 113 107 130 101	128 121 117 113 120 133	152 115 116 117 117 127	154 133 125 119 116 133	154 121 121 122 122 130	
Fotal Mean	662	680	673	639	625	732	744	782 57	770	*****

				Water	content	(mm.)			
Depth cm.			Field /				Fie	ld B	
	ΛA	Λß	AC	AD	Lys.A	BE	BF	BG	ВН
		1	1		2.5.68				
30 60 90 120 150 180	130 127 127 119 128 127	130 140 138 137 142 125	136 139 140 125 119 134	121 133 140 126 115 113	110 153 128 146 178	114 141 143 131 140 152	136 135 143 137 134 139	146 146 145 139 136 155	143 143 142 141 143 147
Total Mean	758	812	793 78	748	715	821	824 84	867	859
			İ		8.5.68				
30 60 90 120 150 180	112 114 124 119 129 136	117 134 130 129 114 131	119 134 132 122 120 139	107 121 132 126 113 119	99 145 121 145 181	107 129 131 131 137 152	125 131 135 133 135 141	129 138 139 134 133 152	132 132 131 137 142 146
Total Mean	734	755 7	766 43	718	691	787	800 80	825 8	820
				2	29.5.68				
30 60 90 120 150 180	84 100 110 110 125 126	93 104 124 122 127 125	101 119 120 113 116 136	87 103 119 117 105 112	83 130 119 144 179	80 113 121 120 129 148	104 114 122 123 125 134	109 117 125 123 123 144	109 117 122 124 131 141
Total Nean	655	695 6	705 76	648	655	711	722 73	741 0	744

78.

APPENDIX 76
Estimates of soil molecure with KTV 90 moisture meter
36an crop 1968

					Water cor	itent (u.,)			*
Depth cm.			Fiel	d A	*******			Fiel	d B	
	ΛΛ	AB	AU	Ali	Lys. A	DE	UP	30	Bli	Lys. B
	1				1.3.63					
30 60 90 120 150 180	45 53 45 42 53 71	108 111 90 105 105 103	125 119 105 105 99 108	45 45 30 35 38 35	32 48 45 43 53	113 96 96 96 105 110	128 117 96 99 102 105	113 114 99 105 105 113	125 117 99 105 108 113	119 122 117 90 108
Total Mean		633 7 exc	661 ludin	228 E hk		616	647	649 64	667 4	556
					9.3.68					
30 60 90 120 150 180	93 116 90 99 102 102	93 99 122 102 90 93	107 105 105 102 96 87	96 107 102 87 90	90 111 107 111 111	99 102 90 99 99 105	107 102 111 99 87 102	113 113 107 105 107 107	116 113 102 99 105 113	107 116 96 82 105
Total Lean	602		602 94	572	530	594	608		648 26	506
-				j	13.3.68		****			
30 60 90 120 150 180	105 111 90 96 93 99	107 102 93 111 111 102	107 111 99 99 105 105	107 107 99 96 90 102	85 113 105 119 153	77 93 87 85 82 93	96 93 85 82 79 96	93 105 96 99 96	96 87 93 90 102 96	102 107 96 82 82
Total Bean	603		626 14	601	575	517	531		564	469

APPENDIX 7b (continued)

					Water cor	itent (ma.)			
Depth on.			Fiel	a A	hedrodreds die direktralende			Fiel	d B	
	AA	AB	AC	Aυ	Lys. A	BE	BF	BG	ВН	Lya. B
			5		16.3.68	1	1			
30 60 90 120 150 180	90 90 96 96 96 99	96 107 105 105 96 104	105 111 102 102 105 102	87 102 105 96 96 111	73 102 116 122 145	93 105 96 99 93 102	96 122 105 93 99	107 107 105 105 102 102	105 105 99 99 113 113	102 113 105 87 102
Total Nean	567	613	627	597	558	588	614 6	628 16	634	509
: (20.3.68	1		1	1	
30 60 90 120 150 180	90 90 90 96 93 96	102 96 87 93 90 96	113 107 105 90 93 96	96 99 102 96 96	87 107 102 119 136	96 102 102 96 99 102	102 99 99 96 99 102	116 105 105 107 111 107	113 99 96 107 105 107	102 99 99 96 99
Total Mean	555		604 77	585	551	597	597 6	651 18	627	495
			1		27.3.68	1				
30 60 90 120 150 180	85 90 87 85 85 85	96 102 102 93 105 92	102 105 93 99 99 102	87 99 96 93 93	79 107 99 107 148	87 107 111 99 105 113	102 105 105 102 99	105 105 105 111 130 116	113 105 96 96 111 105	
Total liean	517	590 5	600 66	558	540	622	612	672 33	626	

APPENDIX 7c

SULLIARY OF SOIL HOISTURE CHANGES (mm) AT DIFFERENCE DEPTHS IN THE SOIL PROFILE, IMASUREMENTS WITH E.A.L. HEUTRON HOISTURE HETER

Negative figures indicate increase in soil moisture.

23.7.68 - 31.7.68

Soil		File	ld A			id a	ld B			Fiel	1 0	
depth										-		
(cm.)	A.A	AB	ΛC	AD	1713	BF	BC	BH	CI	ÇJ	CK	CT
5 15 30 50 70 90 110 130 150	0.6 -0.3 3.6 3.0 2.0 0.8 2.6 1.6	-1.0 1.6 4.8 8.6 2.6 2.4 2.8 0.0 1.6	3.4	26.8 -3.8	-0.3 -1.2 2.4 5.0 2.0 2.6 5.2 2.6 2.4	3.6 3.4 2.0 3.0 4.2 0.6	-1.2 -0.6 2.0 3.0 3.4 3.8 1.0 4.4 2.6	3.8	4.6 1.1 4.0 13.8 6.0 4.4 3.4 3.2 2.4 0.0	0.6 6.6 1.8 3.4 3.0 3.6 -15.0 2.0 0.0	3.7 2.6 8.2 4.6 4.0 3.2 4.0 5.8 0.0 -1.6	4.2 -0.7 3.6 0.0 5.0 15.2 4.0 1.2 -2.2
Total	15.3	18.0	14.7	42.3	20.3	23.8	21.8	دورلا	42.9	7.8	34.5	32.3
Rain Irrig ation	8.4	8.4	8.4	8,4	8.4	8.4	8.4	8.4	8.4	8.4	8.4	8.4
Et	23.7	26.4	23.1	50.7	28.7	32.2	30.2	25.3	51.3	16.2	42.9	40.7
Hean Et				31.0				29.1				37.8
					31.7	68 -	7.8.0	.8				
5 15 30 50 70 90 110 130 150	1.2 -0.3 2.0 2.0 1.4 3.2 3.0 2.4	1.3 2.5 2.4 -1.0 12.4 1.4 0.6 -0.6 2.4	5.6 -4.1 2.4 2.0 2.4 3.0 4.0 0.8 1.6		1.2	3.3 -3.0 4.0 1.0 3.0 3.6 2.2 1.8 0.4 2.0	3.5 3.0 2.4 0.4 2.4 2.6 3.0 4.4 2.6	3.5 0.5 2.6 4.0 3.0 2.6 2.0 3.2 1.4	4.2 5.0 -0.6 -4.8 4.4 2.4 3.4 1.2 1.2	8.5 6.9 2.8 2.0 2.0 3.0 2.6 0.8 3.2 2.2	2.6 4.7 3.4 4.0 3.2 2.4 1.6 2.2	
Total	19.7	23.4	21.7	35.8	23.6	18.3	25.6	25.0	20.8	34.0	27.7	22.8
Rain .	2.8	2.8	2.8	2.8	2.8	2.8	2.8	2.8	2.8	2.8	2.8	2.8
Irrig ation	-	~	-		-	-	-	-	-	-	-	~
Et	22.5	26.2	24.5	38.6	26.4	21.1	28.4	27.8	23.6	36.8	30.5	25.6
liean Et				28.0				25.9				29.1

7.8.68 - 21.8.68

Soil Depth		Fie	ld A			Fiel	Ld B			Fiel	.d C	
(cm.)	ΛΛ	AB	AC	AD	BE	HF	BG	18H	CI	CJ	CK	CL
5 15 30 50 70 90 110 130 150	-3.3 -2.0 0.4 0.6 -0.8 -1.2 -0.2 -1.6 -3.2 -0.8	-3.0 -5.8 -0.8 0.2 -11.6 -0.4 -0.6 2.2 -0.4	-4.6 -0.8 -1.4 -1.0 0.2 -0.8 -1.8 1.4 -0.6 -0.8	2.8 -0.8 -3.8 -3.2 -0.2 -0.8 0.0	0.6	2.6 0.8 -0.2 1.2 0.6 0.8 1.4 -0.2 3.0	4.5 1.2 2.2 1.4 1.0 0.6 0.6 0.2 -1.4 0.2	-7.3 1.1 1.8 -0.8 -0.2 0.6 1.0 -0.8 2.0 2.2	-3.6 -1.1 5.4 -1.2 -1.4 0.4 0.0 3.2 1.8	-5.2 0.2 1.4	-3.6 -3.4 0.4 1.4 -1.8 0.2 2.6 0.4 -0.6 -0.2	-2.1 -4.6 1.2 -0.2 1.8 0.4 3.0 1.6 -0.2
Total	-12,1	-20.6	-10.2	-4.1	-5.4	10.4	10.5	-0.4	4.3	-0.7	-4.6	4.5
Rain	5.3	5.3	5.3	5.3	1.8	1.8	1.8	1.8	5.3	5.3	5.3	5.3
Irrig ation	25.0	25.0	25.0	25.0	12.5	12.5	12.5	12.5	18.8	18.8	18.8	18.8
Et	18.2	9.7	20.1	26.2	8.9	24.7	24.8	13.9	28.4	23.4	19.5	28.6
Hean Et				18.6				18.1				25.0
				21.8.0	68 - 1	8.8.	ių.					
5 15 30 50 70 90 110 130 150	1.4 -1.8 -3.2 -1.2 2.0 2.0 2.0 3.8 3.8	-4.5 -4.2 -1.4 2.4 2.2 0.0 3.8 0.8 5.1	-3.2 -8.6 -8.2 -8.0 -9.0 -7.4 -6.6 -1.8	1.4	3.5	-0.7 24.2 4.8 2.0 0.8 1.2 0.2 0.6	1.2 0.4 -0.8 0.0	0.5 1.8 0.8 0.2 0.4 0.2 -0.2 1.0 0.2	-5.1 1.8 2.4 0.6 0.2 0.4 -0.4	-6.9 -7.2 0.0 0.6 -0.6 2.0 -1.0 0.6 0.0	-2.0 0.0 0.4 0.4 1.2 0.6	3.3 -2.2 1.2
Total	10.2	. 5.4	-66.1	-5.2	70.0	28.4	13.1	4.3	-9.1	-13.5	-6.7	-2.2
Rain Irrig ation Et Mean Et	0.5 36.0 46.7			36.0	18.0	18.0	18.0	18.0		27.0	0.5 27.0 20.8	

APPENDIX 7c (continued)

28.6.68 - 11.9.68

Soil		Fie	eld A			Fiel	ld B		İ	Pie:	ld C	
(cm)	٨٨	AB	AC	M	BE	EF	BG	Вн	CI	CJ	CK	CL
5 15 30 50 70 90 110 130 150 170	-0.5 3.1 7.2 3.0 0.6 0.4 -0.8 -1.2 -1.4 -0.6	4.8 7.3 7.0 2.6 0.0 2.2 ~1.2 0.4 0.8 ~5.3	7.7 4.9 18.8 17.4 11.0 8.4 7.0 6.6	5.1 -0.9 -6.2 -9.2 -7.2 -1.6 1.0 -1.2 -6.2 -7.6	7.7 3.4 -11.8 -6.4 -5.4 1.8 1.0 -0.2 -0.2 -2.0	8.2 4.0 -11.6 2.2 3.2 2.6 1.8 3.8 1.8	6.8 -1.7 10.4 6.6 3.2 2.8 3.0 1.4 3.0 2.0	7.6 1.7 5.4 4.4 2.5 2.2 1.6 1.8 1.4 0.8		3.3 6.0 -39.8 -1.6 -3.6 -6.4 5.2 11.4 12.4 23.4	6.8 4.7 -22.4 -5.8 -0.6 7.4 14.6 14.8 21.6 30.4	4.8 1.7 -28.8 -6.8 0.4 5.8 6.8 13.6 19.6 27.8
Total	9.8	18.6	97.8	-34.0	-12.1	15.0		29.7	94.7	-4.7	71.5	44.9
Rain Irrig ation Et Mean Et	51.8 42.0	60.6 42.0	139.8 42.0	-8.0 42.0	20.0 7.9	20.0 35.0	20.0 57.5	20.0 49.7 37.5	31.0 125.9	31.0 26.3	31.0 102.5	31.0 75.9 82.7
		21.8.6	8 - 11	.9.68	96.4	81.9	89.1	72.5				
				1	1.9.68	3 - 19	9.68					
5 15 30 50 70 90 110 130 150 170	9.4 5.5 6.4 5.2 2.0 1.0 1.6 0.8 1.2	9.8 4.4 6.8 5.6 5.4 1.8 2.0 0.8 0.6 2.2	8.7 6.8 -2.2 -5.2 -2.4 1.4 1.8 -0.8 -2.0 -7.2	8.6 10.4 18.2 16.4 10.2 3.8 -1.2 2.0 4.4 8.2	7.0 3.2 3.8 4.8 1.8 -2.0 0.4 -0.2 -0.6 -1.2		1.4: 3.6 -0.4 -1.0	7.3 2.6 2.2 4.4 -0.2 1.0	-6.8 -9.4 -14.2 -22.0	13.2 4.8 51.2 24.0 8.4 6.6 -3.0 -12.4 -12.2 -23.4	11.6 6.7 35.6 18.2 5.6 -7.0 -15.4 -15.4 -21.6 -29.2	11.0 6.3 42.6 14.6 3.2 -5.2 -7.2 -13.6 -21.4
Total	34.7	39.4	-1.1	81.0	17.0	14.2	16.0	21.3	-39.9	57.8	-10.9	-0.7
Rain Irrig ation Et Hean Et	34.7	39.4	-1.1	81.0	1 - 1 - 1	1 1 1	1 1 1	17.1	1 1 1	1 1 1 1	1 1 1 1	1 1 1 1
]	rriga	otal ation Et an Et	55.6 31.0 86.6	53.1 31.0 84.1	60.6 31.0 91.6	44.6 31.0 75.6 84.5

19.9.68 - 26.9.68

Soil		F1	ald A			Field	В			Fiel	La C	
(ont)	AA	ΛB	AC	VD	BE	BF	BG	ВН	UI	GJ	CK	CL
5 15 30 50 70 90 110 130 150	-4.2 -6.7 -9.6 -0.8 3.4 0.6 1.8 1.4	-5.7 -6.9 -5.6 -0.2 1.0 2.2 0.2 0.6 1.8 0.0	-12.1 -8.5 -11.8 -7.4 -4.2 -6.2 -2.4 -2.8 -5.0 -4.2	-11.2 -9.6 -13.4 -8.8 -1.4 0.8 1.0 0.2 1.4	-10.5 -7.7 1.4 4.0 4.8 4.0 0.6 1.0 0.4	-1.2 -3.4 -0.8 1.8 5.4 2.8 0.4 1.0 -1.0	-4.8 -3.9 0.6 2.4 2.6 1.2 0.2 -0.2	-8.2 -4.4 -4.0 -5.0 -6.8 -1.8 -2.6 -5.4 -2.2 -6.2	-11.2 -7.7 -16.4 -9.8 -2.2 -2.6 -0.8 0.2 -1.6	-8.7 -6.7 1.0 1.6 2.6 2.0 1.4 3.6 -0.6 2.2	-9.6 -5.3 0.2 0.2 2.6 1.8 1.2 0.4 1.0	-11.8 -3.3 1.0 3.6 3.0 0.8 2.6 2.2 7.0 8.4
Total	-1.2.9	-12.6	-65.2	-40.6	-1.0	5.6	-1.5	-46.6	-51.7	-1.6	-7.1	13.5
Rain	0.3	0.3	0.3	0.3	0,3	0.3	0.3	0.3	0.3	0.3	0.3	0.3
Irrig ation Et	57.2 44.6	36.6 24.3	120.7 55.8	48.8 8.5			15.7 14.5	23.9 -22.4	30.2 -21.2	36.8 35.5		39.4 53.2
llean Et				33.3	excl	uding	BH	24.4	exc)	udin	CI	37.5
				26.	9.68 -	3.10	.68					
5 15 30 50 70 90 110 130 150	4.0 5.8 6.2 4.0 3.8 4.2 1.4 0.4 -0.6	5.3 5.8 6.0 2.8 2.8 2.6 3.0 0.6 0.4 4.4	9.1 4.5 8.0 4.8 0.2 3.0 2.8 3.6	9.7 6.2 8.2 8.0 1.2 2.2 2.4 1.0	1.8 9.8 2.2 3.2 7.2 4.4 3.4 1.2	1.4 2.2 2.2 2.8 4.6 6.8 6.4 3.2 2.6 2.8	2.4 2.4 2.8 5.0 3.6 2.6 1.4	7.1 2.1 9.0 11.6 10.4 7.4 5.4 9.0 6.2 8.2	7.4 6.6 8.8 7.4 3.4 2.4 1.0 0.8 0.2	9.4 8.0 2.0 2.0 4.0 3.6 9.8 3.6 0.8	8.8 5.3 1.2 2.2 4.4 3.2 0.0 2.2 0.6 0.6	12.0 0.0 2.0 4.4 7.0 6.0 2.0 0.6 -3.6 -5.0
Total	29.8	33.7	50.6	44.1	39:0	35.0	27.2	76.4	39 • 4	46.6	28.5	25.4
Rain	3.4	3.4	3.4	3.4	3.4	3.4	3.4	3.4	3.4	3.4	3.4	3.4
Irrig ation			-	-		-		-	~	-	-	
Et	33.2	37.1	54.0	47.5	42.4	38.4	30.6	79.8	42.8	50.0	31.9	28.8
llean Et				43.0				47.8				38.4

3.10.68 - 15.10.68

Soil		Pic	oTg V			FLo	ld .			Fie	:ld C	
(ca)	ΔA	Lι	AC	ΔD	BC	ЪF	Ed	BH	CI	ÇJ	CK	CL
5 15 30 50 70 90 110 130 150	-14.5 -2.5 5.4 1.2 3.6 4.8 7.2 5.6 3.8 1.2	-16.1 -4.1 4.2 4.2 5.0 5.2 9.6 9.0 6.8 3.6	-13.8 -3.7 5.2 15.2 5.1 4.0 4.8 3.4 3.0 4.0	-14.2 -2.0 5.8 5.0 6.6 6.0 3.4 1.0 1.2	-15.9 -9.1 -2.1 0.0 2.0 8.0 5.8 4.2 3.8 1.0	-18.0 -3.0 -0.3 0.0 0.0 2.0 4.; 6.0 5.0	-18.8 -7.6 -0.4 1.8 4.2 4.0 2.8 1.8 2.8	-20.3 -0.7 0.8 3.0 5.2 3.4 4.6 0.8 2.0	-14.5 -0.4 5.2 5.8 6.0 7.4 5.2 3.1 4.6	-17.1 -5.2 -1.0 -0.8 2.0 6.2 9.6 1.0	-19.4 -0.9 -0.4 -2.0 1.4 7.0 2.8 -0.4 0.0	-16.0 2.2 0.6 1.2 3.4 8.6 10.2 5.6 5.8
Total	15.8	27.4	27.5	12.8	-2.3	-2.4	-8.6	-0.4	20.5	u.9	-11.9	26.8
Rain Irrig	30.7	30.7	30.7	30.7	30.7	30.7	30.7	30.7	30.7	30.7	30.7	30.7
ation Et	46.5	58.1	58.2	43.5	28.4	28.3	22.1	30.3	57.2	37.6	18.8	57.5
liean Et				51.0				1.5				42.8
					15.	10.63	- 31.1	0.68				
5 15 30 50 70 90 110 130 150	-1.8 -10.9 -17.4 -15.8 -18.4 -8.6 -3.0 2.4 2.4 5.6	-9.1 -20.6 -19.4	-13.0 -22.2 -10.0 -6.8 -4.2 0.0	-7.8 -15.6 -15.6 -11.0 -1.8 1.4 3.4 3.8	-9.5 -17.0 -20.6 -18.4	-11.0 -19.4 -20.4 -15.6 -4.0 2.0 -0.8 2.4	-6.5 -21.6 -22.8 -19.2 -11.0 -2.8 1.0	-20.2 -20.4	-12.1 -21.2	-0.5 -7.7 -16.4 -17.4 -11.8 -2.2 -1.6 -0.8 0.0 0.8	0.7 -12.6 -21.0 -16.8 -13.8 -4.2 1.8 1.6 2.4	-13.2 -23.4 -20.0 -14.2 -2.8 -0.4 1.6
	-65.5	-43.5		-39.7			⇒ 78.1			-57.6	1	
Rain Irrie	134.6	134.6	134.6	154.6	134.6	上 j-ç e ()	154.0	154.0	134.0	104.6	124.6	134.6
ation	-	-	-		~ CO. 4	70.0	56.5	52.1	67.1	77.0	73.9	64.1
Et Hean Et	69.1	91.1	81.4	94.9	58.4	70.9	56.5	59.5	07.1	17.0	12.9	70.6

31.10.68 - 19.11.68

Soil		Pi	eld A		1	Fi	eld B		-	Pi.	eld C	
(On.)	AA	AB	AC.	(A)	DI:	BP	133	1.11	CI	GJ	CK.	CL
5 15 30 50 70 90 110 130 150 170	-3.5 -2.4 -12.6 -3.2 0.6 -3.4 -6.0 -4.8 -1.0 -1.4	-2.9 -3.1 -8.6 0.0 2.6 -7.8 -8.4 -2.4 -1.8 -6.4	0.4 -1.2 0.2 1.2 0.6 -1.2	-0.4 -12.0 -1.4 -2.2 -5.0 -0.8 -2.0	-5.0 0.2 0.4 -2.4 -0.6	-1.2 -1.2 -8.8 -3.8 -2.0 -2.0	-6.4 0.8 -1.4 -3.6 -2.2	-3.9 -13.3 -5.6 -0.8 -1.8 -0.6 -1.4 -2.8	1.8	-2.8 -4.0 -5.4 1.6 -2.2 -8.4 -3.8 0.8 2.2 1.4	-7.0 0.6 -0.6 -7.8 -3.4	-7.4
Total	-37.7	-38.8	-17.7	-38.8	-37.8	-49.7	-34.3	-34.8	-16.2	-20.6	-26.9	-24.9
		104.8	104.8	104.8	104.8	104.8	104.8	104.8	104.8	104.8	104.3	104-8
Irrigation	67.1	66.0	87.1	66.0	67.0	55.1	70.0	70.0	88.6	84.2	77.9	79.9
Mean Et				71.6				75.5		1		82.7
					19.	11.68 -	29.11	1.68	•			
5 15 30 50 70 90 110 130 150 170	-14.B	1.5	4.1 5.4 -3.6 -3.2 -9.4 -9.6 -12.0 -10.0	5.1 6.0 -3.6	5.9 10.8 0.8 -6.4 -9.2 -13.6 -15.6	2.4 10.2 -5.2 -3.4 -12.6 -15.6 -17.0 -17.2	9.8 3.6 -4.6 -5.4 -8.6 -11.2 -12.0	3.1 9.6 2.4 -3.0 -7.0 -11.4 -12.4	3.3 7.8 -0.6 -4.2 -6.0 -9.4 -10.4 -14.8	1.1 2.8 -2.6 -4.6 -8.4 -14.4 -19.6 -19.2	2.2 5.4 -1.8 -3.4 -6.0 -8.5 -10.0 -13.0	2.5 7.4 -2.8 -7.8 -10.0 -13.8 -16.6
	-24.7		-44.4								-41.9	-75.2
Kain	220.3	220.3	220.3	220.3	220.3	220.3	220.3	220.3	220.3	220.3	220.3	220.3
Irrig ation Et	-	128.6	175.9	163.5	165.0	156.5	187.0	178.7	173.4	132.8	178.4	145,1
Mean		1		165.9				171.8				157.4

APPENDIX 7d

Soil moisture changes in Lysimeters A and B

EAL 104 Neutron Moisture Meter

Depth	23/7 -	31/7	31/7 -	7/8	7/8 -	1/8
(cm)	Lys. A	Lys. B	Lys. A	Lys. B	Lys. A	Lya. B
15 30 50 70 90 110 130	1.0) 1.0) 1.6 1.6 0.8 -0.2 -2.6 -3.0	Q.4 4.0 0.4 0.8 0.0 0.2 0.0	2.0) 2.0) 2.4 3.0 0.4 -2.0 0.0	2.4 0.0 3.0 3.8 2.0 0.8 0.6	0.2) 0.2) 1.0 -1.2 -0.8 1.0 0.4	2.8 2.4 -7.8 -1.4 0.2 -3.2 -0.4
Total Rain Irrigation Drainage Et Et column	0.2 8.4 - 6.8 1.8 25.3	4.6 8.4 12.7 0.3	5.8 2.8 1.5 7.1 9.4	11.4 2.8 - 14.2 3.8	1.8 1.8 25.0 2.1 26.5 28.6	-7.4 1.8 12.5 4.5 2.4
	21/8 -	28/8	28/8 -	11/9	11/9 -	19/9
15 30 50 70 90 110 130	-4.0) -4.0) -1.6 -0.6 1.6 0.0 1.8 4.0	3.2 1.4 6.8 1.4 -0.2 0.6 0.2	-1.3) 7.8) 4.6 1.2 -0.6 1.0 0.8 0.2	23.8 4.4 4.8 2.6 2.6 4.8 4.2	13.9) 4.8) 4.8 6.2 0.8 1.2 -3.0 -2.6	-7.4 . 2.6 0.8 -2.4 -0.8 -0.4
Total Rain Irrigation Drainage Et Et column	-2.8 0.5 36.0 - 33.7 17.1	13.4 0.5 18.0 0.5 31.4 44.0	13.7 42.0 - 55.7 60.4	45.2 20.0 - 65.2 41.6	26.1 - - 26.1 40.1	-11.8 - - -11.8 34.4

EAL 104 Neutron Moisture Meter

Depth	19/9 -	26/9	26/9	- 3/10	3/10 .	- 15/10
(cm)	Lys. A	Lys. B	Lys. A	Lys. B	Lys. A	Lys. B
15 30 50 70 90 110 130	-7.8) -7.8) -1.6 0.2 1.8 1.0 3.4 2.2	4.4 4.2 4.0 3.6 2.4 -0.6 3.2	6.6) 6.6) 1.6 2.2 2.4 1.8 2.6	5.0 5.6 5.8 7.6 4.8 4.0	4.2) 4.2) 3.4 4.0 3.8 4.0 7.4	1.0 1.8 2.2 3.8 14.2 21.2
Total Rain Irrigation Drainage Et Et column	-8.6 0.3 71.0 62.7 50.2	21.2 0.3 22.4 43.9	24.4 3.4 - 27.8 52.4	30.6 3.4 - 34.0	43.0 30.7 - 73.7 81.3	58.8 30.7 - 89.5
	15/10	- 31/10	31/10	- 19/11		
15 30 50 70 90 110 130	-16.2) -16.2) -17.0 -15.4 -7.0 1.8 6.2 9.6	-32.0 -17.8 -11.6 0.6 -0.2 3.8 12.4	-7.8) -7.8) -1.0 -0.6 0.2 -0.6 -4.4 -3.4	-2c -7.0 -1.0 -7.8 -3.8 -2.2 0.2		
Total Rain Irrigation Drainage Et Et column	-54.2 134.6 - 80.4 118.6	-44.8 134.6 - 89.8	-25.4 104.8 - 79.4 101.9	-46.4 104.8° - 58.4	-	

APPENDIX 7e

Summary of soil moisture changes (nm) at different depths in the soil profile

NIV 96 Neutron Moisture Meter

7.8.68 - 21.8.68

So11		Fie	1d A			Fie	ld B			Fie:	ld C	
depth (cm)	AA	АВ	AC	AD	BE	BF	БO	PH	CI	CJ	СК	CT
5 15 30 50 70 90 116 130 150	-3.3 -2.0 18.6 14.6 16.4 20.6 30.2 24.0 34.0 26.2	-5.8 24.4 25.8 19.0 25.6 23.2	-0.8 27.0 13.0 15.2 16.2 14.4 19.0	2.8 22.8 25.0 16.8 24.8 28.2	-2.6 -0.8 -1.8 -14.4 -0.2 -1.8	0.8 -15.6 -8.2 -4.6 -20.6 -5.2 -7.4 -5.2	1.2 -13.0 -7.8 -10.2 -0.8 -8.2	1.1 3.6 -2.0 -7.4 -5.8 -6.2 -11.2 -0.6				
Total	179.3	157.8	163.2	181.7	-36.6	-65.0	-32.5	-34.2				
Rain	5.3	5.3	5.3	5.3	1.8	1.8	1.8	1.8				
Irrig ation	25.0	25.0	25.0	25.0	12.5	12.5	12.5	12.5				
Et	209.6	188.1	193.5	212.0	-22.3	-50.7	-18.2	-19.9				
Mean Et				200.8				-27.15				
				1	21.8.6	3 - 28	8.68					
5 15 30 50 70 90 110 130 150 170	1.4 -1.8 -4.6 -0.6 9.4 -10.6 3.8 0.0 -2.2	-4.6 -6.0 3.0 -1.2 10.8 0.4 15.0	-9.9 -3.2 -19.2 -1.2 -3.2 -12.4 5.4 1.0 -8.4 -6.0	-1.4	3.5 8.4 0.6 ~7.0 0.0 10.6 4.2	-0.7 17.4 8.2	-8.6 1.5 -6.8 5.8 8.4 -1.0 4.2 4.4 -2.8 -5.2		-8.6 -5.1 7.0 2.0 16.4 3.0 8.0 5.8 1.0 9.8	-6.9 -7.2 -6.0 13.4 13.8 8.2 -4.4 12.2 -4.8 5.8	-5.1 -1.6 6.0 3.8 13.0 14.8 9.2 11.6 1.4	-7.9 3.3 -8.6 1.2 2.0 -2.0 -7.6 0.0 8.0 4.2
Total	-3.6	0.1	-57.1	-32.8	10.0	42.4	-0.1	6.7	39.3	24.1	57.9	-7.4
Rain Irrig ation	36.0	0.5 36.0	36.0	0.5 36.0	0.5	0.5	0.5	0.5	0.5 27.0	0.5	0.5	27.0
Et Mean Et	32.9	36.6	-20.6	3.7	28.5	60.9	18,4	25.2 33.3	66,8	51,6	85.4	20.1 56.0

NIV 96 Neutron Moisture Meter

28.8.68 - 5.9.68

2.2 6.7 20.0 12.2 0.2 14.0 3.8 10.2 8.0 89.1	14.9 13.6 16.0 12.0 5.8 12.0 -0.2 -2.0 8.0	AC 16.4 16.9 20.0 8.0 14.0 8.0 2.0 6.0 131.3		0.0 -2.0 -12.2 -7.8	0.0 6.0 -12.0 -12.2 -14.0 -15.8 -12.2	-12.2 -8.2 -10.0 -12.2 -10.0 -4.0	5.8 -6.0 -10.2 -15.3 -12.0 -12.2 -8.0 -6.0 -4.0	2.2 -2.2 -2.0 -8.2 -8.0	10.5 8.0 -4.0 -5.8 0.0 -8.0 -10.0 -1.8 -12.0	-12.2 -2.0 2.2	17.2 3.9 7.8 -4.0 -8.0 4.0 -2.2 -1.0 -14.0 -9.8
6.7 20.0 12.2 0.2 14.0 3.8 11.8 10.2 8.0	13.b 16.0 12.0 5.8 12.0 2.0 -0.2 -2.0 8.0	10.9 20.0 8.0 14.0 22.0 8.0 2.0 6.0 131.3	9.1 18.0 -2.2 2.0 0.0 -4.2 -2.0 12.0	0.0 0.0 -14.0 0.0 -2.0 -12.2 -7.8 4.0 -11.8	3.9 -8.0 0.0 6.0 -12.0 -12.2 -14.0 -15.8 -12.2	8.4 9.8 -7.8 -12.2 -8.2 -10.0 -12.2 -10.0 -4.0	5.8 -6.0 -10.2 -15.3 -12.0 -12.2 -8.0 -6.0 -4.0	16.3 4.0 0.2 -18.0 2.2 -2.2 -2.0 -8.2 -8.0	10.5 8.0 -4.0 -5.8 0.0 -8.0 -10.0 -1.8 -12.0	9.3 2.2 2.2 -6.2 -1.8 -8.2 -12.2 -2.0 2.2	3.9 7.8 -4.0 -8.0 4.0 -2.2 -1.0 -14.0 -9.8
-	-	-	41.3	-35.6	-53.8	-36.6	64 1	00	24.2	1 6	1 3 1
-	-		-			,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,,	-04.4	0,9	-14.1	٠.٥	-0.1
	-	[-	-	-	-	-	-	-	-
89.1		_	-	-	-	-	-	-	-	-	-
	82.1	131.3	41.3	-35.6	-53.8	-36.6	-64.4	0.9	-14.1	0.6	-6.1
			86.0				- 7.0				-4.7
			5	.9.68 -	. 11.9.	68					
-2.7 -3.6 -6.0 -6.2 -0.2 0.0 -2.0 -2.0 -2.2 -4.0	-6.3	-12.0	4.0 6.0 12.0 10.0 6.0	3.4 3.8 18.0 13.8 6.0 12.2 6.0	0.1 8.0 18.0 4.0	-2.8 -10.1 6.0 21.8 18.0 8.2 8.2 16.2 14.0 12.2			-5.7 -4.5 18.0 6.0 7.8 4.0 6.2 10.0 11.8 14.0	-8.3 4.6 15.8 -4.0 4.0 -6.2 10.0 6.0 0.0	-12.4 -2.2 6.0 13.8 8.0 3.8 8.0 7.2 15.8 14.0
-28.9	-35.8	-33.9	48.7	88.7	123.8	91.7	111.7	48.2	67.6	12.7	62.0
42.0	42.0	42.0 6.1	42.0 90.7 29.5			- 20.6 111.7	- 20.0 131.7	- 31.0 79.2	31.0 98.6	31.0	31.0 93.0 78.6
	-3.6 -6.0 -6.2 -0.2 0.0 -2.0 -2.0 -2.2 -4.0	-3.6	-3.6	-2.7	-2.7 -10.1 -10.7 -1.3 -0.5 -3.6 -6.3 -12.0 4.6 3.4 -6.0 4.6 -11.8 -4.0 3.8 -6.2 2.0 6.0 6.0 18.0 -0.2 -1.6 -4.0 12.0 13.8 0.6 -7.8 -5.8 10.0 6.0 -2.6 -9.8 0.0 6.6 12.2 -2.6 -4.0 -1.8 14.6 6.0 -2.2 3.8 0.0 -1.8 12.0 -4.0 -5.8 6.2 3.8 14.6 -28.9 -35.8 -33.9 48.7 88.7	-2.7 -10.1 -10.7 -1.3 -0.5 -2.3 -6.3 -12.0 4.6 3.4 0.1 -6.0 4.6 -11.8 -4.0 3.8 8.0 -6.2 2.0 6.0 6.0 18.0 18.0 -0.2 -1.8 -4.0 12.0 13.8 4.0 0.0 -7.8 -5.8 10.0 6.0 18.0 18.0 -2.6 -9.8 0.0 6.6 12.2 18.0 -2.6 -9.8 0.0 6.6 12.2 18.0 -2.2 3.8 6.0 -1.8 12.0 10.0 -4.0 -5.8 6.2 3.8 14.6 18.6 -2.8 9 -35.6 -33.9 48.7 88.7 123.8 -7 -7 -7 -7 -7 -7 -7 -7 -7 -7 -7 -7 -7	-2.7 -10.1 -10.7 -1.3 -0.5 -2.3 -2.8 -3.6 -6.3 -12.0 4.7 3.4 0.1 -10.1 -10.1 -6.0 4.6 3.8 8.0 6.0 -6.2 2.0 6.0 6.0 18.0 18.0 21.8 -0.2 -1.8 -4.0 12.0 13.8 4.0 18.0 21.8 -2.6 -7.8 -5.8 10.0 6.0 16.0 8.2 -2.6 -9.8 0.0 6.6 12.2 18.0 8.2 -2.6 -4.0 -1.8 14.0 6.0 14.0 16.2 -2.2 3.8 6.0 -1.8 12.0 10.0 14.0 -5.8 6.2 3.8 14.0 18.0 12.2 -2.2 3.8 -3.5 6.2 3.8 14.0 18.0 12.2 -2.2 3.8 6.2 3.8 14.0 18.0 12.2 -2.0 -5.8 6.2 3.8 14.0 18.0 12.2 -2.0 -5.8 6.2 3.8 14.0 18.0 12.2 -2.0 12.0 42.0 42.0 42.0 20.0 20.0 20.0 20.0 13.1 6.2 6.1 90.7 108.7 123.8 111.7	-2.7 -10.1 -10.7 -1.3 -0.5 -2.3 -2.8 3.6 -6.3 -12.0 4.6 3.4 0.1 -10.1 -4.1 -4.1 -6.0 4.6 -11.8 -4.0 3.8 8.0 6.0 10.0 -6.2 2.0 6.0 6.0 18.0 18.0 21.8 16.0 0.0 -0.2 -1.8 -4.0 12.0 13.8 4.0 18.0 19.8 0.0 -7.8 -5.8 10.0 6.0 16.0 8.2 16.0 -2.6 -9.8 0.0 6.6 12.2 18.0 8.2 16.0 -2.6 -9.8 0.0 6.6 12.2 18.0 8.2 16.0 -2.2 3.8 6.0 -1.8 12.0 10.0 14.0 12.0 -4.0 -5.8 6.2 3.8 14.6 18.6 12.2 12.2 12.2 -4.0 -5.8 6.2 3.8 14.6 18.6 12.2 12.2 12.2 12.2 12.2 12.2 12.2 12	-2.7	-2.7 -10.1 -10.7 -1.3 -0.5 -2.3 -2.8 3.6 -11.4 -5.7 -3.6 -6.3 -12.0 4.6 3.4 0.1 -10.1 -4.1 -17.2 -4.5 -6.0 4.6 -11.8 -4.0 3.8 8.0 6.0 10.0 6.6 18.0 -6.2 2.0 6.0 6.0 18.0 18.0 21.8 16.0 14.0 6.0 -0.2 -1.8 -4.0 12.0 13.8 4.0 18.0 19.8 15.8 7.8 0.0 -7.8 -5.8 10.0 6.0 16.0 8.2 16.0 14.0 4.0 -2.6 -9.8 0.0 6.6 12.2 18.0 8.2 16.0 14.0 4.0 -2.6 -9.8 0.0 6.6 12.2 18.0 8.2 16.0 2.2 6.2 -2.6 -2.2 3.8 6.0 -1.8 12.0 10.0 14.0 12.0 10.6 11.8 -4.0 -5.8 6.2 3.8 14.6 18.6 12.2 12.2 4.0 14.0 14.0 -5.8 6.2 3.8 14.6 18.6 12.2 12.2 4.0 14.0 14.0 12.0 10.6 11.8 -4.0 -5.8 6.2 3.8 14.6 18.6 12.2 12.2 4.0 14.0 14.0 14.0 14.0 14.0 15.0 10.6 11.8 -4.0 -5.8 6.2 3.8 14.6 18.6 12.2 12.2 4.0 14.0 14.0 14.0 14.0 14.0 14.0 14.0	-2.7

APPENDIX 8

Record of tensioneter readings at 0900 hrs. September - December, 1967

*Negative values, i.e. tensioneter reading less than the contribution from the hanging column of water as measured from mercury reservoir to the depth of tensioneter cup.

L = additional lower tensioneter in lysimeter A: 180 cm.

L = Lower tensiometer: 170 cm in field

150 cm in lysimeter

U = Upper tensiometer: 120 cm in field 100 cm in lysimeter

IRR = fields irrigated.

		Field A			Lysineta	P .n		Field B		I	ysimeter	P
Tate Lyu7	L cm water	U cm water	(L-U) (50) cm/cm	L cm water	U cm water	(L-U 50+1) cm/cm	L cm water	U cm water	(L-U-1) (50+1) cm/cm	cm water	J cm= water	(L-U-1) (50-1) cm/cm
Sep. 1	119.7	149.6	. 40	24.5	129.2	-1.09	194.5	240.2	03	43.5	144.2	-1.01
2	115.6	145.5	.40	36.7	125.1	76	180.9	190.4	.81	42.2	127.8	71
3	112.9	149.6	- 26	21.8	61.6	19	100.6	153.1	1.19	35.4	95.2	19
4	114.2	152.3	.23	27.2	88.4	22	161.8	151.0	1.21	38.1	95.2	14
5	117.0	157.8	.18	34.0	107.4	46	161.8	140.9	1.29	59.8	129.2	36
6	114.2	155.0	.18	23.1	61.6	17	153.7	131.9	1.03	35.4	69.8	08
7	117.0	159.1	-15	29.9	100.5	-, 11	153.7	144.2	1.19	39.4	107.4	30
8	122.4	164.6	-15	32.6	85.7	05	163.2	176.2	.70	36.1	107.4	38
9	122.4	163.2	.18	16.3	76.9	05	161.8	179.5	-64	32.6	100.0	36
10	122.4	161.6	.21	20.4	73-4	06	153.7	175.4	.56	17.7	81.6	27
11	123.2	161.8	-34	23.1	81.6	17	163.2	175.4	.75	27.2	119.7	05
12	129.2	163.2	.32	20.4	77.5	14	160.5	179.5	.62	29.9	106.1	52
IRR 13	129.2	172.7	.13	25.8	80.2	C8	155.0	174-1	-61	35.4	107.4	44
14	129.2	157.3	.23	15.3	73.4	14	156.4	182.2	. 48	32.6	97.9	30
15	126.5	155.0	. 43	24.5	70.7	+.07	159.1	164.9	. 48	36.7	107.4	41
16	121.0	138.7	.64	32.6	80.2	+.04	152.3	175.4	-53	31.3	111.5	60
17	118.3	141.4	-53	29.9	84.3	08	152.3	174.1	.56	31.3	111.5	50
18	122.4	149.5	. 45	27.2	68.4	22	149.6	-	-	27.2	116.9	79
19	115.5	151.0	.29	28.5	84.3	11	149.6	_	-	27.2	103.4	52
20	119.7	137.4	-64	24.5	133.3	-1.17	164.6	167.3	-94	53.0	141.4	-1.16

APPENDIX 8 (continued)

	,					PPENULA O	(0020222	- 47		,		
		Field A		1	Lysimeter	A		Field B		I	ysimeter	В
Date 1967	L cm weter	cm water	(1-U 50 1) cm/cm	L cm water	U cm water	(L-U-1) (50 +1) em/em	L cm water	U CD Water	(L-U 1) (50 +1) em/em	L weter	U cm water	(L-IJ+1) en/em
Sep. 21 22 23 24 25 1RR 26 27 29 30 21 25 29 30	118.3 119.7 113.3 117.0 121.0 125.6 123.8 123.1 121.0 125.1 125.1 125.1 125.1 125.1 125.1	151.0 156.4 155.0 140.9 157.8 157.8 161.8 164.6 163.2 160.5 161.8 170.0 154.6 165.9 163.2 163.2	.34 .26 .26 .40 .26 .47 .24 .26 .25 .23 .21 .26 .10 .21 .18 .23	27.2 27.2 24.5 21.8 21.6 19.0 20.4 23.1 21.7 19.0 19.0 19.0 19.0 19.0 19.0 19.0	111.5 107.4 110.2 100.6 122.4 105.6 119.7 118.3 111.5 97.9 100.6 100.6 112.9 97.9 114.2 99.3 103.4 99.3	66 60 71 57 -1.01 59 95 76 52 63 63 57 57 57 55 85 85	167.3 185.9 169.6 167.3 159.1 176.8 193.1 176.8 160.5 159.1 148.2 141.4 142.8 138.7	179.5 163.6 161.8 160.5 118.3 121.0 115.6 191.8 155.0 157.6 141.4 133.3 125.1 131.9 174.1 131.9 118.3	.75 .04 1.13 1.13 1.31 1.70 2.22 1.43 1.19 1.40 1.54 1.68 1.54 1.19 1.21	46.2 40.8 35.7 35.4 40.8 43.5 50.3 57.1 53.0 55.8 42.3 47.6 44.9 50.3 47.6 44.9	123.8 117.0 110.2 107.4 123.8 118.3 117.0 107.4 100.1 102.0 96.6 100.6 96.5 100.6 97.9 102.0 99.3	55 52 47 44 66 49 33 19 08 .00 19 +.02 06 03 .00
T0.	122.1	160.5	.23	4.1 5.4	95.2 92.5	82 74	133.3	93.8 118.3	1.79	47.6 46.2	95.2 97.9	+.04
IRR 11 12 13 14 15 16 17 18 19 20	140.9 144.2 142.8 141.4 141.4 141.4 152.3 146.9	195.8 196.5 183.0 185.0 186.3 186.3 187.7 179.5 176.8 175.4	.02 08 .18 .12 .10 .10 .07 .45 .40	8.2 8.2 2.7 2.7 5.4 4.1 5.4 5,4	96.6 96.6 91.1 95.2 97.9 92.5 102.0 93.8 104.7	76 76 76 85 85 76 95 76 98 -1.12	167.3 174.1 163.2 163.2 163.2 170.0 160.5 153.7 163.2 165.9	103.4 99.3 69.8 65.7 42.2 104.7 30.1 16.3 20.4 17.7	2.27 2.49 2.46 2.55 3.42 2.30 3.44 3.74 3.85 3.96	59.8 62.6 57.1 40.8 57.1 58.5 59.8 59.8 63.9 63.9	151.0 130.0 134.6 117.0 121.0 114.2 119.6 118.3 119.7	82 46 55 51 27 11 19 17 11

APPENDIX 8 (continued)

		Field	A		L	/sllete	r A			Field	9	L	ysime t	r B
Date 1967	L cm water	U cn water	(1-7+1) (5C+1) em/em	L cr water	L† cm weter	U cm water	() () cm/cm	(L'-L) (30) cm/cm	L cm water	Cm water	(<u>L-U</u> +1)	cm kater	om water	(<u>L-U</u> +1)
	140.9 149.6 150.9 145.5 145.5 129.7 117.0 125.6 Ce.4 E9.a E7.0 74.8 b0.2 B7.0 67.0 67.0	174.1 179.5 162.2 183.6 -2.7 170.0 188.6 125.1 108.6 112.9 99.3 110.2 95.2	.42 .51 .34 .45 .53 .89 1.16 .73 .59 .59 .56 .56 .58	4.1 1.4 1.4 2.7 1.4 0 ** 35.4 40.8 40.8 40.8 40.8		104.7	-193 -1.28 -1.12 -1.06 -1.23 84 90 	1.40 1.45 1.49 1.56 1.49 1 00 	165.9 163.2 171.4 166.6 166.6 170.0 175.4 1120.5 122.4 118.3 104.7 104.7 104.7 195.2 97.9 96.7 106.1	35.4 46.2 55.8 72.1 46.2 29.9 69.4 57.1 63.9 46.2 74.8 58.5 59.2 69.2 69.2 69.2	3.61 3.34 3.31 2.93 3.44 3.77 3.01 2.35 2.60 2.44 1.59 2.14 1.65 1.70 1.97 1.98 2.98	61.2 65.3 63.9 65.3 65.3	73.4 65.3 69.4 74.8 62.6 68.0	17 19 17 17 14 06 08 +.23 +.07 .00 .07 .02 .03 00 03 06 06
	100.6 100.6 106.1 112.8 121.0 121.0	142.8 146.9 160.5 160.5 165.9 172.7 174.1	.29 .23 .15 .16 .04 .21 .10 .04	35.4 39.4 39.4 36.7 -39.4 32.6 39.4 34.0	10.9	60.2 84.3 84.3 83.0 121.0 107.4 93.8 100.6 97.9 107.4	.10 .10 .07 36 22 22 27	0 04 13 04 - 0 +.27 +.05 +.16 +.13	106.5 111.5 118.5 121.0 140.1 138.7 137.4 145.5 149.6 152.3	56.5 53.0 57.1 55.6 9.5 63.0 81.6 35.4 15.0 16.3	2.00 2.17 2.22 2.30 3.61 2.11 2.11 3.18 3.69 3.52	44.9 40.8 47.6	110.2	11 14 22 19 17 19 17 17 12 22

APPENDIX 8 (continued)

		T					APPENDIA	. 0 (052			-			
		Field A				Lysimete	r A			Field .	В		ysime te	r B
Date 1967	L cm water	U cm water	\left(\frac{L-U}{50} + 1\right)\) cm/cm	L cm water	L' cm water	U cm water	(I-U-1) (50-1) em/em	(<u>L'-L</u>) em/cm	L cm water	U cm water	$\left\{ \frac{L-U}{50}+1\right\}$ cm/cm	L cm water	t cm water	$\left\{ \frac{L-U}{50}, 1 \right\}$ em/em
Nov. 21 22 23 24 25 26 27 2 2 3 Dec. 1	129.2 129.2 131.9 130.6 130.6 129.2 16 125.2 114.2 107.4 126.5 107.4	176.8 178.2 183.6 179.5 179.5 179.5 172.7 151.0 140.1 133.3	.04 .02 C1 .C2 .02 00 15 .56 .48 .48	34.0 4.0 35.4 34.0 32.6 34.0 28.6 27.2 24.5 25.9 31.3 28.6 25.8 25.8 25.8	10.9 9.5 9.5 9.5 9.5 6.8 5.4 5.4 5.4 9.5 12.2 9.5 6.8 13.6	104.7 103.4 104.8 95.2 91.1 89.8 65.3 69.4 72.1 76.2 91.1 91.1 95.2 100.6 89.8	413838221711 +.26 +.15 +.04 01925384922	.23 .18 .13 .16 .23 .18 .27 .27 .36 .31	152.3 155.0 152.3 151.0 151.0 151.0 153.7 144.2 125.1 129.2 127.8 129.2 131.9	36.7 84.3 70.7 27.2 29.9 15.0 16.3 10.9 10.9 88.4 77.5 12.2 0	3.31 2.41 2.63 3.47 3.42 3.72 3.74 3.66 3.28 1.81	51.7 53.1 50.3 47.6 48.9 47.6 44.9 32.6 29.9 29.9 38.1 39.4 38.1 40.8	114.2 115.6 110.2 106.1 104.7 104.7 93.8 81.6 60.2 81.6	25 25 19 17 11 14 +.02 +.02 0 03
6	115.6 118.3 122.4 125.1 126.1 127.8 130.6 134.6	161.8 165.9 174.1 176.8 178.2 183.6 186.3 191.8	.07 .04 63 03 03 11 11	28.6 29.9 32.6 34.0 35.4 35.4 36.7	8.2 15.0 16.3 24.5 23.1 23.1 24.5 25.8	89.8 92.5 96.6 100.6 100.6	22 25 28 33 33 33 36 36	.32 .50 .45 .68 .83 .59 .63	137.4 138.7 141.4 144.2 145.5 148.2 149.6 152.3	42.2 49.0 16.3 15.0 9.5 76.2 10.9 5.4	2.90 2.79 3.50 3.58 3.72 2.44 3.77 3.93	36.7 44.9 48.9 54.4 53.0 55.8 59.8	92.5 99.3 102.0 109.4 107.4 110.2 114.2 117.0	11 03 -1.05 06 08 08
14 15 16 17 18 19 20	131.9 137.4 140.1 142.8 144.4 144.2 151.0	194.5 193.1 201.3 209.4 213.5 223.0 228.5	25 11 17 33 44 57 55	36.7 39.4 42.2 43.5 43.5 47.6 50.3	27.2 29.9 32.6 34.0 32.6 39.4 42.2	102.0 107.4 110.2 112.9 114.2 118.3 121.0	26 36 36 38 41 41	.68 .68 .68 .63 .72	153.7 157.8 160.5 160.5 160.5 163.2 164.6	9.5 27.2 8.2 43.5 19.0 23.1 28.6	3.88 3.61 4.04 3.34 3.83 3.80 3.72	72.1 74.8 80.2 84.3 91.1	118.3 123.6 127.6 133.3 136.7 145.5 152.3	06 03 06 06 08 08

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APPENDIX 8 (continued)

SOIL MOISTURE TENSION (CM. WATER) RECORDED WITH TENSIONETERS PLACED AT DIFFERENT DEPTHS IN THE SOIL PROFILE

5 .	Field A			Lysimeter A			Field B		Lysime	ter B		Field C	
Date	170cm.	150cm.	70cm.	175 cm.	150cm.	17º em.	150cm.	70 cm.	170cm.	150cm.	120 cm.	100cm.	50 cm.
AUG. 1968													
1 2 3 4 5	137 145 141 137 137	140 157 155 149 169	203 240 226 206 226	22		146 134 134 127 136	170 163 109 106 109	212 195 217 219 226		47 , 34 45 48 44	286 264 297 286		286 264 264 198 220
6 7 8 9	140 140 149 149 150	154 154 161 160 163	217 212 226 224 235	40 39 40	60 64	132 137 139 132 132	166 168 171 164 163	229 235 235 239 237		42 47 44 43	297 308 308 308 308	165 242 253 253 264	242 264 264 286 286
11 12 13 14	151 150 150 151 151	163 156 161 160 160	235 224 221 219 217	56 56 46 47 39	= 88 60	129 115 124 120 137	144 131	239 220 237 239 253		40 25 48 43 50	297 266 308 320 331	253 242 264 264 264	254 253 297 297 308
16 17 18 19 2	153 153 153 154 154	160 150 159 159 160	221 220 225 229 229	39 35 37 42 40	66 55 64 58	132 137 136 136 141	139 184 193 190 169	249 253 249 244 243		45 50 44 47 49	308 331 320 331 331	132 275 275 285	297 331 331 308 331
21 22 23 24 25	155 154 155 154 156	160 157 157 157 157	230 229 229 227 227	37 39 44 42	55 58 62 63 62	143 121 137 141 140	192 190 188 188	235 216 202 197 197		52 49 50 49 47	331 331 331 331 331	286 286 308	331 331 331 331 308
26 27 28 29 30 31	168 146 158 146 159 153	157 160 159 164 155 152	230 235 235 231 221 219		62 58 52 63 57 59	140 146 146 129 77 73	185 189 189 179 183	200 211 215 182 216 215		52 58 52 - 52	331 353 331 297 331 331	220 286 286 253 264 198	253 209 264 242 320 331

APPENDIX 8 (continued)

Date	Field A			Lysimeter A			Field B		Lysime	ter B		Field C	
	170cm.	150сш.	70cm.	170сш.	150cm.	170cm.	150 cm.	70 cm.	170сш.	150cm.	120 cm.	100cm.	50 cm.
SEP. 1968													
1 2 3 4 5	154 154 148 149 151	155 155 151 155 154	217 221 216 217 229			20	183 186 180 185 188	216 235 207 212 215	54 52 53 58 54		308 331 331 342 342	275 220 2÷2 220	331 342 320 331 342
6 7 8	164 161 161 165 169	169 174 17' 174 175	240 254 246 249 241	49 39 38 47 61		165 61 39 33 11	231 219 204 210 215	237 270 264 260 226	56 51 49 54 62		342 353 364 353 331	253 286 320 331 286	331 364 375 386 331
11 12 13 14 15	166 163 161 151 164	171 166 165 165 164	239 236 229 241 256	48 46 54 51 47	=	30 59 35 28 25	205 202 197 197 199	248 260 251 251 280	56 56 59 61 32		353 353 331 342 364	320 264 286 275 286	375 397 364 364 397
16 17 18 19 20	161 160 159 160 165	161 163 160 159 163	260 268 274 274 287	44 47 47 49 62		32 49 59 61 18	199 199 199 196 204	297 312 331 347 350	56 58 57 73	-	364 364 264 353 375	353 353 264 353 331	408 408 408 375 441
21 22 23 24 25	151 161 160 161 160	175 161 161 157 159	263 255 255 243 246	59 53 49 54 53		18 6 37 29	205 200 199 198 198	345 319 375 403 434	76 71 67 71 72		386 353 386 375 397	463 463 474 441 463	441 430 452 463 463
26 27 28 29 30	161 164 165 161 166	160 163 159 157 160	253 263 274 277 308	61 59 61 63		66 57 74 73	195 171 195 200 200	459 482 517 534 554	74 78 81 83 68		386 397 397 375 397	452 441 496 518 331	441 397 386 375 364

APPENDIX 8 (continued)

2	Field A			Lysime	ter A		Field B		Lysime	ter B		Field C	
Date	170cm.	15° c=.	70 cm.	170cm.	150 cm.	170 cm.	150cm.	70cm.	170cm.	150cm.	120cm.	100еш.	50еш.
OCT. 1968													
1 2 3 4 5	165 166 169 172 173	159 160 166 161 155	333 369 392 429 474	64 66 68 72 77	64 82 86 92 96	15 6 27 46 42 72	203 204 205 205 210	588 502 540 561 575	92 95 103 111 122	266 300 318 331 349	408 419 419 419 419 441	275 551 353 485 375	507 529 529 540 562
6 7 8 c	168 175 177 178 162	156 161 168 170	5° 2 505 484 500 566	77 82 86 88 92	93 100 103 106	40 47 87 77 66	210 209 209 209 212	563 568 598 497 552	129 143 155 170 189	352 357 360 359 244	397 452 463 463 496	573 386 672 397 573	573 584 419 354 496
11 12 13 14 15	180 185 189 190 194	151 169 163 155 160	505 618 628 629 559	97 103 110 119 124	146 152 155	111 111 69 90 59	205 205 212 214 215	564 575 586 590 596	196	194 242 330 364 388	485 485 507 518 551	683 397 639 397 661	551 573 617 617 650
16 17 18 19 20	198 199 2°3 2°4 197	155 149 142 154 146	588 595 599 529 594	131 135 140 144 146	167 163 168 173 170	155 187 200 208 206	213 209 204 213 193	598 598 598 598 598	140 172 165 195 200	397 407 408 409 398	529 551 529 465 485	617 650 716 606 639	661 661 485 507
21 22 23 24 25	207 211 212 209 216	146 147 147 142 156	595 590 487 522 544	153 160 169 169 180	179 184 169 190 199	212 213 182 217 224	189 185 183 178 174	598 596 463 527 550	208 211 214 217 218	407 402 403 399 399	474 551 551 551 573	663 738 573 661 705	485 435 672 661 683
26 27 28 29 30 31	218 206 217 206 187 178	161 169 165 169 165 161	554 474 512 474 437 403	187 248 197 208 208 200	204 163 205 163 126 117	231 235 233 235 237 233	170 160 165 160 159 154	556 459 510 459 419 382	221 223 226 223 102 150	398 394 396 394 328 374	573 573 573 573 562 551	551 331 683 331 353 364	683 617 639 617 595 573

APPENDIX 8 (continued)

2.37	Field A			Lysime	ter A		Field B		Lysine	ter B		Field C	
Date	170cm.	150 cm.	70сп.	170cm.	150cm.	170cm.	150cm.	70cm.	170cm.	150cm.	120 cm.	100cm.	50 cm.
NOV. 1968													
1 2 3 4 5	169 172 168 168	155 146 141 136 132	376 356 338 318 300	130 182 173 168 164	117 120 118 108 101	233 227 226 221 217	154 147 146 139	356 331 312 285 265	180 199 207 223 233	382 375 370 358 348	551 551 551 562 573	364 375 375 364 353	661 529 507 485 463
6 7 8 10	163	132 126	288 275 266 166 205	153 140 134 125 112	96 93 93 91 83	218 213 208 207 199	137 130 126 122	253 234 222 214 200	245 247 243 120 168	344 333 325 313 305	584 463 551 562 551	331 331 308 308 286	452 441 419 419 397
11 12 13 14 15		160 170 188 180 165	240 244 259 245 245	114 146 143 144	86 82 88 87 88	195 197 193 188 188	117 120 117 115 112	197 198 200 198 196	192 211 221 232 240	296 296 291 287 283	551 551 551 529 518	256 266 266 286 275	397 397 397 375 364
16 17 18 19 20	_	160 1 5 9 156 155 166	254 260 261 265 ±46	146 148 146 146 156	91 91 88 92 154	185 182 180 184 145	112 108 108 110 127	198 202 206 217 135	240 242 246 252 342	281 278 277 278 425	518 518 507 518 507	264 253 253 242 220	353 353 342 353 331
21 22 23 24 25		161 160 156 146 142	94 51 108 35 72	146 44 33 35 39	97 45 45 48 55	136 1:0 111 116 122	136 126 118 130 136	71 85 105 111 111	384 387 367 67 51	262 10 24 25 28	507 375 242 242 242	44 77 88 88	331 320 320 286 266
26 27 28 29 30		140 135 132	108 125 139	34 34 34	47 50 52	112 117 117	123 131 132	101 117 132	42 40 34	23 30 34	242 264 220	88 88 99	286 264 275

Date	Field A			Lysime	ter A		Field B			ter B	-	Field C	
Date	170cm.	150 cm.	70 cm.	170cm.	150 cm.	170 cm.	150cm.	70 cm.	170 cm.	150cm.	120cm.	100cm.	50cm
Dec. 1968													
2 3 4 5		120 117 113 108 106	180 183 143 77 166	38 35 32 33 34	52 50 42 38 44	135 139 141 107 103	149 150 155 112 112	163 172 74 74 101	44 44 37 25 34	50 49 6 8	286 286 297 220 220	110 110 110 55 55	264 242 242 264 242
67.690		102 100 93 92 87	130 158 168 180 196	37 40 39 36 38	49 50 44 18	111 117 116 125 126	123 131 132 137 140	22 142 1÷2 163 167	35 37 34 39 38	35 49 53 52 62	253 242 220 286 286	66 88 110 132 110	242 242 209 242 242
11 12 13 14 15		60 58 55 54	201 210 212 230 226	37 38 35 38 37	48 48 44 49 44	134 137 134 143 144	146 151 151 156 157	183 187 176 196 205	40 39 39 39 39	60 66 66 76 66	297 308 286 308 308	132 132 154 154 165	242 254 242 264 264
16 17 18 19 20 21 22 23 24	A LE DUMAN	5° 49	229 237	37 37	43 44	144	159 163	209 217	35 37	67 74	32°.	165 176	264 264
25 26 27 28 29 30													

APPRIDIE 9

MAIRE GROP AT MISA AND THE DETERMINATION OF ARRODYNAMIC RESISTANCE F

Theory

(a) Characteristics of wind profile

Since the atmosphere is viscous, the mean windspeed \overline{U} is not constant, but increases with height as
a result of the surface shearing stress γ_0 dynes/cm²,
which is the force of retardation per unit surface area.

The magnitude of γ_0 depends on geostrophic windspeed, surface roughness and buoyancy; and for laminar flow in which shearing stresses are due entirely to molecular vibrations, the relationship between shearing stress and windspeed gradient is linear (Sutton, 1955).

In the absence of buoyancy,

where wis the dynamic viscosity (gm.cm lee lated to kinematic viscosity y (cm leec l) by

the frictional velocity Un is hence defined by

$$\mathbf{U}_{+}^{2} = \frac{\mathcal{L}}{c} = \mathcal{V} \frac{\partial \mathbf{\overline{U}}}{\partial \mathbf{z}} \qquad \dots (5)$$

and introducing momentum eddy diffusivity K,

$$U_{+}^{2} = (K_{m} + \gamma) \frac{\partial \overline{U}}{\partial z} \sim K_{m} \frac{\partial \overline{U}}{\partial z} \qquad \dots (4)$$

This applies strictly in the surface boundary layer defined as the region where T does not vary by more than 5%. If we now define the rate of viscous dissipation per unit mass of air (cm²sec⁻⁵),

$$ke = \frac{v_{*}^{3}}{s} \qquad(5)$$

where k is von Karnan's constant.

Taylor (1952) observed that the shear-flow energy production P is related to U by

$$P = \frac{\partial}{\partial z} (\overline{v} \, \overline{v}_{\underline{v}}^2) \qquad (6)$$

and balancing production (6) with dissipation (5) gives

$$\overline{U}^{n} = s - \frac{\overline{U}}{s} \qquad \dots (7)$$

If we now make $\overline{\mathbf{v}} = 0$ at $\mathbf{s} = \mathbf{s}_0$, the windspeed profile over a uniform and extensive surface may be represented by a logarithmic function of the form

$$\overline{V} = \frac{V^n}{k} \ln \frac{s}{s} \qquad \dots (s)$$

where s is roughness length, a characteristic of the surface.

Over tall vegetation, a similar equation is applicable but in a modified form:

$$\overline{v} = \frac{\overline{v}}{k} \ln \frac{s-d}{s} \qquad \dots (9)$$

where d is sero plane displacement.

(b) Derivation of serotypanic resistants from the learnithmic wind profile

The aerodynamic resistance r is defined as the time in which I cm³ of air exchanges heat or water vapour with I cm² of surface. It can therefore be shown (Monteith, 1965) that if E is the evaporation rate, % the psychrometric constant, (c specific heat of air and e and e are vapour pressures at height s and at the evaporating surface,

$$\frac{\mathbf{E}(\mathbf{r})}{\mathbf{r}} = \frac{\mathbf{e}_{\mathbf{r}} - \mathbf{e}}{\mathbf{r}} \qquad \dots \dots (10)$$

If also the profiles of e and \overline{v} are the same shape above a uniform crop, the graph of e plotted against \overline{v} is a straight line intercepting the axis $\overline{v} = 0$ at $e = e_o$. Penman and Long (1960) showed that under these conditions the latent heat of evaporation is given by

$$\lambda = \frac{\chi^2 \ell \circ (\circ_0 - \circ) \mathbf{U}}{\sqrt{\frac{(\ln(s - d))^2}{2}}} \qquad \dots (11)$$

Equations (10) and (11) then give

$$\mathbf{r}_{\mathbf{a}} = \frac{\left\{\frac{\ln(\mathbf{z} - \mathbf{d})}{\mathbf{z}_{\mathbf{0}}}\right\}^{2}}{k^{2}\mathbf{U}} \qquad \dots (12)$$

Site of measurements and apparatus

The measurements were made over a maise crop at

Mrea Irrigation Scheme, altitude 1,280 metres, two weeks

after silking. Six small sensitive cup anemometers

(G. F. Cassella & Go.) mounted at 173, 275, 347, 402

and 487 on above the ground were set up on a mast at

a point 20 m from the downwind edge of the field.

The maise was planted 20 on apart on north-south rows 100 on apart. The direction of the wind was E to SE, i.e. mostly at right angles to the maise rows. Anemometer cup rotations were recorded as electrical pulses from which hourly averages of windspeeds were obtained, using calibration curves supplied by the manufacturers. After the experiment, all the anem-

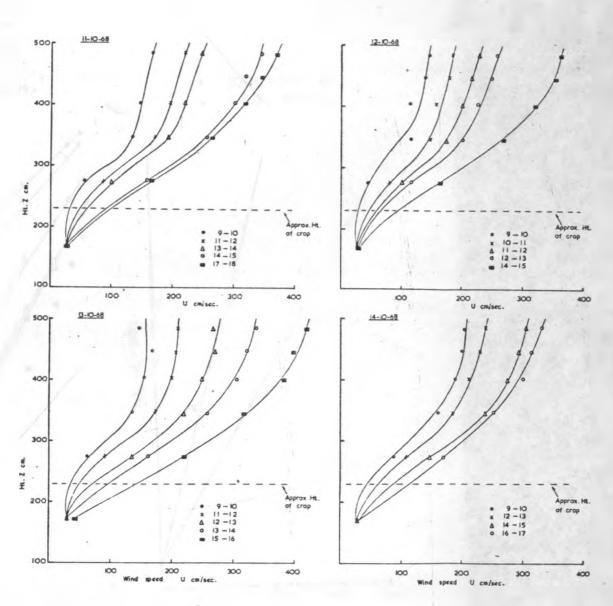
ometers were mounted on a horisontal bar and compared against one another. They all yielded the same windspeed -2% and this was taken into account when constructing the actual wind profiles.

Regults

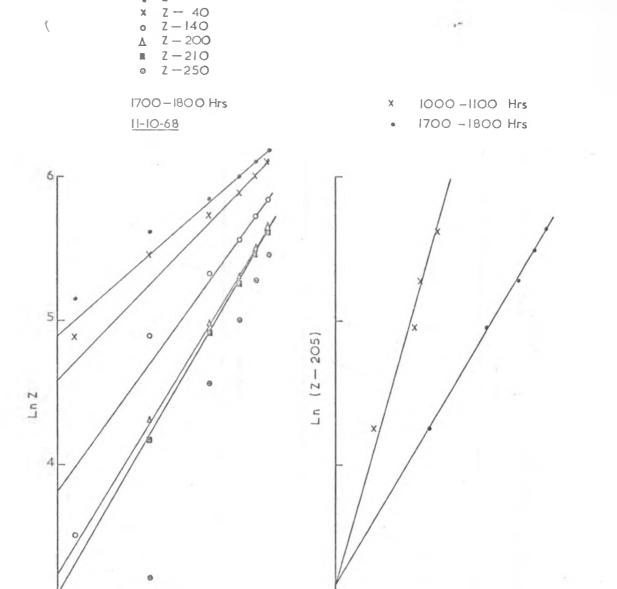
The representative hourly wind profiles, some of which are shown in Fig. 9:1 for clarity, resemble those obtained by Stoller and Lemon (1965), above and within a maise crop, but as in the wheat crop (Perman and Leng, 1960) there is no evidence of increase in windspeed above the crop, and the canopy tends to seal up. This explains the large value of zero plane displacement (205 cm) found necessary in adjusting the windspeed profiles to fit the logarithmic function of equation (9)(Fig. 9:2). The roughness length so was determined from the adjusted logarithmic windspeed profiles above the crop in Fig. 9:5; so was virtually constant at 24.0 - 0.5 cm on 11th and 12th October, but decreased occasionally on 15th and 14th October,

to a minimum of 15.7 cm. The maximum value of s is eless to 10% of average crop height (Slatyer and Mollroy, 1961; Munn, 1966).

Hourly values of r showed a similar trend each day, decreasing gradually from 0.12 - 0.18 at 0900-1000 hours to 0.05 - 0.05 at 1700-1800 hours. These values of r are however much smaller than the minimum values 0.2 - 0.5 quoted by Monteith (1965), and approach 0.05 for a pine forest (Sseics, Endrédi and Tajehman, 1969).



APPENDIX 9 - Fig. 9:1 Windspeed profiles above maize crop at Mwea

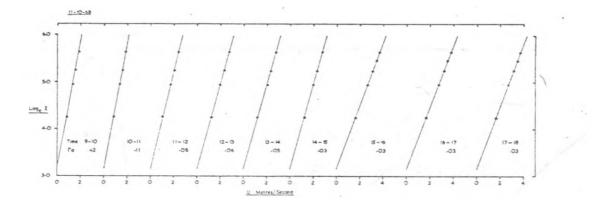


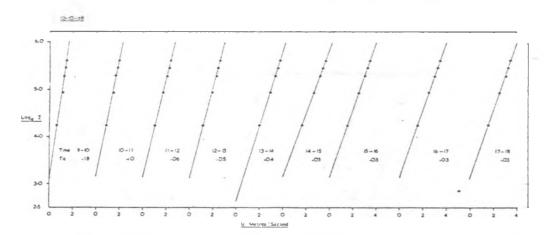
APPENDIX 9 - Fig. 9:2 Fitting logarithmic function on windspeed profiles above maize crop at Mwea

U cm/sec.

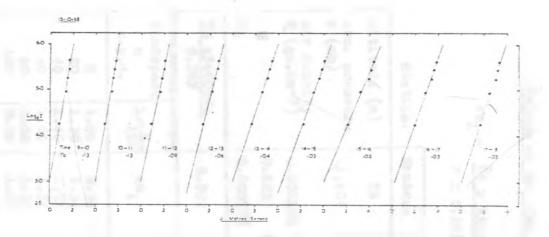
Wind speed

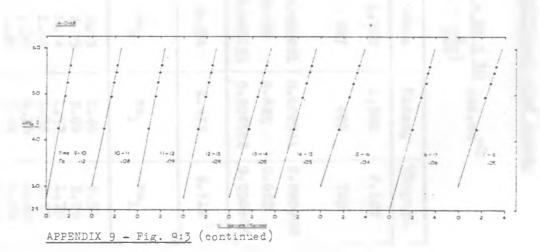
3 L





 $\frac{\text{APPENDIX 9-Fig. 9:3}}{\text{of } r_{\text{a}} \text{ above maize crop at Mwea}} \text{Logarithmic windspeed profiles and determination}$





APPENDIX 10

Values of ra for different wind speeds

 $ra_1 = \frac{e_{\epsilon} \times 3600 \times 24 \times 10}{P \times 0.26(1 + \frac{u}{100})}$ esc. cm⁻¹

Stat	ions	Monbasa	Mwea	Kedong	Muguga	
Altitude	(m)	55	1,280	1,900	2,100	
Atm. pressure P (mb)		1,004	867	803	783	
Air density e (gm.om-3)		0.001163	0.001021	0.000961	0.000952	
E		0.622	0.622	0.622	0.622	
e _e		0.000723	0.000635	0.000598	0.000592	
CE x3600x24 0,267		2.392	2.434	2.475	2.512	
Windspeed u 100		ral	ral	rel	ra _l	
0	1.00	2.39	2.43	2.48	2.51	
20 1.20		1.99	2.03	2.06	2.09	
40	1.40	1.71	1.74	1.77	1.79	
60	1.60	1.50	1.52	1.55	1.57	
80	1.80	1.33	1.35	1.38	1.40	
100	2.00	1.20	1.22	1.24	1.26	
120	2.20	1.09	1.11	1.12	1.14	
140	2.40	1.00	1.01	1.05	1.05	
160	2.60	0.92	0.94	0.95	0.97	
180	2.80	0.85	0.87	0.88	0.90	

APPREDIX 11

WATER INTEROPPED BY MALE PLANTS

Method

Seventeen maise plants of different sizes were uprooted from a field erop and quickly transferred to the laboratory. The roots and first 15 cm of the stem for each plant were then inserted in a flat bottomed container and dry soil added to keep the plants upright. An open ended plastic tube was slipped over the container and tied tightly around the stem to prevent water entry into the container.

Each plant was then weighed, sprayed thoroughly with water, and shaken to remove excess water before being reweighed. This procedure was repeated until a constant maximum weight was achieved.

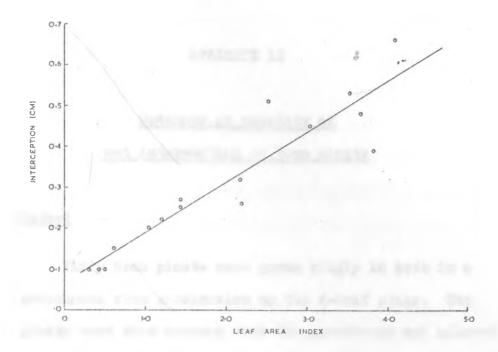
Regul to

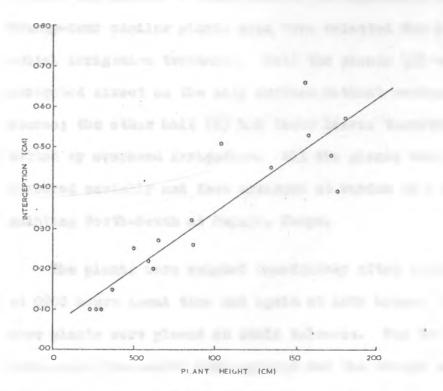
The results are pletted in Fig. 11:1 which also features the straight lines of best fit calculated from linear regression of interception on either leaf area index or height (H) of the plants. The two equations obtained are

- 1. Interception (em) = 0.12 L.A.I. + 0.07, r² = 0.89
- 2. Interception (cm) = 0.005E (cm) + 0.06, r² = 0.82.

Comment

The y-intercepts shown in the above equations have no meaning other than that L.A.I. and H were underestimated.





APPENDIX 11 - Fig. 11:1 Net interception of water by maize plants of different sizes as functions of leaf area index and plant height

Retimate of quantity of net interception by pean plants

Me thed

Fifty bean plants were grown singly in pote in a greenhouse from germination to the 4-leaf stage. The plants were then removed from the greenhouse and allowed to grow for another 2 weeks before the experiment started. Twenty-four similar plants were then selected for differential irrigation treatment. Half the plants (I) were irrigated direct on the soil surface without wetting the leaves; the other half (I) had their leaves thoroughly wetted by overhead irrigation. All the plants were mumbered serially and then arranged at random in a row pointing North-South at Muguga, Kenya.

The plants were weighed immediately after irrigation at 0900 hours local time and again at 1800 hours. Two more plants were placed on scale balances. The two irrigation tresiments were applied and the change in weight of each plant was recorded once every hour.

Results

Results of these experiments are shown in Appendices

12a and 12b. The experiment with two bean plants was

repeated several times, interchanging the plants between

the two irrigation treatments, but the results were

similar.

Evaporation from the plants with wetted leaves was eignificantly higher than water loss from plants irrigated direct on the soil surface, but the difference was not significant when total water loss exceeded 50 gm/plant/day.

Almost all the differences in water loss occurred in the first 1-2 hours.

APPENDIX 12a

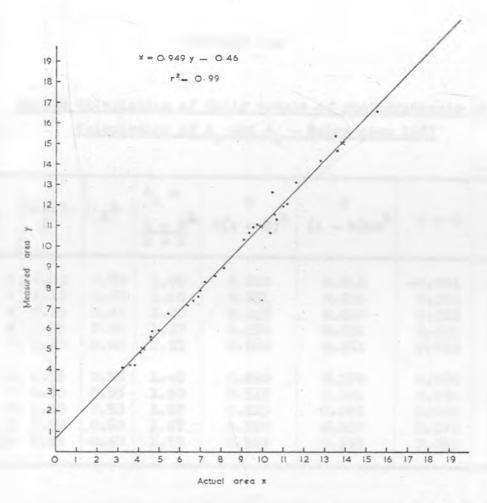
Evaporation (gm/day) from potted bean plants irrigated on sail (X) and on leaves (Y)

Plant No.	26	/8/68	28/	8/68	29/	8/68	30/	8/58	2/	9/68	3/	9/68
riant No.	Х	Y	Х	Y	Х	Y	I	Y	I	I	Y	I
1	33.3		34.5		68.9		76.6		91.6		86.8	
2		31.2		33.4		55.2		57.8	1 32.0	82.9	1	78.
3		42.5		33.6		51.2		43.8		65.0		58.
4	32.6		22.9		61.1		46.1		91.6		85.2	
5 6	43.4		26.7		50.2		48.1		91-1		89.4	
	77.0	46.8		35 . 8		72-2		59.6		82.7		89.
7	37.0 32.1		23.0		55.6		48.5		79.5		79.5	
10	25.1	43.3	23.5	35.5	60.6	55.5	48.6	*50.7	62.1	50.0	92.6	
11		56.6		40.8		66.8		159.3 69.0		58.2 78.8		78. 80.
12	36.5	70.0	29.1	40.5	53.8	00.0	51.6	09.0	91.5		90.8	
13	35.1		26.9		46.2		51.6		84.2		84.1	
14	1	45.0	2017	32.2	7012	51.2	1 72.0	55.4	04.2	57.1	04.1	79.
15	1	30.5		25.9		42.0		46.9		84.6		80.
16	1	28.7		26.7		58.8		40.8		63 - 6		78.
17		39.7		33.0		50.5		45 - 2		82.5		77.
18	1	33.5		26.2		54.2		42.4		70.0		80.
19		40-4		28.9		64.0		46.4		84.1		78.
20	29.2		24.0		70.8		45.0		85.9		64.0	
21 22	35.1		24.6		49-6		48.6		70.3		51.1	
23	42.8	47.0	22.5	05 5	76.7		48.9		69.4		86.6	
24		43.9 32.3		25.5 34.7		49.8		40:2		68.6		77 -
		72.7		34+1		82.5		49.9		66.0		86.
Mean of 1st 6	32.7	44.2	25.2	35.2	57.0	58.7	52.4	57.5	81.8	70.8	84.9	77.
" " 2nd 6	34.6	36.7	25.6	29.0	59.3	57.2	47.8	45.2	77.5	74.3	75.8	80.
Wean Y-X											1	1
lst 6	11.	. 5	10.	. 0	1.	. 7	5.	1	11	. 0	7	. 4
2nd 6	2.	.1	3.	. 4	-2.	.1	-2.	6	3	. 2	4	- 4
	13.	. 6	13.	. 4	-0.	. 4	2.	.5	14	. 2	3	. 0
Significance	S		S		107	S.	N	s.	9 0	t 2%	W	.5.

APPENDIX 12b

Evaporation from two bean plants: watered on the soil (X) and on the leaves (X) on 1/9/68

Time	Evaporation ga/hr						
hre	x	r					
6 - 9	14.5	63.8					
9 - 10	15.7	25.2					
10 - 11	5.8	12.3					
11 - 12	5.0	7.7					
12 - 13	11.2	8.7					
13 - 14	9.5	12.6					
14 - 15	12.4	10.4					
15 - 16	11.9	12.5					
16 - 17	9.3	11.4					
17 - 18	5.1	4.5					
Total	100.4	169.1					



Appendix 13: calibration of the products of maximum length and width against actual leaf area for beans.

All units are eq. ins.

APPENDIX 14a

Sample calculation of daily totals of photosynthesis (P Galculation of A, and A, - Maize crop 1967

	C .
1 1	1 1
m	70 1.
7	1 19
r r	r r
ri	62 1.
1.33	1 69
7	7
1.50	55 1.
1.57	53 1.
1.67	50 1
1.73	19 7

Sample relculation of daily : tosynthesis (P)

Calculation of Post and hence P Maize Crop 1967

$$I^{a} = \frac{1R}{2n} \text{ cal}$$

$$a = 0.25$$

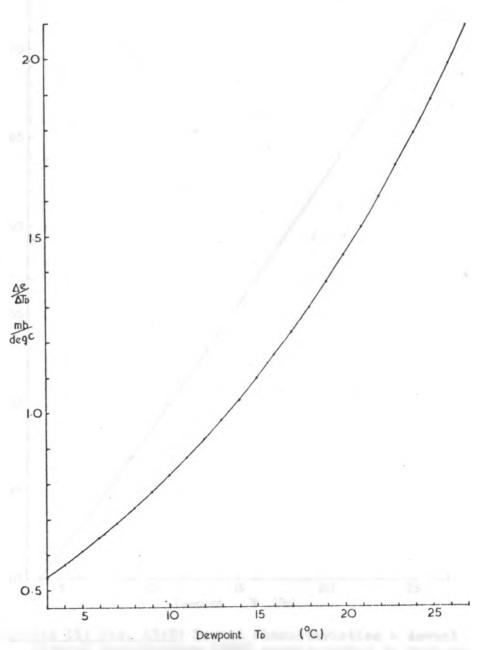
$$\Upsilon = 0.055$$

$$+ P_{1} = \frac{1}{2} \left\{ o \left[1 - f(\eta_{0}) \right] + A_{1} \left[1 - f(\eta_{1}) \right] \right\}$$

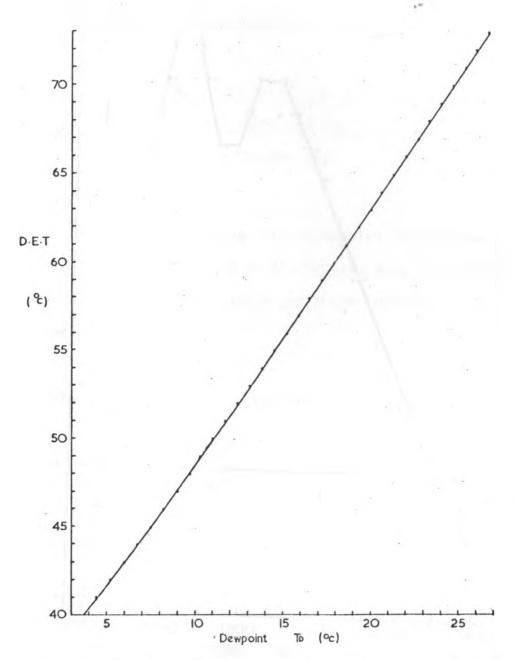
Date	h min.	R Ly/day	I* Ly/min.	4,	1 - r(1 ₀)	Ao	- r(2,)]	h a x 60	Fo gm.m-2day-1	144	11	1 - (41)	Al	$\delta(D_{i} + \tau(\tau_{i}))$	P ₁	P ₀ + P ₁
OCTORER																
5	726 726 726 727	573.8 680.0 656.4 591.5	1.24 1.47 1.42 1.28 1.49	0.32 0.37 0.36 0.32 0.39	0.51 0.54 0.53 0.51 0.54	1.00 1.10 1.20 1.27 1.33	0.510 0.594 0.636 0.648 0.718	48.4 48.4 48.4 48.4 48.5	. 4.68 28.75 50.78 31.35 34.76	0.07 0.08 0.08 0.07 0.08	-0.80 -0.78 -0.78 -0.80 -0.78	0.07 0.08 0.08 0.07 0.08	-0.019 0.029 0.057 0.071 0.090	-0.001 0.002 0.005 0.005 0.007	-0.06 0.11 0.22 0.24 0.35	24.62 28.86 31.00 31.59 35.11
6	727 727 727 727 727	508.9 612.1 668.2 615.1 638.7	1.10 1.32 1.44 1.33 1.38	0.27 0.34 0.36 0.34 0.36	0.48 0.52 0.53 0.52 0.53	1.40 1.50 1.57 1.67 1.73	0.672 0.780 0.832 0.808 0.917	48.5 48.5 48.5 48.5 48.5	52-52 37-75 40-27 42-03 44-38	0.06 0.07 0.08 0.07 0.08	-0.83 -0.80 -0.78 -0.80 -0.78	0.06 0.07 0.08 0.07 0.08	0.114 0.162 0.205 0.295 0.329	0.007 0.011 0.016 0.021 0.026	0.33 0.55 0.79 1.00 1.27	32.85 38.30 41.06 43.03 45.65

Computation of vapour pressure gradient (de), from mean dewpoint (Tp) and the differences in dewcel element temperatures (DET):

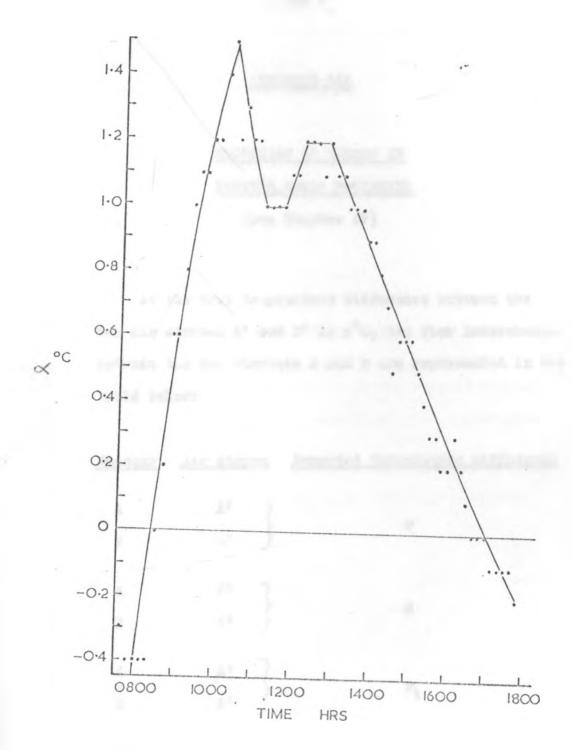
By entering Figs. 15:1 and 15:2 with the mean dewpoing temp. $T_{\rm p}$, values of $\frac{\Delta c}{\Delta T}$ and $\frac{\Delta DET}{\Delta T}$ were obtained. A table of values of $T_{\rm p}$ was then constructed.



Appendix 15: Fig. 15:1 The ratio of vapour pressure to dewpoint gradients As at different dewpoint temperatures.



Appendix 15: Fig. 15:2: Dewcel characteristics - dewcel element temperatures (DET) corresponding to various despoint temperatures (T_D) .



APPENDIX 16a: Variation of absolute temperature difference
between the two thermocouple champers in Bowen's
Ratio equipment

APPENDIX 16b

BOTTH'S PARIO EQUIPMENT.

(see Chapter IV)

If the true temperature difference between the two air streams A' and B' is x'C, two flow interchanges between the two chambers A and B are represented in the table below:

Chamber	Air atream	Recorded temperature difference
A	A' }	*
A	V ₀ }	
A	A* }	r ₁

For the three situations,

$$\mathbf{r} = \beta + \alpha_{\mathbf{r}} + \kappa \tag{1}$$

$$g = \beta + \alpha + \alpha$$
(2)

$$\mathbf{r}_1 = \beta + \alpha_{\mathbf{r}_1} + \mathbf{x} \tag{5}$$

from which it can be shown that

$$\frac{1}{8}\left[\frac{x_{-S}}{2}+\frac{x_{1}-8}{2}\right]=x_{1}+\frac{1}{2}\left[(x_{x}+x_{x_{1}})-2x_{y}\right] \qquad (4)$$

and the same argument applies in the case of humidity measured as difference in dewcel element temperature (D.E.T.).

If & changes linearly with time (see Appendix 16a).

$$\alpha_{\underline{z}} - \alpha_{\underline{z}} = \alpha_{\underline{z}} - \alpha_{\underline{z}} \qquad (5)$$

giving

$$\dot{\mathbf{g}}(\mathbf{x}_{\mathbf{r}} + \mathbf{x}_{\mathbf{r}_{1}}) = 2\mathbf{x} \qquad \dots (6)$$

and therefore

$$\times = \frac{1}{2} \left\{ \frac{\mathbf{r} - \mathbf{g}}{2} + \frac{\mathbf{r}_1 - \mathbf{g}}{2} \right\} \qquad \dots (7)$$

This procedure does therefore considerably reduce the error in the determination of , the error increasing slowly as the increase in with time departs from linearity.

Samule calculation of Bowen's Ratio

STATION: Mwea

SURFACE: Bean crop

SURFACE CONDITION: Pods filling, complete ground cover

DATE: 7th May, 1968

NOZZLE HEIGHT ABOVE CROP: Upper: 70 cm.

Lower: 20 cm.

$$A = \frac{1}{2} \left\{ \frac{(g-r)}{2} + \frac{g_1 - r}{2} \right\}$$

$$A = \frac{1}{2} \left\{ \frac{(g-r)}{2} + \frac{g_1 - r}{2} \right\} \qquad A_1 = \frac{1}{2} \left\{ \frac{(r-g)}{2} + \left(\frac{r_1 - g}{2}\right) \right\}$$

Flow	10-minute period	d(D	.E.T.) oc	•	4-			dT °C.		am	B=	
	starting	chart reading	½(g-r)	A	de d(D.E.T.)	de mb.	chart reading	½(r-g)	A ₁ =dT	dT de	₹ dT de	1+B
g	0900	4.35	-0.01	0.13	0.892	0.12	-0.15	0.01	0.03	0.25	0.15	1.15
r	0910	4.37	0.28	0.27	0.892	0.24	-0.14	0.04	0.05	0.21	0.13	1.13
g	0920	4.92	0.27	0.32	0.902	0.29	-0.21	0.05	0.06	0.21	0.13	1.13
r	0930	4.38	0.36	0.29	0.888	0.26	-0.12	0.07	0.04	0.15	0.09	1.09
g	0940	5.09	0.21	0.25	0.888	0.22	-0.25	0.01	0.02	0.09	0.05	1.05
r	0950	4.67	0.28	0.29	0.888	0.26	-0.24	0.02	0.02	0.08	0.05	1.05
										hourl;	y mean:	1.10
g	1000	5.22	0.29	0.31	0.888	0.28	-0.27	0.02	0.02	0.07	0-04	1.04
r	1010	4.65	0.32	0.29	0.897	0.26	-0.24	0.02	0.01	0.04	0.02	1.02
g	1020	5.28	0.25	0.27	0.892	0.24	-0.24	0.00	0.01	0.04	0.02	1.02
r	1030	4.78	0.29	0.30	0.892	0.27	-0.23	0.01	0.02	0.07	0.04	1.04
g	1040	5.35	0.30	0.29	0.902	0.26	-0.20	0.02	-0.01	-0.04	-0.02	0.98
r	1050	4.75	0.27	0.28	0.892	0.25	-0.16	-0.03	-0.01	-0.04	-0.02	0.98
										hourl	y mean:	1.01
g	1100	5.28	0.29				-0.10	0.01				
r	1110	4.70					-0.08					

WHITERSITY OF NAIROS.

Twited (days from gyrale-

htian]

21-0) 23-76 35-81

61-50 50-55 50-85

78-84 81-288 100-119

Energy storage in photosynthesis

(a) Maize crop 1967

Period (days from germination)	Increase in dry matter	A Bnergy equivalent K.cal./m ²	Total shortwave radiation K.cal./m ²	B Total radiation 0.4 - 0.7 K.cal./=2	A B
0-17 17-32 32-46 46-52 52-59 59-73 -73-81 81-94 94-102 102-111	15 91 230 321 232 537 134 336 216	62 374 955 1,332 963 2,229 556 1,394 896 564	90,080 92,820 83,830 26,700 36,650 71,320 47,780 64,310 50,490 56,800	42,338 43,625 39,400 12,549 17,226 33,520 22,457 30,226 23,730 26,696	0.002 0.008 0.024 0.106 0.056 0.067 0.025 0.046 0.038 0.021
	(b) <u>B</u>	ean crop 196	8		5 1 1
5-9 9-12 12-16 16-23 23-27 27-43 43-48 48-54 54-58 58-65	4 3 23 8 14 162 60 20 16 85	17 12 95 33 58 672 249 83 66	24,370 20,130 18,780 34,640 24,930 77,410 24,430 27,320 21,860 39,170	11,454 9,461 8,827 16,281 11,717 36,383 11,482 12,840 10,274 18,410	0.001 0.001 0.011 0.002 0.005 0.018 0.022 0.006 0.006 0.019

250 -

200

World and

0.107

C-DOS

APPENDIX /8 (continued)

(c) Maise crop 1968

Period (days from germin- ation)	✓I Increase in dry □ natter Δ H gn/m²				Energy equivalent (A) K.cal./m ²			B Total radiation	<u>Å</u> B			
	Wet	Medium	Dry	Wet	Medium	Dry	radiation K.cal./m ²	0.4 - 0.7 k K.cal./m ²	Wet	Medium	Dry	
310	Year											
24-29	20	7	8	83	29	33	11,990	5,635	0.015	0.005	0.006	
29-36	66	52	20	274	216	83	24,710	11,614	0.024	0.019	0.007	
336-44	67	61	89	278	253	559	51,200	24,064	0.012	0.011	0.015	
44-50	288	256	132	1,195	1,062	548	31,740	14,918	0.080	0.071	0.037	
50-58	59	109	115	245	452	477	43,750	20.563	0.012	0.022	0.023	
58-65	306	114	49	1,270	473	203	44,660	20,990	0.061	0.023	0.010	
65-72	380	342	475	1,577	1,419	1,971	45,310	21,296	0.074	0.067	0.093	
72-84	92	272	120	382	1,129	498	75,450	35,462	0.011	0.032	0.014	
84-100	496	465	616	2,058	1,930	2,556	81,180	38,155	0.054	0.051	0.067	
100 - 119 119 - 130	640	321 305	201 354	32,050	1,332	834	97,270 38,880	45,717 18,274	}0.043	0.029	0.018	
										2		